Analysis of the hydrographic conditions and cyst beds in the San Jorge Gulf, Argentina, that favor dinoflagellate population development including toxigenic species and their toxins

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A B S T R A C T
The overlay of cooler nutrient enriched Beagle–Magellan water with warmer nutrient depleted shelf water and a strong stratification of the water column in the San Jorge Gulf region, Argentina, coincided with relatively high dinoflagellate abundances in April 2012, up to 34,000 cells L−1. This dinoflagellate proliferation was dominated by Ceratium spp., but environmental conditions also favored to a lesser amount the occurrence of toxigenic dinoflagellates, such as Alexandrium tamarense and Protoceratium reticulatum, whose toxins were hardly detected in any other areas along the expedition transect of the R/V Puerto Deseado between 38 and 56°S (Ushuaia–Mar del Plata) in March/April 2012. Generally vegetative cells of A. tamarense and P. reticulatum co-occurred with their respective phycotoxins in the water column and their cysts in the upper sediment layers. Two strains of A. tamarense were isolated from the bloom sample and morphologically characterized. Their PSP toxin profiles consisted of C1/2, gonyautoxins 1/4 and to a lesser amount of neosaxitoxin and confirmed earlier data from this region. The ratios between autotrophic picoplankton and heterotrophic bacteria were higher in shelf waters in the north than in Beagle–Magellan waters in the south of San Jorge Gulf.

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1. Introduction

The Southwest Atlantic is well known for its high primary production and therefore presents a very important region for fisheries and seafood production (Bertolotti et al., 1996; Lutz et al., 2010). Zones of highest productivity are upwelling areas along the South American Continental Shelf Edge (Garcia et al., 2008; Romero et al., 2006), but also frontal zones, where current systems with different water masses converge, such as the Brazil Current and the Malvinas/Falkland Current (Brandini et al., 2000; Piola et al., 2010, 2013). In addition, there are smaller local frontal systems in deeper bays and gulfs along the arid Patagonian coastline. These front systems result from isolated waters that are formed by interplay of high evaporation during summer, the shelf current systems and riverine inputs (Acha et al., 2004; Rivas and Pisoni, 2010). Current circulation suggests that tides, wind and bottom topography, for instance the ridge at the mouth of the San Jorge Gulf (SJG), are the main local forcing mechanisms in the Patagonian gulfs (Tonini et al., 2006; Palma et al., 2008; Matano et al., 2010). These hydrodynamic properties lead to distinct water masses characterized by high primary production. Among the phytoplankton species proliferating in these high productive areas there are also harmful algal taxa. Harmful algal blooms of the toxic marine dinoflagellates Alexandrium tamarense, Alexandrium catenella and Gymnodinium catenatum, which cause the neurotoxic paralytic shellfish poisoning (PSP), are a recurrent phenomenon along the Argentinean coast (Carreto et al., 1985, 1998a,b, 2008; Esteves et al., 1992; Santinelli et al., 2002). Toxic diarrhetic
shellfish poisoning (DSP) episodes caused by Dinophysis species have been also reported (Sar et al., 2010, 2012), as well as the production of domoic acid (DA) by Pseudo-nitzschia species under field and culture conditions (Montoya et al., 2006; Negri et al., 2004; Sastre et al., 2007). In contrast, little is known about the occurrence of pectenotoxins (PTxs) and yessotoxins (YTXs). The most recently observed shellfish poisoning syndrome is azaspiracid shellfish poisoning (AZP), first reported in Europe in 1996 and unexplored on the Argentinean shelf. Except for the amnesic shellfish poisoning (ASP), which is produced by several species of the diatom genus Pseudo-nitzschia, all other shellfish poisoning syndromes are caused by dinoflagellates.

Dinoflagellates are bilaterally eukaryotic single-celled organisms that in the most prevalent stage (haploid motile) can reproduce asexually by binary fission. Beyond this relatively simple mode of reproduction, many extant free-living species are known to have a more complex life cycle producing, at times, vegetative or resting cysts. Resting cysts involve a usually obligatory period of dormancy (in contrast to temporary cysts) and commonly result from sexual fusion. The resting cysts eventually sink to the sea floor and remain dormant. The duration of cyst dormancy varies among species and ranges from hours to days or months (Figueroa et al., 2006; Hallegraeff et al., 1998; Kremp and Anderson, 2000). Cyst dormancy is an important stage in the life cycle of dinoflagellates and determines bloom seasonality. However, the relationship between this endogenous (genetic) clock and seasonality is not well understood (Bravo and Figueroa, 2014). Motile dinoflagellates and even temporary cysts are often quickly degraded by bacterial action after death. In contrast, the resting cyst wall is composed of a highly resistant organic material (dinosporin), or of calcium carbonate or rarely, silica. Cysts are protective and usually permit the protoplasm to survive adverse conditions and can remain viable many years after burial and sampling of the sediments (McQuoid et al., 2002; Ribeiro et al., 2011). Excystment, i.e. the release of motile cells can occur when the dormancy is over. Cysts enter in quiescence and when environmental conditions are appropriate, excystment is triggered (Warns et al., 2012). In laboratory experiments germination can be triggered by temperature, light availability, oxygen concentration and endogenous rhythms (Matthiessen et al., 2005 and literature therein). About 200 living marine dinoflagellates, including a number of harmful species, are known to produce organic-walled dinoflagellate cysts (“dinocysts”) which are even preserved in conditions in which calcareous (coccoliths, foraminiferal tests) and siliceous (diatom frustules, radiolarian skeletons) microplankton remains are prone to dissolution (De Vernal and Marret, 2007). Thus, the study of the dinocysts recovered from modern sediments allows to record the dinoflagellate species even in times when the motile vegetative stage of the species is absent in the water column. Furthermore, the information on the distribution and abundance of dinocysts in sediments is essential to understand the ecology and dynamics of dinoflagellate blooms (Blackburn et al., 1989; Dale, 2001; Genovesi-Giunti et al., 2006; Matsuoka and Fukuyo, 2000). Areas with the highest concentrations of dinocysts of potentially toxic species are considered zones of risk for the initiation of new planktonic populations and future blooms of these microalgae. Studies of dinocysts recovered from surface sediments of the Southwest Atlantic are scarce. Offshore the Rio de la Plata a pioneer dinocyst record was obtained in a NW–SE transect (36–39°S) between 500 and 5000 m depth (Wall et al., 1977). Other records are available from the northern Argentine Shelf (Akselman, 1999; Borel et al., 2012) and only two studies exist from the Patagonian Shelf (Akselman, 1987; Orozco and Carreto, 1989).

In order to study the occurrence of harmful microalgae and their toxins in the water column and dinocysts in modern surface sediments along the Argentine Continental Shelf, we conducted an oceanographic expedition with the R/V Puerto Deseado in March/April 2012. In the course of this expedition, we detected high concentrations of dinoflagellates and the occurrence of toxicogenic species in the SGJ area (45–47°S) and related it to the physicochemical characteristics of the surrounding environment.

