Seismic stratigraphy along the Amundsen Sea to Ross Sea continental rise: A cross-regional record of pre-glacial to glacial processes of the West Antarctic margin

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The seismic sediment record of the Amundsen Sea continental rise provides insight into the sedimentation processes from pre-glacial to glacial times, variations in ocean-bottom circulation, early ice sheet growth, and intensification towards the present icehouse regime. Seismic reflection data acquired during the 2010 RV Polarstern and the 2006 RV Tangaroa expeditions, created a ~2000 km long continuous Amundsen Sea to Ross Sea seismic transect. Pre-existing lines linked to this transect, connect key seismic stratigraphy horizons from the Ross Sea shelf to the rise and farther along the West Antarctic margin, up to the Amundsen Sea Embayment. Seismic units AS-1 to AS-5 constitute the Cretaceous to Eocene pre-glacial (PG) sequence (79–34 Ma), units AS-4 to AS-6 the Eocene to mid-Miocene transitional (T) sequence (34–15.5 Ma), and units AS-7 to AS-11 the mid-Miocene to Quaternary full glacial (FG) climate sequence (15.5–0 Ma). The top PG sequence boundary horizon AS-u3/uPG-T, links to unconformity RSU6 of the Ross Sea shelf and to the base of Unit II of the eastern Amundsen Sea, and is interpreted as the first arrival of grounded ice on the shelf. The top T sequence boundary AS-u6/uT-FG, links to RSU4 and the base of Unit III, and is interpreted as the onset of the FG regime with intensified ice sheet advances onto the outer shelves. The Amundsen Sea basin accumulated up to 3.9 km thick sediments in its centre. Seismic facies geometry analysis suggests Paleoocene–Eocene bottom-current activity, late Eocene shelf grounding of the West Antarctic Ice Sheet, and no apparent difference in the deep-sea sediment transport processes or temporal shift in deposition between the Amundsen Sea and Ross Sea. Implications for a Marie Byrd Land uplift starting at ~30 Ma are observed by a progressive change in horizon dip in the central Amundsen Sea seismic sequences.

1. Introduction

West Antarctica has played a vital role in palaeoenvironmental and palaeoclimatic processes due to its low-lying topography sculpted by West Antarctic Rift System Cenozoic tectonics and crustal thinning. It hosts an ice sheet of 3–5 m sea-level equivalent (Bamber et al., 2009; Fretwell et al., 2013) that is potentially vulnerable to global warming and likely to be affected by changes in global ocean circulation (e.g. Pritchard et al., 2012). The answers to an improved understanding of processes have been studied via marine seismic reflection images.

The sedimentary record off the Pacific margin of West Antarctica captured the last 90 million years of syn- to post-breakup pre-glacial evolution, the earliest Antarctic glaciation, ice sheet development and ocean circulation history. At the stages where continental-scale ice sheets began to expand and grounded on the continental shelf, erosion and rapid transport moved large amounts of sediments from inland to the coast, shelf, slope, and deep-sea. Such increased sediment supply and changes in transport/depositional processes combine to form sedimentary features that can be identified in seismic reflection images. The characteristics of these features, such as the geometry and reflectivity, can be related back to a general depositional environment or process (e.g. Shanmugam, 2006).

Whilst records for Antarctic sediment supply and deposition processes have been studied via marine seismic reflection images around Antarctic continental rises (e.g. Nitsche et al., 1997, 2000; De Santis et al., 2003; Rogenhagen et al., 2004; Escutia et al., 2005, 2011; Maldonado et al., 2006; Scheuer et al., 2006; Leitchenkov et al., 2007a, b; Uenzelmann-Neben and Gohl, 2012, 2014; Lindeque et al., 2013), the Marie Byrd Land sector of the southern Pacific between the Ross
Sea and the eastern Amundsen Sea basin remains largely unexplored. Thus its contribution to the West Antarctic sediment budget from pre-glacial to glacial times is unknown. Knowledge of such sediment budgets are important, for instance, to improve reconstructions of Antarctic palaeo-topographies at times of climatic relevance (Wilson et al., 2012) and to constrain boundary conditions for palaeo-ice sheet models (Wilson et al., 2013). Other important questions addressed by studying the sedimentary record of the Amundsen Sea and eastern Ross Sea continental rises include:

- Which part of the sediment record indicates the location and time of the onset of the glacially dominated environment of West Antarctica?
- Where are the major deposition centres for the pre-glacial to glacial periods, and what processes may have caused their formation?
- Is there a difference in the deep-sea sediment transport processes or a shift in deposition between the Amundsen Sea and Ross Sea, and if so why?
- What evidence is there to reveal the development of past bottom current formation?
- Is there regional variation in the acoustic basement topography and if so, what can it reveal about the tectonic history of the Amundsen Sea basin development?

In this work, we use the generally accepted partition of three major Antarctic climate regimes in the Cenozoic (e.g. Lindeque et al., 2013):

- The pre-glacial climate regime is defined as being relatively warm and predominantly ice sheet free, except for smaller glaciers in high altitudes, and open-ocean conditions without major sea-ice cover. The transitional regime refers to a colder climate, analogous to the present day Alpine climates with regional scale ice sheets undergoing multiple advance/retreat cycles to and from coastal areas and inner shelves, and only perennial regional sea-ice cover. The full glacial climate regime describes a climate similar to present day West Antarctica, in which continental scale ice sheets expanded to the coastal regions and grounded on the middle to outer shelf during glacial maxima.

2. Tectonic and palaeoenvironmental setting

2.1. Rifting and basin development

The Amundsen Sea basin formation started in the late Cretaceous during continental rifting, before initial oceanic crust formed, when...
the Campbell Plateau and Chatham Rise of New Zealand separated from the Pacific margin of West Antarctica at western Thurston Island and Marie Byrd Land, between 90 and 83 Ma (Fig. 2; Table S1; e.g. Eagles et al., 2004a; Wobbe et al., 2012). The basin opening propagated from east to west. In the east, multi-stage wide-mode rifting dominated in the Amundsen Sea Embayment (Wobbe et al., 2012; Kalberg and Gohl, 2014), whilst narrow-mode extension took place in the eastern Ross Sea and off western Marie Byrd Land (Luyendyk et al., 2003; Davey and De Santis, 2006; Jordan et al., 2010).

Between about 80 and 62 Ma (C34y and c27o in Fig. 2), the Bellingshausen Plate moved independently from the Antarctic Plate east of the Proto-Antipodes Fracture Zone (FZ) (Fig. 2; Table S1; Larter et al., 2002; Eagles et al., 2004a,b; Wobbe et al., 2012). The southern margin of the Bellingshausen Plate formed a diffuse, possibly distributed plate boundary zone crossing the Amundsen Sea Embayment (Wobbe et al., 2012; Kalberg and Gohl, 2014), whilst narrow-mode extension took place in the eastern Ross Sea and off western Marie Byrd Land (Luyendyk et al., 2003; Davey and De Santis, 2006; Jordan et al., 2010).

2.2. Palaeoenvironmental development

A comprehensive chronology of palaeoclimatic events is compiled in Table S1. The following events are most relevant and intertwined with sedimentation processes along the Pacific margin of West Antarctica:

The Cretaceous–Palaeocene climate is thought to have been warm and relatively ice-free. Most of West Antarctica was elevated above sea level (Wilson et al., 2012) and the ocean gateways between East Antarctica and Australia (Tasmanian gateway), and between South America and the Antarctic Peninsula (Drake Passage/Scotia Sea gateway) were still closed (e.g. Bijl et al., 2013; Eagles and Jokat, 2014). A proto Amundsen Sea–Ross Sea gyre may have developed towards the late Palaeocene (Huber et al., 2004) when surface waters cooled sufficiently to initiate bottom water circulation (Uenzelmann-Neben and Gohl, 2012).

In the Late Eocene, shallow to deep water passages developed in the Tasmanian and Drake Passage/Scotia Sea gateways, and Antarctica's climate became increasingly cooler. By the Oligocene–Miocene stages of the basin development comprised on-going passive margin rifting.

Fig. 2. Isochron compilation map for the Amundsen Sea and Ross Sea region modified after Wobbe et al. (2012). Isochrons are labelled in red and appended with o—old or y—young.

Dashed lines are fracture zones (FZ) that cross the AS–RS transect (thick red/black line). Selected tie lines used in this work are annotated and summarised in Table S2. Main seismic surveys are AWI-2010—red; TAN-2006—green; AWI-2006—yellow; AWI-1994—navy; selected lines from BGR80—labelled. Supporting age information was drawn from boreholes: yellow circle—DSDP Leg 28 Sites 270 to 274; green circle—DSDP Leg 35 Site 324; and shown on the inset: blue circle—ODP Leg 178 Site 1095; purple circle—IODP Leg 318 Sites U1356 and U1359. Velocity data for sediment thickness estimates were taken from: selected TAN-2006 seconubuys—orange squares (supp. Fig. S2); interval velocities from the AS–RS transect seismic data (Tables S3, S4; Fig. S2); and ocean-bottom seismograph (OBS) measurements on line AWI-20060200 (Kalberg and Gohl, 2014). MBS—Marie Byrd Seamounts; ASE—Amundsen Sea Embayment. Red arrows—as in Fig. 1. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)
during this time, favouring the formation of elongate sediment drifts in the eastern sector (Uenzelmann-Neben and Gohl, 2012).

