

Seasonality of Spectral Albedo and Transmission of Sea Ice in the Transpolar Drift, Arctic Ocean 1632 1632 TARIN

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Introduction

The physical and optical properties of Both were measured continuously at high snow and sea ice in the Polar Regions spectral and temporal resolution during the control the amount of solar short-wave ra- transpolar drift of the schooner Tara diation, reflected at the surface, scattered through the Arctic Basin between April and and absorbed within snow and ice, and September 2007. In addition, a nearby ice transmitted into the ocean beneath.

Albedo and transmissivity of snow and sea ice strongly influence heat fluxes Simultaneous in-situ measurements of within the coupled atmosphere-ice-ocean snow and sea ice properties and meteorosystem, and by that the evolution of the logical observations complement the ausea ice

mass-balance buoy (IMB) was used for data analysis.

tonomous datasets [Gascard et al. 2008].

Results

850-

Results show significant seasonal changes and highlight key events during the transitions from spring to summer and summer to autumn.

Melt season began on 11 Jun and lasted 65 days until 15 Aug, including formation, Primary production of organisms and inevolution, and freeze up of melt ponds.

Net short-wave (330-920 nm) flux was 29 W/m² at the surface and 2.6 W/m² under the ice. Resulting potential melt of sea ice was 1 m.

Spectral composition of transmitted light changes significantly during summer, showing increased absorption between of minor importance. 400 and 600 nm. 92% of transmitted light were in the PAR range (400-700 nm).

creased mass of Dissolved Organic Matter caused most likely a mid-summer decrease of transmitted light. Algae growth on the under-ice sensor, and later grazing might have affected the result, too.

Spectral radiation measurements

Sea-ice mass balance was dominated by

surface ablation and bottom ablation was

Surface properties of snow and sea ice represent typical summer conditions: ablation of the entire snowpack, deteriorated sea-ice surface, and isothermal surface temperatures of 0°C.

Textural analyses show that the measurements were performed over and under multi-year sea ice with 88% columnar ice.

Figure 4: Spectral transmission at times of highest sun elevation









The drift of *Tara* (2006 to 2008)



Figure 1: Map of Tara drift during spectral radiation (main figure) and map of the entire voyage of Tara from Tiksi, Russia, to Longyearbyen, Svalbard (inlay). The blue line shows the drift track and red dots give daily positions of Tara at the time of highest sun elevation, green dots and text labels mark every 10th day for better orientation. Ad-ditionally, significant events are labeled. The Tara expedition was initiated as a non-scientific transpolar drift, emulating Nansen's Fram expedition from 1893 to 1896.

Intensive field work was carried out as part of the DAMOCLES project during the drift, most intensively during summer 2007, just before the minimum of Arctic seaice extent

Snow and sea-ice properties were measured in-situ weekly.

Optical measurements



Figure 2: Photograph of the station set-up (28 Apr 2007) showing both. spectral and broadband radiation sensors (photo: M. Nicolaus). The under-ice sensor is lowered at the end of the black cable behind the station and cannot be seen from the surface. The schooner Tara and the field camp are shown in the background. The small inset-picture shows the deployment of the under-ice sensor in its rack (photo: F. Latreille). Inial sea-ice and snow thickness was 2.0 and 0.1 m, respectively

Figure 3: Seasonal changes of surface conditions around the radiation station. Photographs show late winter conditions (18 May), early melt pond formation (24 Jun), further melt pond evolution (02, 16, and 30 Jul), and autumn freeze up (28 Aug). Note that the photo from 18 May was taken from the opposite side. Most pictures were taken by T. Palo (all photos © Tara expeditions). Spectra of each surface condition (day) are given in Figure 5, plotted in the same color

Spectral radiation was measured with 3 Trios Ramses radiometers (320-950 nm, 3.3-nm resolution).

Two sensors were installed above the surface and one was hanging under the sea ice. Broadband radiation was measured at the same rack (Fig. 2)

Spectral radiation was measured in 30-min intervals from 28 Apr to 28 Aug (transmission) and 05 Sep (albedo) 2007.

The station was maintained daily, but the spectral sensors did not need to be cleaned at all during the entire observation period.

18 May 2007

0.12



during the drift (Fig. 1). Left: Daily spectra, above: Selected dates of different surface conditions (Fig. 3). No transmission data are available after 28 Aug 2007, because the under-ice sensor had to be Date in 2007 (dd-mmm)

Characteristic events (see also arrows)

- Melt onset (drop of IR albedo, increase of transm.) 11 Jun: 16 Jul: First melt ponds observed 21 Jun:
- Water standing on surface, BB-albedo minimum 02 Jul:
- 12 Aua: 15 Aug:
- Drop of transmission, no change in albedo Re-increase of transmission, no change in albedo Begin of freeze-up (first new snow, albedo increase)



01-May 15-May 01-Jun 15-Jun 01-Jul 15-Jul 01-Aug 15-Aug 01-Sep Date in 2007 (dd-mmm)

02 Jul 2007

24 Jun 2007

during the drift (Fig. 1). Left: Daily spectra, below: Selected dates of different surface conditions (Fig. 3).



16 Jul 2007



Figure 5: Spectral albedo at times of highest sun elevation





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Conclusions

spectral albedo and transmission measurements on Arctic sea ice.

and sea ice properties.

We have gathered a unique data set of they were made prior to the Arctic sea ice continuous and high temporal-resolution extent minimum in autumn 2007.

Absolute values and timing of energy transfer through snow and sea ice into the Results show how coordinated scientific ocean are now available for biological programs during private expeditions can process studies. Light data are of special contribute to extensive data sets of snow interest, since light is usually the limiting factor for primary production.

Our observations have a large potential to Spectral radiation results show that this be used for upcoming model and remote method might also be used for other studsensing applications. Especially because ies on (semi-) autonomous platforms.

Snow and sea-ice conditions

Figure 6: Time series of (a) broadband albedo and transmission, albedo time series company direct broadband measurements [Vihma et al., 2008] and integration from spectral measure ments (Fig. 2), (b) surface temperature (from snow pits, IMB, and radiation measurements) and snow grain size, (c) sea-ice and total (snow+ice) thickness measured with IMB 2007C, installed close to the radiation station, and (d) snow ablation and thickness. Snow properties were measured weekly in snow pits and along stakes, both within 50 m from the radiation station



Broadband albedo from spectral measurements was on average 0.12 higher than from broadband measurements, as expected because of the limited spectral coverage, but both are consistent in relative changes.

Sea-ice thickness decreased by 0.6 m, with 0.5 m of the loss due to surface ablation. Snow thickness varied between 0 (from 21 Jun to 14 Aug) and 21 cm (2 Sep).

Surface temperature are consistent from snow pits and radiation measurements, mostly 0°C during summer.

Acknowledgement

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References

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Figure 7: Thinsection-photo from an ice core retrieved at the under-ice sensor site (25 Apr 07). Thin sections were photographed under transmitted light with crossed polariz ers (left) and without (right). Scale is in 0.05 m sections, total

core length was 2.06 m.

No.