2. Methods

2.1. Study sites and sampling

During the austral autumn, from March 27th to April 14th 2012, plankton, surface water and sediment samples were collected on board R/V Puerto Deseado during the expedition “Patagonia Austral 2012”. The sampling stations covered a variety of geographical areas at 48 stations along the Argentinean coast. The transect from ~56°S (station 11) to 38°S (station 152) embraced 18° of latitude with an average sampling resolution of about 50′ (Fig. 1). The San Jorge Gulf transect (April 9th to 11th) consisted of six stations (S40, C43, C45, I44, C43N, P45B, Figs. 1A and 2A).

At all stations continuous profiles of temperature and salinity (CTD) were recorded. A vertical net tow was conducted at each station through the upper 20 m of the water column using a 20 μm-mesh Nitex plankton net. In addition, samples were collected using a Niskin bottle for nutrients analysis, pigments and quantitative plankton determination. The samples were grouped according to their oceanographic characterization (see below & Fig. 4). In the SGJ samples were grouped into surface (10 m) and deep water samples (>50 m; Table 1). Surface sediment samples were obtained from the top centimeter of box coring. All samples were kept in dark and cool (4 °C, –20 °C or –80 °C, as appropriate) conditions to preserve the cell integrity, stability of chlorophyll and nutrients, and prevent cyst germination (Matsuoka and Fukuyo, 2000).

2.2. Oceanographic variables and data processing

Data on temperature, salinity, and water depth were obtained using a CTD (Sea Bird model 911 plus with General Oceanic rosette) and calibrated with a final precision of 0.05 in salinity and 0.02 °C in temperature. All parameters were processed and displayed using Ocean Data View Software (ODV-version 4.6, Schlitzer, 2010). To determine the location of ocean fronts, the Simpson’s stability index was estimated at each station (Simpson, 1981). The Simpson parameter, a measure of the mechanical work required for vertical mixing of the water column, is defined as:

\[ \Phi = \frac{g}{\rho} \int_{-h}^{0} (\rho - \rho_0) \cdot z \cdot dz \]

where \( g \) is the gravity, \( h \) is the depth, and \( \rho_0 \) is the mean density of the water column. Small values of \( \Phi \) indicate poorly stratified waters while high values are associated with stratified waters (Simpson, 1981). A value of ~40 J m⁻³ was selected as the limit between homogeneous and stratified waters (Bianchi et al., 2005; Martos and Sánchez, 1997; Sabatini et al., 2000, 2004). The Brun–Väisälä (N²) was calculated as a measure of stratification variability of the water column, according to the equation:

\[ N^2 = -\frac{g}{\rho_0} \frac{\partial \rho}{\partial z} \]

where \( g \) = gravity of the Earth (9.8 m² s⁻¹), \( \rho \) = density of water, \( z = depth \).

2.3. Nutrient and pigment analyses

One liter of surface seawater was filtered (max. 300 mbar) using two precombusted (4 h, 450 °C) glass fiber filters (-0.7 μm, GF/F, Whatman). From the first 0.5 L, 40 mL aliquots were transferred into acid-precleaned HDPE bottles, poisoned with 120 μL of HgCl₂ (35 g L⁻¹) and stored at 4 °C for later nutrient analyses (Kattner, 1999). Duplicate filtrates for DOC and TDN determination were adjusted after filtration.
to \( \text{pH} = 2 \) (with \( \text{H}_2\text{PO}_4 \)) in 10 mL precombusted glass ampoules. These and the two filters with seston from 0.5 L each for pigment determination, were kept frozen \((-20 \degree \text{C})\) in the dark until further processing on land. Nutrients were measured using an autoanalyzer (Evolution III, Alliance Instruments) and standard seawater methods (Kattner and Becker, 1991).

Dissolved organic carbon (DOC) and total dissolved nitrogen (TDN) were determined by high temperature catalytic oxidation and subsequent non-dispersive infrared spectroscopy and chemiluminescence detection (TOC-VPN, Shimadzu). Final DOC and TDN concentrations are average values of triplicate measurements on replicate samples. If the standard variation or the coefficient of variation exceeded 0.1 \( \mu \text{M} \) or 1\%, respectively, up to 2 additional analyses were performed and outliers were eliminated. After each batch of five samples one reference standard (DOC-DSR, Hansell Research Lab, University of Miami, US), one ultrapure water blank and one potassium hydrogen phthalate standard were measured. The limit of detection (3 \( \times \) standard deviation \( \sigma \) of the blank) was 7 and 21 \( \mu \text{M} \) C, respectively. The accuracy was \( \pm 5\% \).

Pigment extraction on duplicate filter samples was performed in 10 mL 90\% acetone during 24 h at 4 \degree \text{C} in darkness. Chlorophyll-\( a \) and phaeopigments were quantified by fluorometry, and concentrations were estimated according to Holm-Hansen et al. (1965) equations. We used a standard of pure chlorophyll-\( a \) (Anacystis nidulans) and a Shimadzu RF-5301 fluorometer at \( A_{\lambda=436}/A_{\lambda=667} \) nm. The correction was made for phaeopigments by acidification with \( \text{HCl} \) (0.1 \text{N}).

Satellite data were retrieved from the Giovanni Ocean Color Radiometry 8-day (6–13 April, 2012) temporal resolution products from MODIS-Aqua 4-km database (http://disc.sci.gswf.nasa.gov/giovanni) on a 0.1° resolution. Chlorophyll-\( a \) (chl-\( a \)) concentrations derived from satellite images were retrieved from the SDN spreadsheet and plotted with the DIVA gridding option (Ocean Data View, version 4.6).

2.4. Size fractionation, cell counts and species identification

The sample volume of each net haul was topped up to 1 L with 0.2 \( \mu \text{m} \) filtered sea water. An aliquot of 20 mL was fixed with 400 \( \mu \text{L} \) Lugol’s solution for species determination and enumeration. The rest was sequentially filtered through Nitex meshes of 200, 50 and 20 \( \mu \text{m} \) by gravity filtration and split into aliquots for extraction of lipophilic and hydrophilic toxins. The aliquots of suspended plankton size fractions were pelleted by centrifugation for subsequent toxin analysis. In addition, plankton samples were collected at 3 and 10 m depth from Niskin bottles. 3.5 L of each depth were mixed and 100 mL of the mixed water was treated with 2 mL of Lugol’s solution for species identification and enumeration. The rest of the samples were pre-screened through a 20 \( \mu \text{m} \) mesh-size Nitex sieve and the filtrate was split in two 3-L aliquots. Each aliquot was then filtered under gentle vacuum through 3 \( \mu \text{m} \) pore-size polycarbonate filters (Millipore, Eschborn, Germany). Dry filters were placed individually in 50 mL centrifuge tubes (Sarstedt, Nümbrecht, Germany) and stored at \(-20 \degree \text{C}\) until analysis.