U–Pb geochronology on detrital apatite grains in the Cape Roberts Projects (CRP) drill cores of the western Ross Sea provides indications for an Oligocene ice sheet with occasional grounding and expansion onto the Ross Sea shelf (Olivetti et al., 2015). They argue that it is a record from the WAIS, although we think it is a combined record from both the EAS and the WAIS. Compared to 70 Ma, the 34 Ma Oligocene West Antarctica topography has been significantly eroded and a great fraction of the present-day Ross Sea shelf submerged (Wilson et al., 2012, 2013). The Eocene–Oligocene transition is marked by high amplitude regional unconformities, visible in the Ross Sea and Wilkes Land margin seismic images (De Santis et al., 1999; Escutia et al., 2005). However, the extent of the Eocene–Oligocene transition is not well constrained on the Antarctic continental margin due to the paucity of drill samples. In Wilson et al. (2012), for example, a large range of WAIS extension at the Eocene–Oligocene boundary is given.

In the Oligocene, the Southern Ocean gateways opened deep-water passages and the Antarctic Circumpolar Current (ACC) developed. This period is marked by an increased amount of sediment deposition, e.g. seismic units RSS-3 and RSS-4 in the Ross Sea (Brancolini et al., 1995; De Santis et al., 1999; Brancolini and Leitchenkov, 2010), WL-S5 and WL-S4 off Wilkes Land margin (De Santis et al., 2003; Escutia et al., 2005, 2011) and Unit II in the eastern Amundsen Sea (Uenzelmann-Neben and Gohl, 2012; Table S1). Elongate drifts and channel–levee drift deposits in the eastern Amundsen Sea increased in response to intensified bottom water circulation (Uenzelmann-Neben and Gohl, 2012). In addition, sediment transport across the shelf and slope increased due to the numerous advance cycles of the continental-scale ice sheets expanding towards the coast, and partially into the inner shelves (Brancolini et al., 1995; De Santis et al., 1999; Haywood et al., 2009; Brancolini and Leitchenkov, 2010; Uenzelmann-Neben and Gohl, 2012; Gohl et al., 2013).

The mid-Miocene to Quaternary period is dominated by the intensified and more frequent advance and retreat cycles of the Antarctic ice sheets onto the middle and outer shelves (e.g. Miller et al., 2008), and the building of progradational shelf sequences (De Santis et al., 1999; Brancolini and Leitchenkov, 2010; Gohl et al., 2013). During this period, the ACC and Antarctic Bottom Water (AABW) circulation developed fully (Uenzelmann-Neben and Gohl, 2012; Table S1). The Antarctic topography gradually changed, especially in the West Antarctic and in the Ross Sea, where much of the region became submerged (Wilson et al., 2012, 2013) to approach the present day topography (Fretwell et al., 2013). Eleven seismic lines were acquired with a 30 ms window, 100 ms operator length and 0.1% white noise, but introduced ringing artefacts in key parts of the deeper sequences and across major sequence boundaries after stacking. It was therefore only applied during velocity analyses and in selected cases, to more clearly image unconformities, drifts feature geometry or the seismic unit boundaries. Static corrections of up to 2 ms were applied to profiles AWI-20100108 and -109, because the 600 m streamer was not equipped with depth-control devices due to sea-ice. The ghost on these two profiles was effectively removed with the application of an F-K filter.

Velocity analysis was applied at every 50th CDP and intervals were picked to match sequence boundary positions closely. After normal-move-out (NMO) correction and spherical divergence adjustment, a Kirchhoff time-migration was applied to the stacked section. A maximum migration angle of 65° and 90% of the NMO velocities were used.

The interpretation of the seismic images is based on the change in seismic characteristics such as reflectivity and variation of internal reflection geometries. No additional filters or amplitude scaling were applied to most sections, so that true amplitudes were largely preserved. In one case, Automatic Gain Control (AGC) was partially applied on line AWI-20100117 between CDPs 15000 and 22988 to image the basement more clearly.

3.1. Seismic reflection data

The multichannel seismic profiles AWI-2010xxxx used to construct the Amundsen Sea to Ross Sea (AS–RS) transect were acquired during the RV Polarstern expedition ANT-XXVI/3 in 2010 (Figs. 1, 2; Table S2; Gohl, 2010). Exceptionally favourable sea-ice conditions allowed the acquisition of a continuous ~2000 km long profile along the 72°S latitude. The TAN-2006-xx lines (Figs. 1, 2; Table S2) were acquired during the RV Tangaroa expedition in 2006 (hereafter referred to as the TAN-2006 survey) and complement the AWI lines (hereafter referred to as the AWI-2010 survey). This combined dataset links pre-existing seismic data of the continental rise and shelf of the eastern Amundsen Sea (e.g. Nitsche et al., 1997, 2000; Scheuer et al., 2006; Uenzelmann-Neben and Gohl, 2012, 2014; Gohl et al., 2013) with data from the Ross Sea rise and shelf (e.g. Anderson and Bartek, 1992; ANTOSTRAT, 1995; Brancolini et al., 1995, 1999; De Santis et al., 1995, 1999; Brancolini and Leitchenkov, 2010). The acquisition and processing parameters of the AWI-2010 and the TAN-2006 survey data are described below and listed in Table S2 for lines used in this work.

3.1.1. The AWI-2010 survey of RV Polarstern

Most of the AWI-2010 profiles (Figs. 1, 2; Table S2) were acquired using a 3000 m long SERCEL Sentinel solid streamer with 240 active channels, towed at a nominal water-depth of 10 m. Profiles AWI-20100108 and AWI-20100109 are the exceptions, where the long streamer was replaced with an older, analogue 600 m long Prakla streamer with 96 channels due to adverse sea-ice conditions. These two profiles are shorter (34 km and 54 km respectively, Table S2) and have a minimal impact on the overall transect data quality or continuity. In all cases, the seismic source consisted of two or three GI-Guns of 4.8–7.2 L (300–450 in³) total volume, fixed to a steel frame and towed in about 3 m water-depth (Table S2). Shots were fired at a 12 s interval in true GI-mode and with 19.5 MPa nominal operational pressure. Data were sampled at a 1 ms interval and recorded for 10 s. Restrictions by the German Federal Environmental Agency prevented the use of a larger airgun source needed to image the basement beneath very thick sedimentary layers. Consequently, basement was poorly imaged in some cases, especially in the easternmost part of profile AWI-20100117.

Seismic processing followed a standard approach using the processing software FOCUS/ECHOS (by Paradigm). After trace editing, a high-pass filter tapered at 5–10 Hz was applied to allow higher frequencies and maximum resolution of the data. Predictive and spiking deconvolution were tested to remove the ghost and internal multiples prior to velocity analyses. Spiking deconvolution effectively removed the ghost when using a 30 ms window, 100 ms operator length and 0.1% white noise, but introduced ringing artefacts in key parts of the deeper sequences and across major sequence boundaries after stacking. It was therefore only applied during velocity analyses and in selected cases, to more clearly image unconformities, drifts feature geometry or the seismic unit boundaries. Static corrections of up to 2 ms were applied to profiles AWI-20100108 and -109, because the 600 m streamer was not equipped with depth-control devices due to sea-ice. The ghost on these two profiles was effectively removed with the application of an F-K filter.

Velocity analysis was applied at every 50th CDP and intervals were picked to match sequence boundary horizons closely. After normal-move-out (NMO) correction and spherical divergence adjustment, a Kirchhoff time-migration was applied to the stacked section. A maximum migration angle of 65° and 90% of the NMO velocities were used.

The interpretation of the seismic images is based on the change in seismic characteristics such as reflectivity and variation of internal reflection geometries. No additional filters or amplitude scaling were applied to most sections, so that true amplitudes were largely preserved. In one case, Automatic Gain Control (AGC) was partially applied on line AWI-20100117 between CDPs 15000 and 22988 to image the basement more clearly.

3.1.2. The TAN-2006 survey of RV Tangaroa

The RV Tangaroa acquired 3409 km of multichannel seismic data and velocity data from 16 sonobuoys during the TAN-2006 marine geophysical survey in the Ross Sea, conducted by NIWA and GNS Science of New Zealand (Fig. 1; Table S2). Eleven seismic lines were acquired with a 300 m streamer with 48 channels (near offset ~129 m and far-offset ~422 m) towed at a nominal depth of 10 m. The source consisted of 4 GI-Guns (9.8 L) towed at 5 m depth and fired in true GI-mode, resulting in a shot spacing of 50 m. Recording was done for 12 s at a sample interval of 1 ms and a 10 Hz low-cut filter applied during acquisition.

The lines vary in geometry and acquisition parameters due to sea-ice conditions, which introduced additional noise and necessitated processing in parts. This particularly affected line TAN-2006-04, where noise...
from sea-ice conditions and damage caused by the snagging and parting of the streamer by an iceberg, reduced sections of the 48 channel streamer to 16 channels on the whole array (Table S2).

