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Permafrost - Current and future challenges to study methanotrophy in permafrost affected tundra and wetlands

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Summary
Arctic permafrost affected wetlands and tundra are known to be sources of the greenhouse gas methane (Christensen et al. 1995, 2000, Wagner et al. 2003, Kutzback et al. 2004, Sachs et al. 2008, Wille et al. 2008). As a consequence of the rising surface temperatures and the degradation of permafrost, methane emissions from these Arctic habitats are predicted to increase in the future. Evolution of methane fluxes across the subsurface/atmosphere boundary will thereby strongly depend on the activity of obligately aerobic methane oxidizing Proteobacteria which are known to be abundant and to significantly reduce methane emissions in permafrost affected soils despite seasonal freezing and thawing and despite the mostly extremely low temperatures. Based on their function as the major (known) sink for methane, research is needed on quantifying in-situ activities of aerobic methane oxidizing Proteobacteria and on understanding their potential to adapt to the currently changing environment. Also, possible unknown sinks for methane in these regions as for example known for marine methane rich habitats need to be explored.
Introduction
Methanotrophy, the consumption of methane, also occurs under the very harsh environmental conditions of cold environments (review by Trotsenko and Khmelenina 2005). Viable methane oxidizers were even detected in deep Siberian permafrost sediments with ages of 1000-100,000 years (Khmelenina et al. 2001). In permafrost affected tundra, methane is constantly bypassing the so-called active layer during the period of Arctic summer (section B15.d.ii.). Wherever oxygen is present in these environments (elevated and dry sites, photosynthetically active polygonal ponds, rhizosphere), methane is oxidized to carbon dioxide. Though not yet quantified in-situ for permafrost affected soils, between 76 % and up to more than 90 % of the methane produced in wetlands gets oxidized before reaching the atmosphere (Roslev and King 1996, Le Mer and Roger 2001). During the last years, the group of microorganisms capable to consume methane (methanotrophs) has been observed to be very diverse comprising obligate aerobic members of the phyla Proteobacteria (Bowman 1999, 2006), and Verrucomicrobiaea (Dunfield et al. 2007; Pol et al. 2007), as well as anaerobically methane oxidizing Archaea in marine habitats (e.g. Boetius et al. 2000), and bacteria of a yet unknown phylum carrying out methane oxidation in the presence of very high nitrate and methane concentration in freshwater habitats (Raghoebarsing et al. 2006). In permafrost affected tundra, however, only aerobic methane oxidizers of the Proteobacteria phylum, described in detail elsewhere in this book (section C20c.i.) have been observed so far. Based on their pronounced distribution and function as the primary sink for methane in Arctic permafrost affected tundra (Wagner 2008), they are of major importance for the greenhouse gas (GHG) budget of these often methane rich environments.

Ecology of aerobic methanotrophic Proteobacteria in permafrost affected soils
last two which are also known as type X MOB. The families Methylocystaceae, and Beijerinckiaceae include the genera Methylosinus, Methylocystis, Methylocella, and Methylocapsa. Members of the Methylocystaceae and Beijerinckiaceae are termed type II MOB and belong to the alpha-subdivision of the Proteobacteria phylum.

Almost 200 isolates of MOB are known (http://www.ncbi.nlm.nih.gov/Taxonomy/) suggesting that MOB are ubiquitous in many environments. Some psychrophilic and psychrotrophic strains as well as several uncultivated methanotrophs were obtained from Arctic permafrost habitats (Figure 1 of this chapter). *Methylobacter psychrophilus* (type I) is a true psychrophilic organism with optimum growth between 3.5 and 10 °C that was isolated from Siberian tundra (Omelchenko et al. 1996). *Methylobacter tundripaludum* (type I) with an optimum growth at 23 °C and *Methylocystis rosea* (type II) with an optimum growth at 27 °C are psychrotrophs isolated from Arctic wetland soils in Svalbard (Wartiainen et al. 2006a, b), and *Methylocella tundrae* (type II) with an optimum growth between 15
and 25 °C, as well psychrotrophic, was isolated from *Sphagnum* tundra peatlands (Dedysh et al. 2004). Additional psychrotrophic methanotrophs are *Methylosphaera hansonii* (type I) isolated from Antarctic, marine salinity, meromictic lakes (Bowman et al. 1997), and *Methylomonas scandinavica* (type I), isolated from deep igneous rock ground water (Kaluzhnaya et al. 1999). In general, type I methanotrophs were reported to dominate over the type II group in Arctic permafrost affected soils (Wartiainen et al. 2003, Wagner et al. 2005, Liebner and Wagner 2007).