For phytoplankton quantitative estimations of Niskin bottle samples, cells were enumerated with a phase contrast Leica DMILL LED inverted microscope according to the procedures described by Utermöhl (1958). Subsamples of 50 mL were settled for 24 h in a composite sedimentation chamber. For the analysis of net samples, subsamples were counted in Sedgewick–Rafter chambers under the same microscope according to LeGresley and McDermott (2010). The whole chamber bottom was scanned at 200 \( \times \) magnification. Densities are expressed as cells in net tow (cells \( NT^{-1} \)). Additional qualitative observations of dinoflagellates were done by UV epifluorescence microscopy (Leica DM 2500) of thecal plates stained with Calcofluor according to Fritz and Triemer (1985). For diatom frustules observation, organic material was removed from net subsamples using sodium hypochlorite and mounted on permanent slides as described in Almandoz et al. (2011). Additionally, clean material was filtered onto 0.2 \( \mu \text{m} \) polyamide filters, sputter coated with Au–Pd and examined by scanning electron microscopy (Jeol JSM-6360 LV) to confirm the identity of Pseudo-nitzschia species.

For picoplankton abundance determination, 3 mL of surface seawater sample were preserved with 1% paraformaldehyde + 0.05% glutaraldehyde (final concentration), frozen and stored at \(-20 \degree \text{C}\). The samples were later thawed, stained with SYBR-Green I (SYBR-I, 1:30 dilution of commercial stock; Invitrogen, USA; \( l_{\text{ex}} \times 495 \text{ nm}, l_{\text{em}} \times 525 \text{ nm} \) diluted in dimethyl sulfoxide (DMSO, Merck, Germany) (Marie et al., 2005). Ten
microliters of fluorochrome was added to 1 mL of bacterial sample. Samples were incubated in the dark for 15 min at room temperature. Flow Cytometry analyses were performed with an Apogee-A40 flow cytometer (Apogee Flow Systems, UK) equipped with an argon laser (488 nm). Cells were counted following Marie et al. (2005). Briefly, heterotrophic bacteria (HB) were detected by their signature in a plot of side scatter (SSC) versus green fluorescence (FL1). Autotrophic picoplankton (AP) were detected by direct fluorescence in a plot of SSC versus red fluorescence (FL3).

2.5. Isolation of A. tamarense

At station C43, 100 mL of seawater from 3 m depth was filled into a polyethylene flask and kept at 4 °C in the dark until inspection in the laboratory eleven days after sampling. Single A. tamarense cells were isolated from this sample under a stereomicroscope (Olympus SZH-ILLD) by micropipette. The cells were transferred into individual wells of 96-well tissue culture plates (TPP, Trasadingen, Switzerland) containing 300 μL of medium, prepared from 0.2 μm sterile-filtered natural Antarctic seawater with 1/10 of the original concentration of K medium (Keller et al., 1987). Isolated cells were then incubated at 15 °C under artificial light at a photon flux density of 50 μmol photons m⁻² s⁻¹ on a 16:8 light:dark (L:D) photocycle. After 3 to 4 weeks, two unialgal and clonal isolates (provisionally named H5 and H7) were transferred to polystyrene cell culture flasks each containing 50 mL of ½ strength.
K medium and were maintained thereafter under the same conditions as described.

For toxin analysis, one culture of each isolate was grown in 250 mL Erlenmeyer glass flasks at 15 °C under a photon flux density of 50 μmol m$^{-2}$ s$^{-1}$ on a 16:8 h light:dark photocycle. Cell density was determined by settling lugol-fixed samples and counting $N > 800$ cells under an inverted microscope. Dense cultures (densities ranging from ca. 5000 (H5) to 10,000 (H7) cells per mL, and thus most probably representing late exponential phase) were harvested in 4 × 50 mL in 50 mL centrifugation tubes by centrifugation (Eppendorf 5810R, Hamburg, Germany) at 3220 × g for 10 min. All four pellets from one isolate were combined in a microtube and again centrifuged (Eppendorf 5415, 16,000 × g, 5 min) and stored frozen (−20 °C) until use.

Observation of live or fixed cells was carried out with a stereomicroscope (Olympus SZH-ILLD) and with an inverted microscope (Axiovert 200 M, Zeiss, Germany) equipped with epi-fluorescence and differential interference contrast optics. Light microscopic examination of the thecal plate tabulation was performed on formalin-fixed cells (1% final concentration) stained with calcofluor white (Fritz and Triemer, 1985).

Photographs were taken with a digital camera (Axiocam MRc5, Zeiss, Germany).

2.6. Toxin extraction and determination

Cell pellets from the plankton net tows and A. tamarense cultures were harvested by centrifugation (3220 × g, 15 min at 4 °C), suspended in 500 μL methanol for lipophilic toxins and 0.03 M acetic acid for hydrophilic toxins, and subsequently transferred into a FastPrep tube containing 0.9 g of lysing matrix D (Thermo Savant, Illkirch, France). The samples were homogenized by reciprocal shaking at maximum speed (6.5 m s$^{-1}$) for 45 s in a Bio101 FastPrep instrument (Thermo Savant, Illkirch, France). After homogenization, samples were centrifuged at 16,100 × g for 15 min. The supernatant was transferred to a spin-filter (0.45 μm pore-size, Millipore Ultrafree, Eschborn, Germany) and centrifuged for 30 s at 800 × g, followed by transfer to autosampler vials. The filtrates of the acetic acid extraction were analyzed for PSP toxins after separation of target analytes in reversed-phase mode by high-performance liquid chromatography with post-column derivatization and fluorescence detection (LC-FD), according to the method given in Fig. 4.

Table 1

<table>
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<tr>
<td>SW</td>
<td>8</td>
<td>78 ± 7</td>
<td>7.0 ± 3.6</td>
<td>2.0 ± 4.2</td>
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<td>SJG 10 m</td>
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<td>1.8 ± 1.8</td>
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<td>1.3 ± 1.3</td>
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<td>SJG &gt; 50 m</td>
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<td>65 ± 10</td>
<td>22.7 ± 1.7</td>
<td>17.8 ± 3.8</td>
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<td>10.9 ± 5.5</td>
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<td>0.1 ± 0.0</td>
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<td>52 ± 6</td>
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<td>5.1 ± 4.6</td>
<td>1.3 ± 0.2</td>
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<td>BMW</td>
<td>27</td>
<td>63 ± 6</td>
<td>17.1 ± 3.5</td>
<td>1.6 ± 2.5</td>
<td>0.3 ± 0.2</td>
<td>2.4 ± 0.6</td>
<td>1.0 ± 0.1</td>
<td>1.1 ± 0.9</td>
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Fig. 4. T–S diagram of CTD casts of the entire expedition showing the different water masses. Encircled are the SJG stations, which consist of both water types: SW (top) and BMW (bottom), Low-salinity coastal waters (LSCW) and high-salinity coastal waters (HSCW). Box, Temperature–salinity diagrams of selected stations occupied in the transect I (stations I44, C45, C43N and C 43) and transect II (stations S40, C43N and P45B) (see Fig. 2).
in Krock et al. (2007). Analysis of multiple lipophilic toxins in the methanolic extracts was performed by liquid chromatography coupled to tandem mass spectrometry (LC-MS/MS), as described in Krock et al. (2008).