Preliminary seismic data processing to final migrated stacks was completed on-board. Zero traces were inserted for missing shots to account for <10 km wide gaps and jumps in the line geometry. Line gaps of >10 km were processed as separate lines (Fig. 2; Table S2). A 25 m CMP spacing ensured a good signal-to-noise (S/N) ratio. A standard final processing flow was used with GLOBE CLARITAS seismic processing package and included: resampling the data down to 2 ms, applying a 225–250 Hz low pass filter, standard muting of noisy traces and data spikes before velocity analyses, followed by stacking. After stacking, a band-pass filter of 4–24 Hz was applied, followed by finite difference migration with a maximum migration dip-angle of 45°, and 100% of the NMO velocities were used. The finite difference post-stack migration produced a clearer true-dip image. The velocity models were constrained by records from sonobuoys deployed along some profiles.

3.2. Interpretation approach

3.2.1. Type section

Prior to interpreting the relevant AWI-2010 and TAN-2006 lines (Table S2) or linking it with the interpretation of previous surveys, we identified the prominent horizons, unit, or sequence boundaries observed in the processed seismic images alone. To achieve this unbiased approach, we selected a representative type section in the central part of the Amundsen Sea basin on line AWI-20100110 (Figs. 1, 2). It was chosen at this basin-centre location, under the assumption that the sediments were likely to be the thickest and least disturbed, and would as such present the best possible opportunity to identify all the potential basin-wide units (Fig. 3; see Fig. 1 for location). In mapping the initial prominent first and second order sequence boundary horizons, we looked for variations in seismic facies characteristics such as reflectivity, reflection continuity and internal geometry (Fig. 3).

Fig. 3. Type section of the seismic reflection units AS-1 to AS-10, observed in the Amundsen Sea rise basin (Fig. 1 for location), with unit characteristics on the right. The units are separated by unconformities or boundaries of abrupt reflection characteristic change, labelled AS-u1 to AS-u10. The pre-glacial (PG), transitional (T) and full glacial (FG) regime sequences are separated by the base reflector discontinuities uB-PG, uPG-T and uT-FG, respectively. Black lines in the basement indicate block faulting associated with the Endeavour fracture zone (EFZ).
3.2.2 Seismic characterisation

Subsequently, we compared our initial horizons in the type section interpretation (Fig. 3) to studies in other deep sea basins around Antarctica that used changes in seismic facies and internal geometry to identify pre-glacial to glacial sedimentary sequences in seismic reflection data, some of which are tied to drilling information (e.g. eastern Amundsen Sea: Nitsche et al., 1997, 2000; Uenzelmann-Neben and Gohl, 2012; Bellingshausen Sea: Scheuer et al., 2006; Weddell Sea: Maldonado et al., 2006; Lindeque et al., 2013; Riser Larsen Sea: Leitchenkov et al., 2008; Cosmonaut Sea: Leitchenkov et al., 2007a; Mawson Sea: Leitchenkov et al., 2007b; Wilkes Land: De Santis et al., 2003; Escutia et al., 2011; Ross Sea: De Santis et al., 1995; Brancolini and Leitchenkov, 2010). We used these collective interpretations and seismic horizon characteristics of the pre-glacial to glacial units as analogues to assess if similar packages and depositional processes apply to the AS–RS transect and relevant TAN-2006 lines. We base our
comparison on internal geometry such as drifts (i.e. mound shaped asymmetric ridges with internal strata pinching out on the steepest side of the mound) and channel–levees (i.e. mound shaped asymmetric ridges aggregated through repeated overspill events from the channel axis created by turbidity flows). The assumption is that if these sequences have already been identified on a circum-Antarctic scale in several of the deep-sea basins, two of these locations being on either side of the study area, it would be very likely that similar sequences could be present in the AS–RS transect and connecting lines.

In some parts of the eastern Amundsen Sea and around the Marie Byrd Seamounts the FG and T sequences contain distinctive mass sediment transport bodies, such as sheeted or mounded drifts, interpreted to have formed by contour and coast parallel bottom currents (e.g. Uenzelmann-Neben and Gohl, 2012). Hence identifying drifts in the seismic images of the AS–RS transect, could similarly provide an indication of the depositional processes and palaeo-bottom current activity in the central Amundsen Sea basin.

3.2.3. Horizon mapping

The type section interpretation (Fig. 3) was expanded into the regional basin context along our 2000 km long AS–RS transect (Fig. 1). The AS–RS transect was constructed from 8 profiles of continuous seismic data from the AWI-2010 survey (Fig. 4a). To test the initial unit boundary horizons interpretation in the type section (Fig. 3) within the AS–RS transect, not only the marker horizons but also every internal reflection, were traced up to the farthest western and eastern flanks of the basin and for the full extent of the transect dataset. In doing so, we were able to identify low-angle onlap or downlap unconformities and depositional angle changes (Figs. 4b, 5; Fig. S1a, b). Lastly we mapped the basement, the corresponding fracture zones (Figs. 2, 4), and larger internal structures such as sediment drift bodies (Figs. 4b, 6; Fig. S1b). Independent analyses of the central transect placed us in a position to assess if the unit boundaries and preliminary sequence boundaries correlate with the existing interpretations in the Ross Sea and eastern Amundsen Sea (Fig. 1).

3.3. Integration with pre-existing seismic data

3.3.1. Seismic reflection surveys

There are three points where the AS–RS transect ties into the regional seismic survey network. The first is at the western end of the AS–RS transect, linking the Amundsen Sea seismic profile with that of the eastern Ross Sea (Fig. 2). Lines AWI-20100107 and AWI-20100108 connect to line TAN-2006-06a, whereas line TAN-2006-06c.
crosses line AWI-20100109 (Figs. 1, 2, 7). Since the TAN-2006 dataset has no previously published interpretation, the AS–RS transect horizons were extended onto these lines.

The second key seismic stratigraphic link is the Ross Sea shelf-slope-rise connection between the AS–RS transect interpretation and the interpretation of the BGR80 seismic lines (Hinz and Block, 1985; ANTOSTRAT, 1995), in particular, line BGR80-005 connecting to TAN-2006-06 (Figs. 2, 8; Table S2). Line ATC82B-203 (Fig. 2; Wannesson et al., 1985) provided additional information on the shelf break where the reflections are obscured by multiples or unclear. The BGR80-005 interpretation was used as published (Hinz and Block, 1985; ANTOSTRAT, 1995). The AS-ux unconformities were traced up the slope, across the shelf along lines BGR80-003 and BGR80-004 (Fig. 2; Table S2) and to line PD-30, interpretation used as published (Anderson and Bartek, 1992).

The third key seismic stratigraphic link lies in the eastern Amundsen Sea, at the eastern end of the AS–RS transect (Figs. 1, 2). At this intersection, profile AWI-20100117 of the AS–RS transect crosses lines AWI-94050 and AWI-94054 (Fig. 9, Table S2; Gohl et al., 1997; Nitsche et al., 1997, 2000; Uenzelmann-Neben and Gohl, 2012), linked to line AWI-20060200 (Table S2; Uenzelmann-Neben and Gohl, 2012; Kalberg and Gohl, 2014) and to AWI-20060023, crossing the eastern Amundsen Sea drift deposits (Uenzelmann-Neben and Gohl, 2012, 2014). These lines connect farther east to profiles extending to the Bellingshausen Sea (Nitsche et al., 1997, 2000; Scheuer et al., 2006). The AWI-94050 and AWI-94054 horizons and interpretations were used as published (Uenzelmann-Neben and Gohl, 2012) and match the proposed central Amundsen Sea basin horizons (Figs. 3, 4; Fig. S1). Line AWI-20060200 was newly interpreted, because a detailed seismic horizon interpretation is as yet unpublished.

### 3.3.2. Borehole information

Since boreholes are absent from the Amundsen Sea continental rise and shelves, we used the borehole controlled age model from the eastern Ross Sea outer shelf, derived from Deep Sea Drilling Project (DSDP) Leg 28 Sites 270–272 (Figs. 1, 2; Hayes et al., 1975; Bart, 2003). This age model serves as a guide to estimate the age ranges of the interpreted slope and deep-sea horizons in the AS–RS transect via a series of seamlessly linked profiles and extended horizons (Figs. 2, 4–9; Fig. S2; Table S2 and references therein).

DSDP Leg 28 Sites 273 and 274 have been correlated to the ANTOSTRAT seismic sequences (Brancolini et al., 1995) via a series of acquired and archive seismic reflection tie lines (Fig. 2 in Granot et al., 2010). Good data continuity in seismic lines BGR-001, 13 and 14 (Fig. 4 in Granot et al., 2010) ties the shelf-slope correlation of the mid-Miocene RSU4 unconformity to DSDP Leg 28 Site 273 located on the shelf (Figs. 1 and 2), but the shelf-slope correlation of the older, prominent unconformity RSU6 was not possible due to seafloor...
multiple. Line 9, however, ties the Eocene/Oligocene transition RSU6 unconformity to DSDP Leg 28 Site 274 (near shot point 450, Fig. 5 in Granot et al., 2010). Since DSDP Leg 28 Site 274 is located on the lower slope/continental rise and the seismic facies characteristics of their deeper lines (e.g. line 4a, shot point 8500 – 9000, Fig. 10 in Granot et al., 2010) appear similar to those observed in the AS–RS transect, we also used the Granot et al. (2010) chrono-stratigraphic model as a guide for our interpretation.