Methane oxidizing *Proteobacteria* were shown to be highly abundant in permafrost soils of the Lena Delta, Siberia, with cell numbers ranging between $3 \times 10^6$ and $1 \times 10^8$ cells g$^{-1}$ soil and contributing up to 10 % to the total number of microbial cells (Liebner and Wagner 2007). In the same area, specific clusters of methane oxidizing *Proteobacteria* closely related to *Methylobacter psychrophilus* and to *Methylobacter tundripaludum* were detected indicating a micro-diverse community on the species level (Liebner et al. in press). Highly divergent functional gene sequences of methanotrophs were found in soils of the high Canadian Arctic (Pacheco-Oliver et al. 2002). In contrast, the diversity of methane oxidizing *Proteobacteria* in wetland samples from Svalbard was observed to be restricted to only two genera (Wartiainen et al. 2003), whereas most methanotrophic *Proteobacteria* were detected in a Russian sub-arctic tundra (Kaluzhnaya et al. 2002). A recent study observed a shift between a mesophilic methanotrophic community near the surface and a psychrophilic methanotrophic community near the permafrost table of Siberian permafrost soils (Liebner and Wagner 2007). This indicates that depending on the environmental conditions both mesophilic as well as psychrophilic methanotrophs are active in Siberian permafrost soils. Still, it remains to be proven which group is responsible for methane oxidation at low and subzero temperatures in deep permafrost sediments (Trotsenko and Khmelenina 2005).

**Methane oxidation in the context of thawing permafrost**

Model calculations suggest that methane currently emitted from Arctic permafrost environments may enhance the greenhouse effect with a portion of approximately 20 % (Wuebbles and Hayhoe 2002). Arctic surface temperatures on average increased to a greater extent than those of the rest of the earth (IPCC 2001, 2007), causing a particular susceptibility of Arctic permafrost to degradation (Nelson et al. 2001). Global warming is assumed to degrade 25 % of the total permafrost area by 2100 (Anisimov et al. 1999). Serious concerns are
thus associated with the potential impact that thawing permafrost may have on the global climate system through release of greenhouse gases (Friborg et al. 2003, Christensen et al. 2004, Wagner et al. 2007) highlighting the importance of the current and of the future methane oxidation capacity in these regions. In models on modern methane emissions from Arctic wetlands, methane production and oxidation rates are primarily determined as functions of substrate availability, substrate concentration, and temperature as well as indirectly of water table and thaw depth (Walter et al. 2001, Zhuang et al. 2004, Anisimov 2007). Changes of these parameters will consequently lead to short-term alterations of methane production and methane oxidation rates. Whether, however, the currently observed climate change will effectively alter modern methane fluxes from Arctic permafrost affected wetlands will particularly depend on its long-term impact on the methane cycling communities and on their ability to adapt to the new environmental conditions. This ability is very likely dependant on the level of specialisation and diversity of the indigenous microbial communities. It was, for example, observed that an increase of temperature and precipitation altered the community structure and relative abundance of methane oxidizers in rice, forest and grassland soils (Horz et al. 2005, Mohanty et al. 2007). Also, the overall relative abundance and diversity of methanogenic archaea in a high Arctic peat from Spitsbergen increased with increasing temperature, in accordance with a strong stimulation of methane production rates (Høj et al. 2008). In contrast, the population structure of methanogenic archaea in a permafrost affected peat in Siberia remained constant over a wide temperature range (Metje and Frenzel 2007). There is, however, a lack of experimental research investigating the long-term effect of simulated climate change on the methane cycling communities in permafrost affected soils. Also, an account of the entire plant-microbe-animal-system and the interactions between metabolic networks, which are important for the production of methane (Panikov 1999), are missing in modern methane flux models.

Research needs
Though well known that methanotrophy by aerobic Proteobacteria functions as a major sink for methane in permafrost affected tundra and wetlands, experimental data on the actual amount of methane flux reductions by methanotrophs are completely lacking. In-situ rate measurements of methane oxidation activities such as push-pull-tests (e.g. Urmann et al. 2005, 2007, 2008) as well as below ground flux measurements would essentially improve our knowledge in this
context and would help to validate current methane flux models from Arctic wetlands. Also, as discussed in the introduction to this chapter, methanotrophy can be carried out by a variety of different bacterial phyla and even by members of the domain Archaea. So far the group of aerobic methanotrophic Proteobacteria is the only one known to be responsible for methane consumption in permafrost affected wetlands and tundra. Nevertheless, much more attention should be paid on exploring yet unknown communities possibly capable to consume methane in these habitats as well. These additional methane sinks might be of particular benefit considering how important the future methane oxidizing capacity will be in the scope of continuously degraded permafrost. We will finally have to investigate how methane consuming communities respond to (simulated or actual) global climate change. This includes a generally improved understanding on what determines the stability of methane oxidizing communities and what implications may arise from that knowledge to the future GHG budget of Arctic permafrost affected wetlands.

References