For AZA analysis, polycarbonate filters with 20 μm pre-screened samples from the rosette bottles were repeatedly rinsed with 1 mL methanol until complete decolourisation of the filters. The methanolic extract was filtered through spin-filters as described above. AZA profiles were determined as previously described (Krock et al., 2009).

2.7. Sediment analysis for dinocysts

An aliquot (5–10 g wet weight) of the sediment sample from each station was sieved through 150 μm and 10 μm Nitek® screens to eliminate coarse sand, fine silt and clay. One calibrated tablet of Lycopodium clavatum (University of Lund, BATCH N° 483216) containing 18,583 spores was added to each sample as an internal standard (Stockman, 1971) to allow for the calculation of concentrations in cysts g⁻¹ sediment. The 10 to 150 μm grain size fraction was treated with cold 10% hydrochloric acid (HCl) to eliminate carbonates. A zinc chloride (ZnCl₂) solution adjusted to a density of ρ = 2.0 g cm⁻³ was used to separate the organic fraction based on a density gradient. Residues were sieved and collected on a 10-μm mesh and then mounted between slide and cover-slide in Kaiser gelatinized-glycerine (De Vernal et al., 1996).

Dinocysts were identified and counted using a transmitted-light microscope (Nikon Eclipse 600) at a magnification factor of 600× and 1000×. The nomenclature of the dinocysts follows Fensome et al. (1993) and the taxonomy for the motile dinoflagellates was used. A minimum of 200 dinocysts was counted per sample. Using these counts, the abundance of each taxon and the concentration of total dinocysts per mass unit of wet sediment [g] were determined.

Two cyst morphotypes of the genus Spiniferites are present (known as Spiniferites ramosus and Spiniferites mirabilis from the paleontological literature) and are grouped in this study as dinocysts of Gonyaulax aff. spinifera (following Head, 1996 and literature therein; Rochon et al., 2009).

The cysts referred in this study as A. tamarense type are morphologically identical to the cysts of A. tamarense but were not incubated to confirm their identity. Other species as A. catenella, Alexandrium acatenella and Alexandrium fundyense produce similar cysts hardly distinguishable from A. tamarense (Head et al., 2006) and literature therein) although A. catenella and A. fundyense are probably co-specific and molecular studies point in that direction that they form the same clade using theca-based phylogeny and the near-identical cyst morphologies reflect this close phylogenetic relationship (Matsuoka and Head, 2013).

3. Results

3.1. Hydrography of the San Jorge Gulf (SJG)

Further insight into the dynamical characteristics of the hydrographic fields can be obtained from the Simpson's stability parameter and Brunt–Väisälä frequency. The water column in the gulf region (Fig. 2) showed a strong stratification in the inner part of the gulf (values higher than 100 m⁻²) and outside of the gulf at the easternmost station (transect I, Fig. 2A, Fig. 2B/C) (values higher than 60 m⁻²).

The water column appeared highly stratified and stable in the first 50 m of stations I44, C45 and C43 with values of Brunt–Väisälä frequency greater than 155 cycles h⁻¹ (transect I, Fig. 2A, Fig. 2B/C). The stabilization (and density structure) of the water column was mainly determined by temperature and salinity, with deeper water showing temperatures as low as -8 °C and the upper water layer reaching almost 14 °C. In contrast, at stations S40, C43N and P45B (transect II, Fig. 2A, Fig. 2D/E) the Simpson parameter was only 0.4, 4.7 and 0.7 m⁻², respectively. The Brunt–Väisälä frequency was very low and the water column was mixed (transect II, Fig. 2A, Fig. 2D/E). The transition of water masses is further illustrated by a T–S diagram of selected stations (Fig. 4). On average, the Subantarctic Waters (SAW) were more saline, with salinities >33.50 and presented changes in the vertical temperature. The Beagle–Magellan Waters (BMW) were less dense than 25.50 kg m⁻³, with constant temperature (~8 °C), and influenced by low-salinity coastal water (LSCW). The Argentinean Shelf Waters (SW) showed the highest temperatures and, at the northern stations were influenced by high-salinity coastal water (HSCW). The stations of the study area (Fig. 4, box) are illustrated in more detail. The stations of transect I showed variations in salinity and temperature in contrast to the stations of transect II.

3.2. Nutrients

Inorganic nutrients were largely depleted and DOC was slightly elevated at the surface (10 m water depth) of all stations in the SJG (Table 1), similar to all other stations north of the SJG at which SW was sampled. Only in the surface water of station P45B nitrate was elevated (11.4 μM nitrate). At water depths below 50 m inorganic nutrients were high (>14 μM nitrate >1.4 μM phosphate). Station C43, east of the SJG, also showed higher nutrients in the deeper water (88 m: 17.7 μM nitrate and 1.5 μM phosphate). Nitrate, phosphate and DOC concentrations in the

### Table 2

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<th>Water mass</th>
<th>n</th>
<th>Chl-a [μg L⁻¹]</th>
<th>Phaeo [μg L⁻¹]</th>
<th>Ratio Chl-a/Phaeo</th>
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<td>SW</td>
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<td>1.74 ± 0.6</td>
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<td>BMW–SAW</td>
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<td>0.45 ± 0.2</td>
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### Table 3

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<tr>
<th>Station</th>
<th>Total dinoflagellates</th>
<th>Ceratium tripus</th>
<th>Ceratium lineatum</th>
<th>Ceratium fusus</th>
<th>Total toxic dinoflagellates</th>
<th>Proteroceratium reticulatum</th>
<th>Gonyaulax aff. spinifera</th>
<th>Dinophysis acuminata</th>
<th>Alexandrium tamarense</th>
<th>Total diatoms</th>
<th>P. pungens</th>
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<td>S40</td>
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<td>33.85</td>
<td>25.56</td>
<td>5.70</td>
<td>9.06</td>
<td>3.56</td>
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<td>0.22</td>
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<td>16.01</td>
<td>0.72</td>
<td>1.58</td>
<td>0.06</td>
<td>0.04</td>
<td>0.06</td>
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<tr>
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<td>–</td>
<td>–</td>
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### Table 4

<table>
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<th>Water mass</th>
<th>n</th>
<th>Chl-a [μg L⁻¹]</th>
<th>Phaeo [μg L⁻¹]</th>
<th>Ratio Chl-a/Phaeo</th>
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<td>SJG</td>
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<tr>
<td>BMW–SAW</td>
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<table>
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<tr>
<th>Station</th>
<th>Total dinoflagellates</th>
<th>Ceratium tripus</th>
<th>Ceratium lineatum</th>
<th>Ceratium fusus</th>
<th>Total toxic dinoflagellates</th>
<th>Proteroceratium reticulatum</th>
<th>Gonyaulax aff. spinifera</th>
<th>Dinophysis acuminata</th>
<th>Alexandrium tamarense</th>
<th>Total diatoms</th>
<th>P. pungens</th>
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<td>S40</td>
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<tr>
<td>I44</td>
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<tr>
<td>C45</td>
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<td>0.40</td>
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BMW were moderately high. Compared to the SJG and SW, inorganic nutrients in the SAW samples south of the SJG were higher and DOC was lower.