The nearest other relevant deep borehole on the continental rise, International Ocean Drilling Program (IODP) Leg 318, Site U1356 lies off the Wilkes Land margin, approximately 1300 km to the west of the most western end of our AS–RS transect (Fig. 2 inset; Escutia et al., 2011). We used the seismic reflectivity characteristics of the dated and biostratigraphically analysed pre-glacial, transitional and full-glacial sequences as an analogue to place observations of the AS–RS transect in a provisional chrono-stratigraphic framework.

To the east of the Amundsen Sea, the nearest borehole DSDP Leg 35 Site 324 (e.g. Tucholke et al., 1976) is located in the Bellingshausen Sea, about 1000 km from the eastern end of the AS–RS transect (Fig. 2). This site only provides information for the uppermost 200 m Pleistocene and Pliocene sediments. DSDP Leg 35 Site 323 lies on the abyssal plain and contains predominantly pelagic deposits. Neither of these two drill sites was considered useful for the objectives of this work.

Farther to the east on the Antarctic Peninsula margin, Ocean Drilling Project (ODP) Leg 178 Site 1095 is more applicable (Fig. 2 inset; Barker et al., 2002). Sequences of mid-Miocene age in a sediment drift deposit were drilled, dated and mapped (e.g. Iwai et al., 2002; Uenzelmann-Neben, 2006). Although the drill site lies more than 2000 km from the eastern end of our AS–RS transect, it was used in Scheuer et al. (2006) and Uenzelmann-Neben and Gohl (2012) for the horizon correlation and age model verification in the Bellingshausen Sea (Figs. 1, 2), and is therefore considered as supporting information for the interpretation of our transect.

3.3.3. Seismic velocity measurements for sediment thickness estimates

Sediment thickness and depth sections were constructed from 12 points along the AS–RS seismic transect (Figs. 1, 4; Figs. S1, S2). The points were chosen to represent trends of changes in sediment thickness and basin geometry and are therefore irregularly spaced. Due to distant borehole information and absent seismic refraction data in the central Amundsen Sea basin, we depended on interval velocities derived from RMS stacking velocities for time-to-depth conversions (Fig. S2; Table S3). Supporting velocity information was extracted from sonobuoy 5 and 8 collected on the eastern Ross Sea rise during the TAN-2006 expedition (Fig. 2) and compared to a seismic refraction model from ocean-bottom seismometer (OBS) data along profile AWI-20060200 east of the AS–RS transect (Fig. 2; Kalberg and Gohl, 2014). The interval velocities used are consistent with the sonobuoy velocities and compare well to the refraction velocities.

3.4. Age model and sedimentation rate estimates

Oceanic basement age control of the central Amundsen Sea was obtained from the magnetic seafloor spreading anomalies of Wobbe et al. (2012; Fig. 2). The top-of-basement reflection in each of the 12 points used in the sediment thickness and depth sections (Fig. 4; Fig. S1), were assigned the age of the closest isochron, or the age was linearly interpolated between the two closest magnetic spreading anomalies. This age was then used to place the basement horizon in each of the 12 sections at the corresponding age position on the geological time scale, and the sections vertically stretched so that seafloor matched present time (Fig. 10). The first and last occurrence of each sequence, across all 12 points, were projected back onto the time scale in order to estimate the relative depositional time span (thin colour bars on the left in Fig. 10). The deduced age range for each sequence was subsequently used in conjunction
with the 12 point-based sediment thickness estimates (Fig. S2; Table S3) to calculate the sedimentation rates (Table S4).

4. Results

4.1. Seismic interpretation and horizon characteristics

4.1.1. Type section and nomenclature definition

Prominent horizons or unconformities were assigned up-sequence as AS-u1, AS-u2 etc., where AS stands for Amundsen Sea and -ux for the unconformity number. Similarly, units between two boundary horizons were assigned AS-1, AS-2 etc. and numbered up-sequence (Figs. 3, 4 and Fig. S1).

We started the interpretation with the delineation of the crystalline basement horizon and observed faulting in the type section (Fig. 3). Continuing up from the top of basement, the first high amplitude, continuous reflection was assigned as AS-u1, followed by AS-u2. Both reflectors stood out prominently against the low amplitude to transparent background. The next unit transitions were identified at reflections AS-u3, AS-u4, and AS-u5, as amplitudes increased consistently up-sequence, and the internal unit reflection geometry changed to become progressively more closely spaced (Fig. 5; Fig. S1). AS-u6 is mapped along a low angle unconformity with onlap up-dip (6.2–6.4 s and shot numbers 17000–18000 in Fig. 3; Fig. S1) and also marks a strong amplitude increase of all the units above this horizon. AS-u7 and AS-u8 are high-amplitude unconformities as well, and -u9 was chosen at the base of sediment drifts, but above what seems to be a truncated drift (at 5.75 s and shot numbers 15000–16500 in Figs. 3, 4 and Fig. S1).

To identify the pre-glacial to glacial sequences and characteristics, we compared our type section to other type sections of deep-sea basins around Antarctica. These studies typically describe the pre-glacial (PG) seismic sequences as the initial series of discontinuous reflections from low amplitudes to transparent reflectivity, occurring just above the basement. Internal reflections are predominantly parallel to each other and the basement topography. If plotted without filtering the display, this sequence appears light grey to white on a greyscale seismic image e.g. units below WL-u3 in Escutia et al. (2011), unit WL-S3 in De Santis et al. (2003), Unit I in Uenzelmann-Neben and Gohl (2012) and units WS-S1 to WS-S3 in Lindeque et al. (2013). The full glacial (FG) sequences were mostly characterised by the highest amplitude reflections and strong lateral continuity. Internal reflections
are parallel and very closely spaced, with a collective dark grey to partly black appearance on a greyscale seismic image without display filters e.g. WL-S6 to WL-S8 in Escutia et al. (2011) and De Santis et al. (2003), Units III and IV in Uenzelmann-Neben and Gohl (2012; Fig. 9) and unit WS-S5 to WS-S7 in Lindoque et al. (2013).

Using the seismic characteristics identified in these other studies as a guide, we visually identified the UPG-T and UT-FG boundaries in our type section, interpreted as separating the 1st order pre-glacial, transitional and full-glacial sequences (annotated in Fig. 3). In our central Amundsen Sea type section (Fig. 3), we describe the pre-glacial, transitional and full glacial sequences as follows:

- The pre-glacial, PG sequence, displays a low to transparent reflectivity and lack of internal structure, which indicates well-sorted sediments and relatively uniform sedimentation processes over long times, probably associated with the greenhouse period of Antarctica.

- The transitional, T sequence, has medium to low amplitude reflectivity and reflections are discontinuous, parallel and evenly spaced. The higher amplitudes are probably due to the interlayering of strata with different grain sizes (e.g. turbidites) or alternation of well and poor sorted sediment strata (e.g. diamictites-debris flows and hemipelagites) associated with increased down-slope gravity-driven, and along-slope bottom current driven sediment transport processes. It is important to note that these interlayering sequences are widely distributed along the AS–RS transect, which discounts the possibility that they could be point sources causing diffractions.

- The full glacial, FG sequence, shows the highest continuous reflectivity and the reflections are parallel and very closely spaced. Drifts occur in units AS-9 and AS-10. These deposits are interpreted as being indicative of the continued supply of predominantly poorly sorted material through mass down-slope sediment transport processes as grounded ice sheets repeatedly advanced across the shelf, followed by redistribution along the continental rise by strong bottom-current activity. Sediment drift building by contourite currents is a major characteristic observed for the FG phase (e.g. Uenzelmann-Neben, 2006; Rebesco et al., 2014; Figs. 4–6; Fig. S1).

Although these type sections in the Weddell Sea, eastern Amundsen Sea and Wilkes Land studies sections are far apart from each other, and from this study’s central Amundsen Sea type section, the visual resemblances of the PG, T and FG sequences between the type sections are striking when placed directly next to each other.

4.1.2. Seismic characterisation and horizon stratigraphy

We used the nomenclature and marker horizons from our type section (Fig. 3) to map the remaining full length of the AS–RS transect and present new basin-wide seismic stratigraphy for the Amundsen Sea deep-sea basin (Figs. 4, 5; Fig. S1).

The acoustic basement topography is irregular and rugged along the western basin flank (line AWI-20100107) as well as on the eastern Amundsen Sea flank (line AWI-20100117). The basin geometry deepens towards middle of the transect, forming a classic ‘saucer shape’ and a valley in the centre (line km 1000 to 1200 in Fig. 4; Fig S1). To the west of the centre point (line km 700 to 1100), typical block faulting was mapped. The fracture zones predicted in the magnetic data (Wobbe et al., 2012) match major basement faults observed in the AS–RS transect (Kohiko-, Pahemo- and Endeavour FZs in Fig. 4; Fig S1). In the region east of the centre point (line km 1100 to ~1650) and up to the first seamount (line km 1700 in Figs. 4, 5), the basement topography is smoother and without faulting. The slope angle of this anomalous part is ~1° (Fig. 5).

The seismic reflectivity characteristics of units AS-1 to AS-10 in the AS–RS transect (Fig. 4; Fig. S1) match the characteristics for each unit described in the type section (Fig. 3) well and lateral variation is negligible. Marker horizons AS-u1 to AS-u9 are laterally continuous and easy to trace along the ~2000 km AS–RS transect distance. Every internal reflection was mapped over the full transect length and low-angle unconformities and onlaps or downlaps identified, which increased confidence in deciding on the PG, T and FG sequence boundary horizons (Figs. 4b, 5b; Fig. S1b).