3.3. Phycopigments and satellite images

Mean chl-a values derived from satellite images of the SJG for the week between 6th and 13th April 2012 were highest (up to 5 mg m⁻³ chl-a) to the east of SJG mouth (Fig. 3A), where the water column was stratified (Fig. 2B/C). In general chl-a values were lower in the SJG area during the two weeks before and after 6th to 13th April 2012 (data not shown).

Relative low chl-a concentrations were detected in the inner SJG (Fig. 3A). Chl-a concentrations measured in situ (Table 2) were generally in the same range with the data derived from satellite images. Ratios of chl-a to phaeopigments throughout the entire expedition transect ranged between 0.08 and 4.92 and were highest in the SJG (Table 2).

3.4. Phytoplankton composition in the SJG

3.4.1. Nano- and microplankton

Abundance of phytoplankton (>2 μm), estimated by microscopic analysis following the Utermöhl method, showed similar values in the SJG and adjacent waters, ranging between 9.0 × 10⁴ and 36.4 × 10⁵ cells L⁻¹ (Fig. 3B). However, assemblages dominated by dinoflagellates and diatoms were respectively observed inside and outside the SJG (Fig. 3B). A bloom of Ceratium spp. (mainly Ceratium tripos but also Ceratium lineatum and Ceratium fusus), representing up to 85% of total phytoplankton abundance, was found towards offshore stations of the SJG (C43 and C43N). Small gymnodinioid dinoflagellates (<20 μm) exceeded Ceratium abundances at the inner stations of the SJG. In contrast, samples collected both north and south of the SJG were dominated by diatoms, which represented 58% (station P45B) and 80% (station S40) of total abundance. Station P45B, with a total diatom abundance of 6.1 × 10⁵ cells L⁻¹, was dominated by centric diatoms including a small Thalassiosira (10–30 μm) and Paralia sulcata.

On the other hand, station S40 showed the highest diatom abundance (27 × 10³ cells L⁻¹). It was dominated by centric taxa, mainly Rhizosolenia pungens, Paralia sulcata and Thalassiosira sp. (40–50 μm), but pennate diatoms of the genus Pseudo-nitzschia were also abundant (Tables 3 and 4).

3.4.2. Heterotrophic bacteria and autotrophic picoplankton

Highest numbers of heterotrophic (HB) (average of 1.18 × 10⁶ cells mL⁻¹) and autotrophic picoplankton (AP) (average of 2.64 × 10⁴ cells mL⁻¹) were found at stations to the north of SJG. Lower levels of bacteria were detected in the SJG (Table 5). Our data (Fig. 5) suggest a latitudinal trend with decreasing HB abundances from the north to the northern edge of the SJG (42°S to 45°S) and increasing abundances from the southern edge of the SJG to the south and a subsequent decrease (47.5°S to 55.5°S). The AP abundances seem to follow the same trend, and this seems to depend on a physical variable such as temperature (see data subgroup in Fig. 5 and below). However, the ratios between AP and HB are lower in SW in the north than in BMW in the south.

In the SJG and SW regions HB and AP showed a positive correlation (Table 6) with surface seawater temperature (SST), and HB also with surface seawater salinity (SSS). No significant correlations were found between HB and chl-a, phaeo. However, AP showed a positive correlation only with phaeo (Table 6). An exception was the station I47, (Valdés Front), this area is characterized by highest productivity (Acha et al., 2004) and densities of bacterioplankton on the Argentine shelf (Alder and Franzosi, 2004) (Fig. 5, dashed circle). South of the SJG, in the Beagle–Magellanic and Subantarctic waters (BMW–SAW) HB and AP showed a positive correlation with SST. They also showed a strong negative correlation with SSS. Significant correlations were found for picoplanktonic abundance and chl-a and phaeo (Table 6).

3.4.3. Description of the isolate of A. tamarense

Based on the thecal plate pattern determined under the fluorescence microscope (exemplarily shown for isolate H7 in Fig. 6), both isolates were identified as A. tamarense. During exponential growth in culture, cells mainly occurred as single units or in short two-cell chains. Cells with the typical horseshoe-shaped central nucleus were nearly spherical (Fig. 6A), with length and width of about 35 μm. The first apical plate consistently had a small ventral pore (Fig. 6D).

The PSP toxin profiles of strains H5 and H7 are almost identical consisting of 57% GTX1/4, 40% C1/2 and minor proportions of NEO, GTX2/3, B1 and STX (Fig. 7). Total PSP toxin content of strains H5 and H7 was 74 pg cell⁻¹ and 60 pg cell⁻¹, respectively.
3.4.4. Dinocysts in surface sediments

Photosynthetic taxa in the surface sediments were represented by cysts of *Gonyaulax aff. spinifera*, *Protoceratium reticulatum* and *A. tamarense* type (Fig. 8A). Heterotrophic assemblages included cysts of diverse species of *Protoperidinium*, *Archaeperidinium minutum*, *Diplopsalidaceae*, and the *Polykrikos kofoidii/Polykrikos schwartzii* complex. These cysts were grouped as heterotrophic cysts for the purpose of this work (Table 7). The total dinocyst abundance ranged from $0.64 \times 10^3$ to $2.1 \times 10^3$ cysts g$^{-1}$ sediment wet weight (with a mean of $1.4 \times 10^3$ cysts g$^{-1}$), except for station C43N where a much higher value ($17.1 \times 10^3$ cysts g$^{-1}$) was reached (Fig. 8B). The bottom sediment texture is heterogeneous (Fig. 8B) consisting of sabulitic (coarse) sandy silt, sandy silt and clayey silt, suggesting different hydrodynamic conditions with a probable control over the distribution of dinocysts, similar in size to fine silt particles of the sediment.