In general, the horizons are mostly parallel to the general basin geometry and top-of-basement trend, except above the anomalous
part of the basin (line km 1100 to ~1650). This part was mapped in
greater detail (Fig. 5). Up to AS-u3, the reflections are parallel to the
basement slope trend and the internal unit reflections onlap against a
drift and channel–levee structure in AS-3 (Figs. 4, 5; Fig. S1). A
prominent reflection in the middle of AS-3 (6.4–6.8 s, shot 20000–25000 in Fig. 5) pinches up-dip against the large sedimentary
-drift body and is laterally discontinuous. It was therefore interpreted
as a sub-unit within AS-3.

Above AS-u3, the reflections become increasingly angular compared
to the reflections below AS-u3, for example, the thinner red lines in unit
AS-4 and AS-5 (Figs. 4b, 5b; Fig. S1). The internal unconformities are at
their largest internal angles up to AS-u6; this is especially visible in the
internal geometry of unit AS-6. Above AS-u6, the horizons trend chang-
es to progressively more horizontal and parallel towards the top units,
with complex internal geometry in the youngest unit, AS-u10 (Fig. 5b,
between seafloor (sflr) and AS-u9; Fig. S1b).

After mapping all the marker horizons and every internal reflection,
more low angle unconformities were noticed, which confirmed the ini-
tial marker horizon decisions chosen in the type section interpretation
(Sections 4.1.1 and 3.2.1). It also cleared up the earlier ambiguity around
AS-u3 (see unconformities in Fig. 5b onlapping against AS-u3; Fig. S1ab).

The resulting PG sequence, units AS-1, AS-2 and AS-3, are bound by
the basement reflection (bsmnt) below and the AS-u3 discontinuity above
(Figs. 4, 5). AS-u3 or uPG-T was picked as the top boundary of this 1st order T sequence be-
cause it marks a sharp change into the highest amplitude, finely layered
reflective package above (Fig. 4a). Again, this is consistent with the de-
scriptions of the pre-glacial to transitional unit progression, seen in type
sections of the other Antarctic margin deep-sea basins. In addition, this
horizon is a pronounced unconformable high amplitude reflector with
reflections terminating against it and it delineates a prominent change in
the reflection dip angles on either side (Fig. 5), which collectively
may also indicate a tectonic uplift change. We interpret this sequence
and associated seismic characteristics to indicate a rapid increase in
the supply of large volumes of poorly sorted sediments, pushed across
the shelf and spilled down the slope and into the abyssal plain, in a re-
response to the first ice sheet advancements onto the shelf.

The FG sequence includes units AS-7, AS-8, AS-9 and AS-10. These
units show more complex internal second and third order unconfor-
mities towards the top, and several internal drift structures (Figs. 3–6;
Fig. S1). Increased occurrence of the complex internal drift structures
represents changes in the depositional processes and bottom water
circulation. This is consistent with the interpretations of Uenzelmann-
Neben and Gohl (2012) in the adjacent eastern Amundsen Sea rise
basin. The sharp transition at AS-u6 is interpreted to mark the onset of
the full glacial polar climate regime and the establishment of a modern
oscillatory polar ice sheet, consistent with drilling information from the
IODP Leg 318, Site U1356 off the Wilkes Land margin (Escutia et al.,
2011; Fig. 2 inset).

4.1.3. Drifts

Distinct sediment bodies are observed in the AS–RS transect (grey
wedge-shaped structures in Figs. 4–6; Fig. S1) There are too many

<table>
<thead>
<tr>
<th>Wilkes Land Basin</th>
<th>Ross Sea shelf and slope</th>
<th>central Amundsen Sea basin</th>
<th>far eastern Amundsen Sea basin</th>
</tr>
</thead>
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<td>AS-u9</td>
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<td>RSU11</td>
<td>AS-u5</td>
<td>Unit I</td>
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Fig. 11. Seismic stratigraphy correlation chart that illustrates the links of the boundary horizons and the seismic units in this work, to previous studies of the Ross Sea shelf and eastern Amundsen Sea basin. Thick black lines—lower boundary of the proposed full glacial and transitional units; red numbers—ages of the horizons, showing the lateral variation across the basin from west to east. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)
Fig. 12. Model showing the changes in plate-tectonic motion, pre-glacial to glacial sediment deposition, ice sheet development and ocean circulation in the Southern Pacific, and greater Amundsen Sea region since the Paleocene. The corresponding sediment deposition for each regime is illustrated in the AS–RS transect diagram below. Plate rotation adapted from Wobbe et al. (2012). Thick dashed black lines—MOR—mid-oceanic ridge, black lines—fracture zones, EFZ—Endeavour Fracture Zone, PAC—Pacific Plate, ANT—Antarctic Plate, BEL—Bellinghausen Plate, COB—continent–ocean boundary. The AS–RS transect is shown as a red line, linked with the TAN-2006-06 seismic profile (green line). Observed sedimentary drifts and mounds are indicated by ovals, and the arrows depict the inferred sediment supply and ocean circulation. Red and blue polygons illustrate the schematic extent of the East Antarctic Ice sheet (EAIS), West Antarctic Ice sheet (WAIS) and Antarctic Peninsula Ice sheet (APIS) expanding to the continental shelf during the transitional and full-glacial regimes. ChP—Challenger Plateau, ChR—Chatham Rise, CaP—Campbell Plateau, BT—Bounty Trough, NNZ—North Island of New Zealand, SNZ—South Island, AUS—Australia, WL—Wilkes Land, MBL—Mary Byrd Land, AP—Antarctic Peninsula. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)
mounds, drifts and channel–levee systems to show high-resolution images of their internal geometry or to give a detailed interpretation of each one. This would by far exceed the scope and size of the paper, and were purposefully left for future work. Here, only the main ones are briefly described in the context of the units within which they occur.

The PG sequence contains two sedimentary bodies. The first feature is interpreted as a buried giant channel–levee system on the eastern flank of the basin, indicative of large turbidite flows possibly related to glacial conditions on land (Escuita et al., 2000). However, such detailed classifications remain ambiguous in this case as the features are only imaged with a single 2D profile. The giant channel–levee system occurs near the basement and within the older pre-glacial unit AS-3, with its base on AS-2 (Fig. 5). It is ~3000 m long and almost 400 m high at the highest point (5.9–6.3 s TWT at line km 1475 in Fig. 5 and Fig. S1). The internal seismic reflections are distinct and continuous, forming a parallel lamination that follows the overall drift geomorphology, shown in more detail in Fig. 5b. The stoss side pinches out to the west.

The second PG sedimentary body is a smaller, approximately 80 km wide mound or drift, just to the west of the basin centre near the Pahemo FZ (line km 750 in Fig. 4b). It lies directly on the 68–65 Ma basement and is constrained to AS-2 (Fig. 6c, line AWI-20100110, shot 6804–9883). This feature is interpreted as a mound or drift, since it appears internally stratified. It could also be interpreted as a slope-slide block, or gas-fluid intrusion that folded overlying sediments. However, the resolution of this feature is limited and it is difficult to observe the actual internal stratification geometry; hence this single line only allows for an ambiguous or non-unique interpretation.

Smaller drifts or mounds occur in the T sequence and are concentrated on the western flank of the basin (profiles AWI-20100107, -108 and -109 in Figs. 4b, 6b; Fig. S1b). The base seems internally stratified and terminates on the eastern part of the base lines directly on the 63–65 Ma basement (Fig. 6b, line km 520 onwards, in line AWI-20100110). The tops extend into units AS-5, AS-6 and AS-7 (line km 350–550 in Figs. 4b, 6b; Fig. S1b). The height is estimated as 400–500 m (Fig. 6b, 6.3–5.8 s TWT).

Fig. 6a illustrates an extract of the elongate sediment bodies observed in the FG sequence in line AWI-20100110 (shot 5209–14075 or CMP 7481–19312; Fig. 4). The number of drifts increases from one in AS-8 to two in AS-9, and to several stacked, more widely distributed ones in AS-10 (e.g. at shot 9900 and 5.6–5.9 s TWT; Fig. 6a; Fig. S1b). The base seems amalgamated and terminates on AS-u4, but the eastern part of the base lies directly on the 63–65 Ma basement (Fig. 6b, line km 520 onwards, in line AWI-20100110). The tops extend into units AS-5, AS-6 and AS-7 (line km 350–550 in Figs. 4b, 6b; Fig. S1b). The height is estimated as 400–500 m (Fig. 6b, 6.3–5.8 s TWT).

4.2. Seismic stratigraphic links

4.2.1. Ross Sea rise link

The TAN-2006 and AWI-2010 surveys link seamlessly (Figs. 7, 11) and the horizons from the AS-RS transect were extended into lines TAN-2006-06, -06a,b,c, -04, 04p1, -p2 and -p4 in the Ross Sea (Table S2; Figs. 2, 7 inset, 11). Both lines TAN-2006-06a and -06c show drifts or mounds in their T units as well, and these tie with the mounds seen in lines AWI-20100108 and -109 (Fig. 7). The PG sequence (Fig. 7) pinches out against a basement high in the eastern Ross Sea (see bathymetry image in Fig. 1).