3.4.5. Occurrence of toxic phytoplankton species

Six potentially toxic phytoplankton species were found in the SJG, four of them dinoflagellates (*A. tamarense*, *Dinophysis acuminata*, *Gonyaulax aff. spinifera* and *P. reticulatum*) and the other two diatoms (*Pseudo-nitzschia australis* and *Pseudo-nitzschia pungens*). *A. tamarense*, the putative source of PSP, was present in all net samples examined except in P45B located north of the SJG and C43N, where the lowest concentration of PSP toxins was detected. Moreover, the highest abundance of $148 \times 10^3$ cells NT$^{-1}$ coincided with the maximum PSP concentration (Table 4, Fig. 9A). *Gonyaulax aff. spinifera* and *P. reticulatum*, two potential producers of YTX, were present in the SJG stations and also northward, and coincided with the occurrence of yessotoxins (Fig. 9B/C). However, no clear correspondence was found between cell abundance and toxin cell quota (Fig. 7).

![Fig. 6. Micrographs of *A. tamarense* isolate H7. (A) Live cell. (B) Lugol-fixed cell. (C-F): Calcofluor-stained cells with UV excitation. (C) Whole cell overview. (D) Epitheca in ventral view. Note the small ventral pore (vp, arrow). (E) Epitheca in dorsal view. (F) Hypothecal plates. Note that the dorsal postcingular plate 3‴ is not visible. Scale bar = 10 μm.](image)

![Fig. 7. PSP toxin profiles and toxin cell quotas of *A. tamarense* strains H5 and H7 isolated from station C43.](image)

<table>
<thead>
<tr>
<th>Variable</th>
<th>HB$_{SGJ-SW}$</th>
<th>AP$_{SGJ-SW}$</th>
<th>HB$_{BMW-SAW}$</th>
<th>AP$_{BMW-SAW}$</th>
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<td>0.67*</td>
<td>0.43*</td>
<td>0.78**</td>
</tr>
<tr>
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<td>ns</td>
<td>-0.55**</td>
<td>-0.46*</td>
</tr>
<tr>
<td>Chl-a</td>
<td>ns</td>
<td>ns</td>
<td>0.49*</td>
<td>0.56**</td>
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</tbody>
</table>

*p < 0.05; **p < 0.01; ns, not significant.

Table 6

The Pearson’s product–moment correlation coefficient. The p value is based on significance test for the correlation coefficient. SST, surface seawater temperature; SSS, surface seawater salinity; Chl-a, chlorophyll a; HB, heterotrophic bacteria; AP, autotrophic picoplankton; SJG, San Jorge Gulf; SW, shelf water; BMW, Beagle-Magellan water; SAW, Subantarctic water.
...density and toxin concentration for any of the two species. A similar case was found for *D. acuminata*, the putative source of PTX-2. While a general coincidence in the presence of *D. acuminata* and PTX-2 was observed, no clear relation was found between cell density and toxin concentration. The diatom genus *Pseudo-nitzschia* was found in low densities in the two southernmost stations of the SJG. In contrast, a maximum was represented by *Domoic acid* and dinophysistoxins were not detected along the entire transect. Only in the SJG region the 50–200 μm fractions contained DA and PTX-2 up to approximately 500 ng NT⁻¹ and in lower concentration YTX. In addition, 2400 and 230 ng NT⁻¹ PTX-2sa were measured at stations I50 and I52, respectively, but no PSP toxins were found in the 50–200 μm size fractions. In the >200 μm fraction only sporadic and low (max. 100 ng NT⁻¹) levels of PTX-2sa were detected at few stations.

### 4. Discussion

#### 4.1. Water masses of the Argentinean shelf

The oceanographic data collected during the expedition (data not shown) were consistent with previous descriptions of Argentinean coastal water masses. The main source of water masses in the Argentinean Shelf is the Malvinas/Falkland Current (MC) that originates from Subantarctic Waters (SAW). SAW is derived from the north of the Drake Passage, flows as Cape Horn Current (Guerrero and Piola, 1997; Piola et al., 2010) and turns northward as the MC between the eastern edge of the Argentinean shelf with the Islas Malvinas and the Argentinean coast (Bianchi et al., 1982). Significant contributions of low-salinity waters enter the MC from continental run off and melt water through the Strait of Magellan and the Beagle Channel (BMW), which is indicated by lowest salinities in the region at the eastern mouths of the Beagle Channel and the Strait of Magellan (Fig. 1B). The BMW causes the formation of a coastal low-salinity tongue (LSCW) (~33.4, see Fig. 1B), which can be detected up to about 45°S, and then continues with shelf water or midshelf water (SW), ranging from 33.4 to 33.8, along the central shelf up to 40°S (Bianchi et al., 2005). High-salinity coastal waters (HSCW) close to north of 42°S and continue to the north to cold Poles with Rio de la Plata estuary.

#### 4.1.1. Oceanographic parameters and biogeographical association of planktonic species with particular water masses

The general water mass characteristics in the San Jorge Gulf were particularly driven by four factors: (i) the Argentinean Shelf Water (SW), (ii) the Beagle–Magellan Water (BMW), (iii) the Malvinas Current (MC) and (iv) the topographical depressions in the inner part of the gulf (Matano et al., 2010).

#### Table 7

<table>
<thead>
<tr>
<th>Sta.</th>
<th>Total autotrophic cysts</th>
<th>Total heterotrophic cysts</th>
<th>Protoceratium reticulatum</th>
<th><em>Gonyaulax aff. spinifera</em></th>
<th><em>Alexandrium tamarense</em></th>
<th>Total cyst</th>
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<tr>
<td>S40</td>
<td>0.04</td>
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<td>0.008</td>
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<td>0.33</td>
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<td>0.006</td>
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<td>11.75</td>
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<td>11.35</td>
<td>0</td>
<td>17.09</td>
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<td>nd</td>
<td>nd</td>
<td>nd</td>
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<td>nd</td>
</tr>
<tr>
<td>C45</td>
<td>1.52</td>
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<td>0.90</td>
<td>0.37</td>
<td>1.85</td>
</tr>
<tr>
<td>P45B</td>
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<td>0.81</td>
<td>0.22</td>
<td>1.00</td>
<td>0.06</td>
<td>2.09</td>
</tr>
</tbody>
</table>