4.2.2. Ross Sea shelf-slope-rise link

The Ross Sea shelf seismic lines ATC82B-203 and BGR80-005 are nearest to line TAN-2006-06 (Figs. 2, 8 inset). The stratigraphy of the Ross Shelf was taken as published from Brancolini et al. (1995, 2010) and the link to line PD90-30 taken from Anderson and Bartek (1992).

The interpretation of DSDP Leg Sites 270, 271 and 272 (Hayes et al., 1975; Bart, 2003) via BGR80-003 and -004 was adopted as is (Figs. 2, 8). Most of the AS-ux horizons on the TAN-2006-06 line could be traced and linked to parts of the interpreted Ross Sea shelf horizons, apart from a few slumping structures on the continental slope (Figs. 8, 11). As the shelf sequences are different from the deep-sea sequences in compositional and depositional processes, we deemed the only age and nomenclature correlation of the horizons of relevance. Seafloor multiples in seismic records from the shelf made some of the correlation ambiguous, but nearby line ATC82B-203 provided useful additional information so that we are fairly confident on the correlation.

The AS-u1 reflector inside the PG sequence may have a similar strong reflectivity character as RSU6 in the eastern Amundsen Sea part of the AS-RS transect, interpreted as the top of volcanics and of Eocene age (Brancolini et al., 1995; Brancolini and Leitchenkov, 2010; Wilson et al., 2013). AS-u1, however, does not display the same ‘volcanic’ reflectivity character throughout. In the centre of the basin near the Endeavour fracture zone, the AS-1 seismic facies character rather suggests a sedimentary unit. AS-u1 and AS-u2 are not observed on the TAN-2006-06 line or western part of the AS-RS transect, but AS-u3 is. AS-u3 is also observed for the length of the AS-RS transect, into the eastern Amundsen Sea. Based on this regional continuation and consistent, strong prominent reflector characteristics (see detailed line drawing of each reflector and related discontinuities in Fig. S1b), AS-u3 rather than AS-u1 or AS-u2, is interpreted to represent the margin wide onset of glaciation process in this deep sea basin.

The horizon RSU6 is interpreted to be comparable to the AS-u3/upG-T sequence boundary, and RSU4 to the AS-u6/upG-T sequence boundary (Figs. 8, 11). Shelf horizon RSU4a appears to correlate with horizon AS-u5. This discontinuity is considered to be part of the transitional sequence as opposed to a first-order sequence boundary, since it terminates against AS-u6 (Figs. 5, 8 and Fig. S1). The suggested correlations of the other RSUx to AS-x horizons are annotated in Fig. 8. We use this slope-shelf interpretation to support our proposed chrono-stratigraphic model for the entire AS-RS transect, because it is currently the best available link to an age-controlled stratigraphy for the entire Ross Sea–Amundsen Sea region (Fig. 11).

Shelf-slope-rise correlations are in general difficult to pursue on most margins because of stacked down-slope transport processes. We selected the highest quality shelf-slope seismic profile nearest to our AS-RS transect whilst being aware of the large uncertainties. It would be desirable, but beyond the scope of this paper, to focus on a more constrained shelf-to-slope stratigraphy correlation of the eastern Ross Sea in a separate study, with more seismic lines and possible future drill sites.

4.2.3. Eastern Amundsen Sea link

Line AWI-20100117 forms the far eastern part of the AS-RS transect and crosses older seismic lines AWI-94054 and AWI-94050 (Table S2; Gohl et al., 1997; Nitsche et al., 1997, 2000; Uenzelmann-Neben and Gohl, 2012) as well as AWI-20060200 (Figs. 2, 9; Table S2; Gohl, 2007; Uenzelmann-Neben and Gohl, 2012; Kalberg and Gohl, 2014). The base of Unit II in Uenzelmann-Neben and Gohl (2012) correlates to the uP-GT/AS-u3 unconformity, whilst the base of Unit III corresponds to the uP-FG/AS-u6 sequence boundary (Figs. 9, 11). All three units occur throughout and the correlation was recognisable without discrepancies or ambiguities (Fig. 11).

4.3. Sediment thicknesses and deposition volumes

AS-1 is the thickest unit in the central part of the basin (~1.3 km; Table S3) at the Endeavour FZ where it in-fills the rugged block faulted basement (Figs. 3, 4; Fig. S1). It is therefore interpreted as a syn- or post rift sequence (line km 1100 in Fig. 4). Little internal faulting is present and AS-1 seems to drape over the basement. However, it thins towards the basin flanks (~133 m thick) and terminates at the Pahemo Fault in
the west, and against the Proto-Antipodes Fault in the east (Fig. 4; Fig. S1). AS-2 follows the trend of AS-1 but is more uniform in thickness in the centre of the basin (-500 m; Table S3; Fig. 4b and Fig. S2) and levels out above the block-faulted basement topography. It terminates against the Kohiku FZ in the east (Fig. 4b and Fig. S2). In contrast, AS-3 is anomalously thick in the east where the levee-drift is located (680 m at point 6, Fig. 4b; Fig. S2) and thins to 108 m thick in the west. But, similarly to AS-2, it terminates at the Kohiku FZ.

AS-4 varies from 155 to 360 m thick (Table S3 and Fig. S2) and thins noticeably upslope (Fig. 4b; Fig. S2). Similar to AS-4, units AS-5, AS-6 and AS-7 thin above the AS-2 PG-drift location to less than 100 m thick (Table S3; Fig. S2) and then propagate eastward to the buried channel–levee system, where they thicken again. All three units have a maximum thickness of 350–400 m (Table S3; Fig. S2) and are parallel to each other, as well as to the units above and below.

AS-8 varies greatly in thickness throughout the basin (45–179 m thickness; Table S3) and is absent between line km 1750 and 2000 (Fig. 4; Figs. S1, S2). Following the trend of units AS-4 to AS-6, it also thins markedly above the AS-2 drift location (Figs. 4b, 5; Figs. S1, S2). AS-9 and AS-10 are similar in thickness (50–200 m thick; Table S3; Fig. S2) but varies in geometry. Both contain drifts, mostly concentrated in the western half of the basin (Fig. 4b; Fig. S1b). Units AS-4 to AS-10 are continuous along the entire transect, parallel and uniform in thickness up to the region above the anomalous basement topography and the AS-2 PG-drift.

The total sediment thickness ranges from 805 to 3909 m (Table S4; Fig. S2) with the thickest part being in the centre of the basin (Fig. 4; Fig. S1). Under the assumption that all units occur throughout the Amundsen Sea basin extent, an area of 1.2 to 1.4 × 10^6 km² (Fig. 1), and levels out above the block-faulted basement topography, we estimate the pre-glacial, transitional sequence (PG = 1281 m, T = 672 m, FG = 631 m) and the full glacial sediment volumes as: PG = 1.5 to 1.8 × 10^6 km³, T = min/max of the basin area, we estimate the pre-glacial, transitional sequence (PG = 1281 m, T = 672 m, FG = 631 m) and the full glacial sediment volumes as: PG = 1.5 to 1.8 × 10^6 km³, T = 0.6 to 1.0 × 10^6 km³ and FG = 0.8 to 0.9 × 10^6 km³ (Table S4; Fig. S2). Our presented volumes and age model is the highest for the FG sequence (1.5 × 10^6 km³) and is the lowest for the PG sequence (680 m, Table S3) and contains a large, ~3000 m thick basement from transect line km 500 to 2000 (unit AS-1) should be as old as 79.5–64 Ma.

The acoustic basement can generally be divided into four regions based on their distinct structural characteristics. (i) The transect segment from the eastern Ross Sea up to the Kohik FZ (line km 0–650 in Fig. 4; Fig. S1) shows a rugged and irregular topography, possibly associated with a decrease in spreading rates from 74 to 22 mm/yr, between chron 27 and 22 (63–50 Ma) (Wobbe et al., 2012; Fig. 2). Such low spreading rates are often described as the mechanism to generate rough basement surfaces (e.g. Malinierno, 1991; Small and Sandwell, 1989). (ii) From the Kohiku FZ to the Endeavour FZ (line km 650–1200 in Fig. 4; Fig. S1), the basement deepens and exhibits major faulting, possibly associated with a number of previous fracture zones in the late Cretaceous oceanic crust. (iii) From the Endeavour FZ to the Proto-Antipodes FZ (line km 1200–1650 in Fig. 4; Fig. S1), the basement topography is relatively smooth and no faults are visible. The top-of-basement has an east–west dip angle of about 1° (Fig. 5). A reason for this part of the basement being smoother than that in the west could be that the spreading rates are greater than 60–70 mm/yr for crust as old as 71 Ma (Wobbe et al., 2012). (iv) The basement east of the Proto-Antipodes FZ (line km 1650–2000 in Fig. 4; Fig. S1) exhibits a very rugged and faulted surface. Here, the transect crosses over to the elevated Marie Byrd Seamount (MBS) province with a greatly structured basement, interleaved with magmatic intrusions (Wobbe et al., 2012; Wobbe et al. 2012) also suggest that continental fragments from the breakup process may be embedded in this province.

5.2. Pre-glacial (PG) sequence

Using the sequence age ranges (Fig. 10) and sediment thicknesses (Table S4; Fig. S2) we determined sedimentation rates for the PG, T and FG sequences (Table S4). The age model contains large uncertainties, because borehole age-control is lacking and we therefore only outline the general trend for the first-order sequences. For the total sediment volume of 2.8 to 3.3 × 10^8 km³ in the Amundsen Sea basin, a min/max sedimentation rate of 1–3 cm/k.y. is estimated (Table S4). The PG sequence has the lowest min/max sedimentation rate estimate at 0.6–2.1 cm/k.y. The rate progressively increases to 1.0–2.5 cm/k.y. for the T sequence and is the highest for the FG sequence (1.5–3.6 cm/k.y.).

5. Discussion

5.1. Oceanic basement off the Amundsen Sea margin

Oceanic basement age is one of the better-constrained parameters for our seismic stratigraphic model of the Amundsen Sea continental rise. Basement age in the Amundsen Sea basin (Wobbe et al., 2012) increases becoming younger from 84 Ma (C34y) in the east to 42 Ma (C20n) in the northwest (Fig. 2). In our stratigraphic age model, the selected 12 points along the AS–RS transect range in basement age from 79.5 Ma at point 12 in the east, to 64.6 Ma at point 1 in the west (Figs. 4b, 10; Fig. S1). Therefore, sediments in direct contact with the
suggests a bottom-current flow from west to east (Fig. 12b; Stow et al., 2002; Shannugam, 2006; Shannugam and Camerlenghi, 2008). The occurrence of a larger and additional PG drift at ~55.8 Ma (Figs. 6c, 12a) supports our hypothesis that bottom-current circulation may have already been active in the Eocene and probably earlier, which is consistent with the findings of Uenzelmann-Neben and Gohl (2012). They observe drifts in Unit I (60–21 Ma) near the Marie Byrd Seamounts in the eastern Amundsen Sea, which also implies Paleogene bottom-current activity in this part of the basin. They infer that the likely cause and driving force were related to proto-Antarctic Bottom Water (AABW) deflected against a basement high. The latter is thought to have been caused by thermal uplift related to the emplacement of the Marie Byrd Seamount province. This may explain why it is the only drift in unit AS-3 and at this particular location, which indicates a shift of the PG depo-centre to the east (Fig. 12b).

5.3. Transitional (T) sequence: first WAIS advances to the shelf

The seismic stratigraphy established on the continental rise off Wilkes Land identified the first arrival of the EAlS to the margin at 33.5–30 Ma, marked by the regional unconformity WL-U3 (Fig. 11; Escutia et al., 2011). Comparing the type sections for the Amundsen Basin with the seismic line that crosses on IODP Leg 318, Site U1356 (De Santis et al., 1995, 2003; Escutia et al., 2005, 2011), it is visually evident that the lower sequences are similar in interval thickness, geometry and seismic character. We thus tentatively correlate our interpreted end of the pre-glacial sequence unconformity AS-u3/uPG-T to unconformity WL-U3 (Fig. 11). The WL-U3 age is in good agreement with our estimated AS-u3/uPG-T model horizon age of 34–28 Ma for the basin flanks (Figs. 10, 11).

From the seismic reflection shelf-slope-rise correlation between the TAN-2006-06 line and drill records of DSDP Leg 28 Sites 270–274, the pre-glacial sequence unconformity AS-u3/uPG-T is tentatively linked to unconformity RSU6 (Figs. 8 and 11). RSU6 has an age range of 34–29 Ma (Brancolini et al., 1995; De Santis et al., 1999; Brancolini and Leitchenkov, 2010; DSDP 270; Fig. 11). Our hypothetical model age of 34–30 Ma for the AS-u3/uPG-T horizon (Figs. 8, 10 and 11), broadly falls within the RSU6 age range given the variation in shelf and slope processes, the large uncertainties in our age model, and the absence of drilling information to constrain the age of AS-u3/uPG-T. We further speculate that the younger trend in the centre (~21 Ma; Fig. 11) may be associated with a delay in the arrival of the ice sheet along this part of the shelf, conceivably due to elevated topography such as a mountain range that it had to overcome before reaching the coast and shelf. The reconstructed 34 Ma maximum paleoaltotopography model of West Antarctica does suggest an elevated land area opposite the centre of the transect (Wilson et al., 2012), though the possible doming effect of Marie Byrd Land (e.g. LeMasurier, 2008) has not been taken into account. The minimum model however, included an elliptical dome with up to 1000 m of uplift. Alternatively, low-lying palaeo-river drainage systems from the PG climate regime may have redirected the ice streams and drainage towards basin flanks instead.

In the eastern Amundsen Sea, the pre-glacial sequence unconformity AS-u3/uPG-T is mapped on line AWI-20100117 of the AS–RS transect, matches the base of Unit II in crossing line AWI-95054 (Figs. 2, 9, 11). The base of Unit II was assigned a hypothetical age of 21 Ma in Uenzelmann-Neben and Gohl (2012) by a jump-correlation to the Ross Sea shelf horizon ages in De Santis et al. (1995). Our age-model however suggests an older age of at least early Oligocene, 30–28 Ma (Figs. 10 and 11). The Eocene–Oligocene transition age of 34 Ma for the uPG-T in our age model and in Leg 318, Site U1356 (Escutia et al., 2005, 2011) is favoured above the younger DSDP ages, because of its similar distal setting and thus assumed similar processes. Based on the seismic horizon mapping, the constructed age model (Figs. 10, 11) and borehole information, we therefore infer an age of about 34 Ma as the suggested timing of the first arrival of glaciers to parts of the coast, which is consistent with the reconstructed high paleoaltotopography of West Antarctica (Wilson et al., 2012) and the inferred extent of a modelled ice sheet at that time (Wilson et al., 2013).

By the late Oligocene to mid-Miocene, the deep-water Tasmanian and Drake Passage/Scotia Sea gateways had opened permanently, enabling a constant flow of the ACC (Fig. 12c, Table S1). Seismic records from various Antarctic margins indicate that from the mid-Miocene onward, ice sheets grounded more frequently on the continental shelf and pushed sediments across the slope to form progradational shelf sequences (e.g. De Santis et al., 1999; Brancolini and Leitchenkov, 2010; Gohl et al., 2013). This is in agreement with higher amplitudes and strong reflectivity seen in units AS-4, −5 and -6 above AS-u3/uPG-T (Figs. 4, 6; Fig. S1), which we associate with a larger sediment influx of poorly sorted material such as debris transported onto the rise through downslope processes, and an increase in the sedimentation rate to 1–2.5 cm/ky (Table S4).

In the Ross Sea shelf-slope-rise seismic horizon correlation (Fig. 8), the 26–18 Ma units AS-4 to AS-5 of our transitional regime, match well with units RSS-4a to RSS-3 of the Ross Sea (Brancolini et al., 1995; De Santis et al., 1999), and with WL-SS at the Wilkes Land margin (Fig. 11; Escutia et al., 2011). In the eastern Amundsen Sea seismic horizon correlation (Fig. 9), AS-4 to AS-6 collectively correlates to Unit II of Uenzelmann-Neben and Gohl (2012) but with a slightly older age range of 30–15 Ma for the eastern Amundsen Sea basin flank (Figs. 10 and 11). However, stratigraphic information from drill holes is absent in the Amundsen Sea and western Bellingshausen Sea and implies large uncertainties in both our and Uenzelmann-Neben and Gohl’s (2012) chrono-stratigraphic models.

Sediment drifts occur on the western flank of the basin on a basement high (Figs. 4, 6b; Fig. S1). The amalgamated base of the mounds/drifts lies on AS-u4, in AS-5, as well as on the basement, hence their ages should be at least younger than the youngest age of AS-4, which is 26–24 Ma (Figs. 4b, 6b, 10; Fig. S1). ODP Leg 178 Site 1095 reached sequences of this mid-Miocene age in an ice-distal drift deposit near the Antarctic Peninsula (e.g. Iwai et al., 2002; Uenzelmann-Neben, 2006). Elongate drifts and levee drifts were mapped in Unit II in the eastern Amundsen Sea in Unit and inferred to be 21–15 Myr old (Uenzelmann-Neben and Gohl, 2012; Fig. 11). This age difference of the same distance-correlated sequences could either be explained by large uncertainties in such a long-distance correlation, or by a change in the timing of the same sedimentation processes along the margin. The latter would imply similar WAIS-triggered sediment flux, but arrival on the shelf time-shifted from the Ross Sea to the Amundsen Sea embayment (Fig. 12c).

5.4. Full glacial (FG) sequence: intensified WAIS dynamics

The seismic stratigraphy work on the ice-distal Wilkes Land continental rise identified WL-US as the boundary horizon at which the EAlS grounded on the shelf (Escutia et al., 2011). The age of their horizon is consistent with our estimated AS-u6 horizon age of ~15 Ma for the Amundsen Sea basin flanks (Figs. 10, 11).

In the shelf-slope-rise seismic reflection correlation between the TAN-2006-06 line and DSDP Leg 28 Sites 270–272 of the Ross Sea, the transitional sequence unconformity AS-u6/uT-FG links to RSU4 (Figs. 8, 11). RSU4 has an age of 15–17 Ma, constrained by the DSDP Leg 28 boreholes (Brancolini et al., 1995; De Santis et al., 1999; Brancolini and Leitchenkov, 2010). This age range is again, similar to our model age of ~15 Ma on the basin flanks. The upper boundary of the transitional sequence (horizon uT-FG/AS-u6) follows a similar age trend than uPG-T. We infer that uT-FG/AS-u6 probably appears older on the basin flanks (15.3 Ma; Figs. 10, 11) and younger in the middle (9 Ma; Fig. 10) due to later bottom-water circulation and reworking of the Ross Sea gyre and/or AABW generation.