...only detected in the Eastern Beagle Channel and in the mouth of the Beagle Channel (stations I2 and I3) and in the SJG region with a maximum value of 295 ng NT⁻¹ at station C45. Yessotoxin (YTX), a lipophilic algal toxin usually co-extracted with diarrhetic shellfish poisoning (DSP) toxins, was exclusively found in the SJG region (24–343 ng NT⁻¹), whereas pectenotoxins (PTXs), also regarded as a toxin group of the DSP complex, were occasionally found in varying abundances (4–680 ng NT⁻¹) along the entire expedition transect from latitude 56°S (south of Tierra del Fuego) up to latitude 38°S (Mar del Plata). Paralytic shellfish poisoning (PSP) toxins represented the most abundant toxin group with up to 5.3 μg NT⁻¹ in the SJG region (Fig. 9A, Table 8). Yet, lower levels were also found south of Tierra del Fuego (56°S) and in the Bahía Grande area (51–52°S). PSP toxin profiles consisted only of the equilibrium ratio of GTX2/3 and traces of STX. Azaspiracids, cyclic imine toxins like spirilides and gymniodimines, pinnatoxins, okadaic acid and dinophysistoxins were not detected along the entire transect. Most toxins were found in the 20–50 μm size fractions, where most toxigenic plankton species are found. Only in the SJG region the 50–200 μm fractions contained DA and PTX-2 up to approximately 500 ng NT⁻¹ and in lower concentration YTX. In addition, 2400 and 230 ng NT⁻¹ PTX-2sa were measured at stations I50 and I52, respectively, but no PSP toxins were found in the 50–200 μm size fractions. In the >200 μm fraction only sporadic and low (max. 100 ng NT⁻¹) levels of PTX-2sa were detected at few stations.

#### 3.5. Phycotoxins

Four toxin groups were detected in phytoplankton samples along the entire cruise transect from Ushuaia to Mar del Plata: domoic acid (DA) causing the amnesic shellfish poisoning (ASP) syndrome was...
A

B

C

Station
The SJG water mainly originates from SW modified by the contributions of coastal waters, whose salinity spatially and temporally varies between 33.0 and 33.6 (Fernández et al., 2007). Coastal waters flow northward from the Strait of Magellan along the coast (Fig. 1B) and consist of a mixture of SAW and low-salinity BMW waters derived from continental input (Brandhorst and Castello, 1971) (Fig. 4). In the SJG region, it splits into two branches: one branch enters to the gulf in the southeast and influences the entire region throughout the year. The other branch flows away from the coast towards N–NE, thus separating the water masses inside the SJG from those, east of the gulf. Between these two stratified water bodies lies a zone of well mixed waters above a ridge in the mouth of the SJG (Fig. 2) driven by winds. Our CTD profile data agree well with previous findings on tides (Piola and Rivas, 1997) and the general oceanography of the SJG (Baldoni, 1990). Tidal fronts generate seasonal development and retention during the spring and late autumn (Carreto et al., 1985, 1986; Glorioso and Flather, 1997; Guerrero and Piola, 1997).

At the surface of all stations inside and near the gulf, the water was derived from the warm and saline Brazil Current referred as SW (Brandini et al., 2000). In the deeper layers of the stratified stations the water was derived from cold and less saline BMW (Piola and Rivas, 1997). The influence of the Patagonian Current was verifiable at intermediate depths (45 m) and characterized by a particularly low salinity. The occurrence of this water mass in the SJG was previously described in Gómez et al. (2011) and our results confirm this. Even though the water in the deep basin of the SJG (station I44) was derived from nutrient enriched BMW, total inorganic nitrogen (TIN) and phosphate in the basin was higher compared to BMW. As the nutrient input from rivers in the arid climate around the gulf is low, higher nutrients in the bottom water can only be explained by remineralization.
of organic matter. These additional nutrients may also be explained by sedimentation of decaying plankton blooms, which frequently occur in the SJG. In addition, high nutrient conditions can also only build up if water residence times in the deep basin are relatively long compared to the overlying shelf water. This interpretation is supported by the formation of quasi stationary (counterclockwise) gyre and boundary currents in SJG generated by the wind-driven circulation as is predicted by the numerical model of circulation of Palma et al. (2004).

The BMW entering the SJG generates an intense thermohaline front from its southern end, which would stay at variable positions in the SJG during southern winter, spring, and summer, but moves to the north in autumn as has been proposed by several authors (Akselman, 1996; Guerrero and Piola, 1997).

These conditions of the water column define environmental conditions that favor dinoflagellate blooms (Gargett, 1997; Smayda, 1997). At the stations within the SJG, where BMW was superimposed by SW, the dinoflagellate proportion was always >80% and dinoflagellates dominated diatoms. At the adjacent stations in the north and the south of the SJG, which were not stratified and influenced only by one water mass, the dinoflagellate:diatom ratio was almost reversed with diatoms contributing up to 60 and 80% of the total plankton community at station P45B and S40, respectively (Fig. 3B).

It has to be taken into account that the determination of plankton communities is highly dependent on the applied methods. Plankton cell counts of a defined volume of sea water are regarded to be quantitative values, even though they represent the sampled depth and are not necessarily representative for the whole water column. In contrast, vertical net hauls can be regarded as more representative of the water column at a certain sampling point. However, plankton smaller than the mesh size of the net (here: 20 μm) are not retained and plankton net hauls are not quantitative in a strict sense, because it is difficult to precisely estimate the volume sampled by the net. For these reasons, ratios of net haul and sea water plankton counts can be quite different (Tables 3 and 4).

It has been reported before that stratified water masses favor photosynthetic dinoflagellates, which, in contrast to diatoms, due to their flagella show active swimming behavior and perform diurnal migrations in order to take up nutrients in deeper parts of the water column during night and to photosynthesize close to surface during day-time (Cullen and Horrigan, 1981; Eppley et al., 1968; Ji and Franks, 2007; Smayda, 1997). Since a preferential abundance of dinoflagellates for stratified water was also reflected by the distribution of phycotoxins. The only phycotoxin sporadically detected throughout the entire expedition transect (Fig. 9A) was pectenotoxin-2 (PTX-2), produced by several species of the genus Dinophysis. However, Dinophysis, even though they contain chloroplasts, are obligate predators and for sustained growth need the dinoflagellate:diatom ratio was almost reversed with diatoms contributing up to 60 and 80% of the total plankton community at station P45B and S40, respectively (Fig. 3B).

4.2. Abundances of heterotrophic bacteria and autotrophic picoplankton

A difference between the HB and AP abundance was observed that seemed to depend on the physicochemical conditions of the water masses in the south (BMW) and north (SW) of the SJG (Figs. 1B, 2B/C). The SJG proved to be a transition zone between these large water masses.

The positive covariation found to the south of 47°S for picopelagtonic abundance and chl-a (Table 2), showed the importance of picopelagton for the contribution to primary productivity in this zone. However, this scenario is not the same for the Argentine shelf north of 47°S, where no correlations between AP and chl-a were found.

Our results show the coincidence of decreasing bacterial and increasing dinoflagellate abundances, and this seems to be characteristic during blooms dominated by dinoflagellates in the SJG. It has been shown that bacteria generate favorable environments for the development of phytoplankton, especially dinoflagellates and the formation of ecdysis of their cysts (Mayali et al., 2007).