In the eastern Amundsen Sea, the transitional sequence boundary AS-u6/uT-FG in line AWI-20100117, matches the base of Unit III in
crossing line AWI-95054 (Figs. 2, 9, 11). The base of Unit III was assigned a hypothetical age of 14 Ma (Goh et al., 1997; Nitsche et al., 1997, 2000; Uenzelmann-Neben and Gohl, 2012), which is similar to our model age. Though these ages are almost identical, it is a yet unproven hypothesis that needs to be tested by drilling.

From the Miocene to the Pliocene–Pleistocene, the ACC developed fully, bottom water formation and circulation intensiﬁed, and the topography of West Antarctica evolved close to the present situation (Fig. 12c,d). After the mid-Miocene, the WAlS advanced across most of the continental shelves in times of glacial maxima (Fig. 12d). This period sees the highest sedimentation rate (1.5–3.6 cm/ky), probably due to the frequent advance/retreat cycles on the shelves. The seismic reﬂections show complex internal formation geometry, which is an indication for dynamic sediment inﬂux and intensiﬁed bottom-current activity (Figs. 4, 6a; Fig. S1).

Our model ages of –9 to 12 Ma, –5 Ma and –1 to 2 Ma for AS-u7, –8 and –9, respectively, are in good agreement with the DSDP Leg 28 drilling data and the seismic chrono-stratigraphy of the Ross Sea shelf (Figs. 8, 10, 11). In the eastern Amundsen Sea, seismic units AS-7 and AS-8 correlate with Unit III, whilst AS-9 and AS-10 correlate with Unit IV (Figs. 9–11), although our age model suggests an older age of 8 Ma as opposed to 4 Ma proposed by Uenzelmann-Neben and Gohl (2012). However, given the large uncertainties in age estimates in long-distance seismic correlation, and the absence of drilling information to support an age model, the differences are within reasonable range. The units are relatively thick along the eastern basin ﬂank (Fig. 4; Figs. S1, S2), which is consistent with depo-centres identiﬁed by Uenzelmann-Neben and Gohl (2012). This deposition geometry probably formed due to the close proximity to the glacial outlets of the Amundsen Sea Embayment (red arrows in Fig. 1). A series of stacked drifts lie on unconformity AS-u9. The stoss side of all the drifts trend west, suggesting an eastward ﬂowing current (Stow et al., 2002; Shannumagam, 2006; Shannumagam and Camerlenghi, 2008), which is in agreement with the post mid-Miocene (~15.5 Ma) bottom-current directions in Uenzelmann-Neben and Gohl (2012). The bases of the FG drifts in unit AS-8 are estimated to be –9 to 7 Ma, those in AS-9 in the ~8 to 4 Ma age range, and the ones in AS-10, younger than ~4 Ma (Figs. 4b, 10; Fig. S1). The drifts and stacked buried channel–levee systems in all the upper units (Fig. 4; Fig. S1), indicate that sedimentation transport and deposition processes of the Amundsen Sea continental rise occurred in an intense, dynamic and variable manner throughout the late Miocene, Pliocene and Quaternary. Most sediment was likely supplied by the major ice stream outlets of the Amundsen Sea Embayment to the southeast, and the Ross Sea region to the southwest.

5.5. Geodynamic implications for Marie Byrd Land

An interesting observation in the central Amundsen Sea rise is the down-to-west dip of the basin and the lower to middle sedimentary sequences, AS-1 to AS-3, deposited parallel to it (Fig. 5). Above AS-u3, the reﬂections become increasingly tapered, reaching their largest internal angles up to AS-u6, and then change to progressively more horizontal and parallel above AS-u6 (Fig. 5b). Subtle low-angle onlaps in the internal geometry of the second order sequences (e.g. AS-6 in Fig. 5b) lead us to speculate that this change in horizon angles may be attributed to tectonic uplift. Alternatively, we cannot rule out that subsidence may be one of the possibilities for the implied vertical motion. The negative anomaly in the basement residuals, modelled by Wobbe et al. (2014), and the deep Amundsen Sea basin geometry (Fig. 4; Fig. S1) collectively indicate some degree of local subsidence that may have been controlled by mantle processes. The sequences up to AS-u2 were either present before the uplift, or, the rate of uplift or subsidence was similar to the rate of deposition. The angle changes gradually towards AS-u3, indicating that uplift/subsidence may have initiated and accelerated up to AS-u6. Then, uplift or subsidence may have stopped and the units above AS-u6 deposited in a horizontal-parallel manner. Imposing our horizon age model on these observations, we imply that the start of vertical movement occurred before or near AS-u3/upG-T at about 30 Ma and continued until AS-u6/ut-FG at about 13 Ma, when the rate of a tectonic episode slowed down signiﬁcantly. The down-to-west dip could be related to the elevated Marie Byrd Sea-mount province to the east, which was emplaced by thermally driven magmatic intrusions and underplating of the crust as a product of partial melting (Kalberg and Gohl, 2014; Kipf et al., 2014). Kipf et al. (2014) however, derived an emplacement age of 65–56 Ma for the seamounts, which is about 20 myrs before our estimated initiation of the uplift. Instead, we speculate that increasing dynamic topography gradually tilted the area shown in Fig. 5, resulting in the change in horizon dip angles. Such tilting may perhaps be related to the Marie Byrd Land uplift caused by a mantle plume (e.g. LeMasurier, 2008). LeMasurier (2008) estimated the beginning of the MBL uplift to be at about 30–25 Ma, which broadly agrees with our postulated uplift age range. Uplift could have also caused simultaneous subsidence of the adjacent over-deepened western Amundsen Sea basin, in response to lithospheric bulging. However, without mantle-dynamic modelling, the extent and timing of the MBL uplift remain far from being constrained.

6. Conclusions

Our collective interpretation of a continuous Amundsen Sea to Ross Sea seismic transect and supporting seismic lines, enhances the understanding of the Cretaceous to present Amundsen Sea continental rise pre-glacial to full-glacial sedimentation processes. The major seismic units (AS-1 to AS-10) and unconformities (AS-u1 to AS-u9) were grouped into three ﬁrst order sequences, related to the pre-glacial (PG), transitional (T) and full glacial (FG) climate regimes of Antarctica.

The PG sequence comprises units AS-1 to AS-3, bounded by the basement and unconformity AS-u3/upG-T, with an age range of ~79 to 34 Ma. AS-3 is thicker than all the other units and contains a large, ~300 km long and up 400 m high, elongated levee drift on the eastern basin ﬂank. It is estimated to be of Eocene age and an eastward ﬂowing Paleocene to Eocene bottom current is inferred from its geometry. The PG sequence is 358–2203 m thick with an estimated sedimentation rate of 0.6–2.1 cm/ky.

The T sequence (units AS-4 to AS-6) has a collective age range of ~34 to 16 Ma and is present throughout the basin, from the Ross Sea to the eastern Amundsen Sea. Sediment drifts, mounds and sediment waves were identiﬁed and these are mostly conﬁned to the far western basin ﬂank, perched on a basement high. However, the geometry does not allow for a current direction interpretation on this scale. The T sequence ranges in thickness from 274 to 1070 m, thins from west to east and has an estimated sedimentation rate of 1.0–2.5 cm/ky, which is only slightly higher than that for the PG sequence. This transitional sequence is interpreted to indicate the Late Eocene onset of a glacially dominated environment of West Antarctica.

The FG sequence (units AS-7 to AS-10) ranges in age from ~16 to 0 Ma and is present along the whole transect length, throughout the Ross Sea and Amundsen Sea basins. Sediment drifts are typically 50–100 km wide and 200–300 m high. These are mainly concentrated on the western ﬂank of the basin and are stacked above the Pahemo FZ. The FG drifts’ stoss sides pinch out to the west, thus a strong eastward ﬂowing, mid-Miocene to Pliocene bottom-current can be inferred. AS-8 is the only FG unit that has a drift body on the eastern basin ﬂank of the Amundsen Sea basin. The FG sequence sediment thickness ranges from 374 to 889 m and has the highest sedimentation rate (~1.5–3.6 cm/ky) compared to the PG and T sequences.

During the pre-glacial to glacial periods, the major deposition centres remained in the basin centre and in front of Ross Sea outlet. No apparent shift in deposition, or change in deep-sea sediment transport processes between the Amundsen Sea and Ross Sea, is obvious.
The seismic sections show continuous sequences with consistent internal geometry and drifts in the centre, and on both basin flanks. The progressive change in horizon dip in the eastern part of the AS-RS transect, near the Pahemo Fracture Zone, may be attributed to vertical tectonics such as the Marie Byrd Land uplift, or magmatism during the emplacement of the Marie Byrd Seamounts.

The sediment rates and sediment rate changes, temperate versus cold-polar regime, age of onset of glaciation, onset and direction of bottom currents, as well as mapping the spatial distribution of the sequences, volume and depocentre shift extrapolations are currently constrained by good quality data, and the only available in this region. Our continuous seismic sections, the models, ages and rates can thus be viewed as an initial regional framework. Future model improvement can only be obtained by drilling and additional regional seismic data acquisition.

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