4.3. Dinocyst composition

The dinocyst composition of surface sediments showed a dominance of photosynthetic taxa in the SJG except for station S40, where a high number of cysts of heterotrophic species occurred. High abundance of heterotrophic species might indicate high abundance of prey organisms, and many of the Protoperidinium species have been shown to preferentially feed on large diatoms (e.g. Sherr and Sherr, 2007; Taylor and Pollinger, 1987). The analyses of the phytoplankton community in the water column are consistent with this observation showing an important increase in the number of diatom cells (26.8 × 10⁷ cells L⁻¹ at station S40, Table 3), sustained here by the nutrient-rich surface BMW. The abundance of Gonyaulax aff. spinifera cysts predominated followed by cysts of P. reticulatum. The only exception was observed at station C45 where cysts of the A. tamarense type reached the highest values exceeding those of P. reticulatum. Cysts of A. tamarense were also registered at C43 east of the SJG, at S40 south of the SJG and at P45B north of the SJG. This was in agreement with the detection of the related motile cells in the phytoplankton community.

The elevated cyst concentration of the photosynthetic group (11.75 × 10³ cysts g⁻¹) at station C43N, represented almost exclusively by a high value of G. aff. spinifera cysts (11.35 × 10³ cysts g⁻¹), correlated with high numbers of motile stages of this taxon (76.0 × 10⁵ cells NT⁻¹, see Table 4). However, this species was more abundant in the motile stage of station C45 (82.0 × 10⁵ cells NT⁻¹, see Table 4) and no clear relation was observed with the cyst abundance (0.9 × 10⁵ cysts g⁻¹). It is necessary to consider that the water-column analysis of planktonic communities are only temporal snapshots whereas sediment assemblages reflect an average of cysts produced by surface communities over time (Matthiessen et al., 2005). Other factors, such as granulometry and accumulation rate of bottom sediments, need to be taken into account. Dinocysts are comparable in size to medium-coarse silt particles and their hydrodynamic behavior is similar to that of other small sedimentary particles and thus become selectively concentrated in fine-grained sediments (Tyson, 1995). Previous work has demonstrated that the frontal regions encompass the location of peak cyst concentration and the examination of sediments underlying bloom areas indicated that the cysts seeding the bloom are not exactly within the bloom area (Tyler et al., 1982). At the Patagonian Shelf, Carreto et al. (1998b) described an important concentration of A. tamarense cysts in the tidal front localized at Valdés Peninsula.

Within the region where the motile phase forms blooms, hydrographic and sedimentary processes commonly deposit the cysts in areas of fine-grained sediment termed “seed beds”. The granulometry for the analyzed sites is consistent with the spatial distribution of the sediments described by Fernández (2006) with a sand fraction in the southern coast, predominance of gravels and sands in the extreme
south of the gulf and more content of silt and clay in the central and deeper zones. At station C45 the sediments had a coarse fraction (sand) besides a fine content (silt) whereas sediments of station C43N consisted only of a fine fraction (clay and silt). This difference could explain the order of magnitude increase of the cyst abundance at C43N. Moreover, the lowest total abundances in stations C43 and S40 (with 0.64 × 10^3 and 0.89 × 10^3 dinocysts g^-1, respectively) were correlated with coarser size-grains in the sediments (sabulitic). As we do not have data of sedimentation rates for the area, we cannot directly compare cyst concentrations among sampling sites. However, the studied cyst beds in the bottom sediments of the area probably reflected concentrations over many years indicating that recurrent blooms of Gonyaulax aff. spinifera are common in the area. Another possibility is that the high abundance shows a selective deposition of cysts in the thermohaline frontal system in the mouth of the SJG. The abundance of P. reticulatum cysts reach a value of 0.4 × 10^3 cysts g^-1 at station C43N that means twice as many than inside (station C45) or north (station P45B) of SJG. The benthiic cysts recovered in this study inside and at the edge of the SJG highlighted the latent possibility for the development of local blooms of potentially toxigenic dinoflagellates.

4.4. Phycotoxin composition and distribution in field plankton samples

Even though the dinoflagellate bloom was dominated by Ceratium spp., which formed up to 89% of the total dinoflagellate abundance in the net haul of station C43, many other dinoflagellate taxa including toxigenic species such as A. tamarense, P. reticulatum, and D. acuminata were present at lower abundances (Table 4). Generally, the toxigenic species co-occurred with their related toxins, yet with highly varying ratios between producing species and associated toxin groups (Fig. 5B/C). This variability may be due to the generally low abundances of toxigenic species. In addition, predation among planktonic species is a common feature and thus may jeopardize the ratio of toxin producer and their toxins. However, there were less toxins detected in the larger 50–200 μm and >200 μm size fractions (Table 8), which indicates that toxin transfer to higher trophic levels did not play an important role. However, heterotrophic and mixotrophic dinoflagellates are known for a low predator:prey size ratio (Hansen et al., 1994; Jeong et al., 2010) and thus may transfer toxins among species of the same size.

The PSP toxin profiles of all field samples consisted only of GTX2–3, which contradicts the toxin profiles of the two strains of A. tamarense isolated from station C43. Both strains displayed almost identical PSP profiles consisting of approximately 41% C1/2, 57% CXT1/4 and 2% NEO. This profile has been reported several times for Argentinean strains of A. tamarense (Carreto et al., 2001; Montoya et al., 2010). The fact that the field samples displayed a different toxin profile than the local A. tamarense isolates, may be due to a 3-months delay of the sample transport from Argentina to the laboratory in Germany. Apparently, during this time a conversion took place from the original toxins to GTX2/3. It is well known that the N-sulfocarbamoyl C-toxins easily cleave off the N-sulfonyl group to form GTX2/3. Interestingly, in addition to the desulfonilation of the C-toxins, N-dehydroxylation of GTX1/4 and NEO to GTX2/3 and STX, respectively, seems to occur under higher ambient temperatures during several months’ periods. However, this observation is only a first indication for such processes and has to be studied in more detail under controlled conditions.

5. Conclusions

Hydrographic conditions favor dinoflagellate blooms in the SJG region in general. Even though the plankton community was almost monospecific and dominated by Ceratium tripos, many other dinoflagellate taxa, including HAB species, were present at lower concentrations in the SJG, but were hardly detected outside the Gulf at that time. However, despite the fact that the superposition of stratified water masses favor dinoflagellate proliferation in general, the precise conditions that determine the proliferation of one or the other dinoflagellate species are not well understood. But the presence of toxigenic taxa in the water column and their cysts in the sediments of the SJG clearly indicates the potential for harmful algal blooms in this area. For this reason a better understanding of the processes leading to these dinoflagellate blooms in this highly productive area is necessary, due to the importance of the SJG region for fishery industries and marine exploitation. For a better understanding of the plankton dynamics of the SJG, more integrated studies are needed to increase the temporal data resolution for this region. More detailed sediment and cyst bed analyses, determination of sedimentation rates and drift studies on occurring blooms would complement the data presented in this study.

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