

# Resolving eddies by local mesh refinement

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## Abstract

Nesting in large-scale ocean modeling is used for local refinement to resolve eddy dynamics that would not be accessible otherwise. Unstructured meshes offer this functionality too by adjusting their resolution according to some goal function. However, by locally refining the mesh one does not necessarily achieve the goal resolution, because the eddy dynamics, in particular the ability of eddies to release the available potential energy, also depend on the dynamics on the upstream coarse mesh. It is shown through a suite of experiments with a zonally re-entrant channel that baroclinic turbulence can be out from equilibrium in wide (compared to a typical eddy size) zones downstream into the refined area. This effect depends on whether or not the coarse part is eddy resolving, being much stronger if it is not. Biharmonic viscosity scaled with the cube of grid spacing is generally sufficient to control the smoothness of solutions on the variable mesh. However, noise in the vertical velocity field may be present at locations where the mesh is varied if momentum advection is implemented in the vector invariant form. Smoothness of vertical velocity is recovered if the flux form of momentum advection is used, suggesting that the noise originates from a variant of the Hollingsworth instability.

*Key words:* Unstructured meshes, nesting, Hollingsworth instability

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## 1. Introduction

Nesting is a widely used tool in studies of large-scale ocean circulation, helping to resolve eddy dynamics over a limited area. The interest to nesting is motivated by several factors. For one thing, running a global fine-resolution model can still be prohibitively expensive if one's focus is on the regional dynamics. For another, the Rossby radius of deformation is rather small at high latitudes, so that resolving eddies there may require excessively fine resolution elsewhere if the resolution is uniform. There are numerous examples in the literature showing the success of the nesting approach (see, e. g., Chanut et al. (2008), Durgadoo et al. (2013), Mertens et al. (2014)), while the general principles of two-way nesting algorithms are reviewed by Debreu and Blayo (2008).

Unstructured meshes offer geometric flexibility and freedom with respect to mesh design, and may serve as an alternative to the nesting approach for structured meshes. In addition to applications where the unstructured meshes are used

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22 to better represent the domain geometry (see, e. g., Wekerle et al. (2013), Tim-  
23 mermann and Hellmer (2013)), the use of mesh refinement as a tool to resolve  
24 eddies is already a proven concept (see Ringler et al. (2013)). However, if the  
25 mesh resolution is allowed to vary, questions arise about the optimal way and  
26 consequences of varying it. Physical principles governing the selection of mesh  
27 resolution depend on applications, and not surprisingly, there is no unique solu-  
28 tion. The review by Greenberg et al. (2007) mentions some aspects, and each  
29 real application may add new details.

30 We focus below only on one aspect of the problem related to the use of locally  
31 refined meshes to resolve eddying regimes. The amplitude of eddy motions sim-  
32 ulated by a numerical model in a particular subdomain depends not only on the  
33 local resolution, but also on the presence of upstream perturbations, which serve  
34 as the seeds from which perturbations grow, and facilitate the release of available  
35 potential energy. While this remark may seem trivial, its implications can be very  
36 easily underestimated, and this study seeks to address them in a qualitative way.  
37 Although we deal with unstructured meshes, the results reported below can be of  
38 interest to a wider community of ocean modelers working with nesting tools on  
39 standard structured meshes.

40 We consider a baroclinically unstable eastward flow in a zonally-reentrant chan-  
41 nel, where the baroclinicity is maintained through forcing at its southern (warming)  
42 and northern (cooling) walls. A linear equation of state is used with the temper-  
43 ature being the only scalar field influencing the density. The flow is simulated on  
44 triangular meshes composed of nearly equilateral triangles. The resolution varies  
45 in the zonal direction, and by observing the flow variability along the channel the  
46 effect of the change in the mesh resolution is assessed.

47 A remark is due from the very beginning. Although the mesh refinement is dis-  
48 cussed, the dissipative operators are always varied accordingly, and the refinement  
49 means not only smaller scales but simultaneously smaller coefficients in explicit  
50 dissipative operators, and similar reduction in effective implicit dissipation asso-  
51 ciated with upwinding or flux limiting in transport equations. These two aspects  
52 (refinement and reduced dissipation coefficients) are inseparable, for dissipative  
53 operators are always designed to dispose of eddy variance of scalars and the eddy  
54 enstrophy on the grid scale. According to linear instability theory the wavelength  
55 of the most unstable wave (we take the Eady instability problem as an example)  
56 scales as  $\lambda \approx 3.9\pi L_R$  where  $L_R = NH/\pi f$  is the first internal Rossby radius,  $N$   
57 the buoyancy frequency,  $f$  the Coriolis parameter and  $H$  the fluid thickness. On  
58 meshes called eddy-permitting ( $1/3$ - $1/4^\circ$  at midlatitudes), eddies with the size of  
59  $\lambda/2$  can already be well represented, and yet it is well known that this resolu-  
60 tion is by far insufficient. The point is that the accompanying subgrid dissipation  
61 still turns out to be too high so that only a part of the extracted available po-  
62 tential energy (APE) is fluxed back to maintain kinetic energy at large scales,  
63 while the other part is lost to subgrid dissipation on small scales (see Jansen and  
64 Held (2014) for the spectral analysis of the APE release rate and energy transfers  
65 on eddy-permitting and resolving meshes). According to the results obtained in

66 Jansen and Held (2014) in simulations with a biharmonic Leith subgrid operator,  
 67 the APE release rate saturates at resolutions between 2 to 3 grid intervals per  $L_R$ ,  
 68 which as we shall see, also agrees with this study. Note also that this correlates  
 69 with the analysis of Hallberg (2013) for a related topic.

70 Our main goal below is to explore the response of turbulent flow to changes in  
 71 mesh resolution, concentrating on the retardation and overshoots in eddy variabil-  
 72 ity, and also on the ability to maintain smooth solutions in domains where reso-  
 73 lution varies. Since mesh refinement also implies reduced dissipation and higher  
 74 variability, a question on whether the dissipative operators can control the smooth-  
 75 ness of solutions in regions where the resolution is adjusted back from fine to coarse  
 76 one is tightly linked to the main goal.

## 77 2. Configuration and model

78 Most of the experiments are carried out in a zonally-reentrant channel  $L = 40^\circ$   
 79 long ( $0^\circ\text{E} - 40^\circ\text{E}$ ) occupying the latitude belt between  $30^\circ\text{N}$  and  $45^\circ\text{N}$ . The geom-  
 80 etry is spherical. There are 24 unevenly distributed layers going down to 1600 m.  
 81 Triangular surface meshes of variable resolution are used. The basic coarse resolu-  
 82 tion is  $1/3^\circ$ , and the basic fine resolution is  $1/12^\circ$ , giving the mesh refinement (or  
 83 stretching) factor, measured as the ratio of the largest to the smallest mesh edges,  
 84  $r = 4$ . Meshes are refined via relatively narrow transitional zones centered in most  
 85 cases at  $\phi_w = 7.5^\circ\text{E}$  and  $\phi_e = 32.5^\circ\text{E}$ , so that more than a half of the domain is well  
 86 resolved, and the other part is left coarse. The mesh resolution (edge length)  $h$   
 87 varies according to the hyperbolic tangent,

$$h = h_0(r + 0.5(r - 1)(-\tanh((\phi - \phi_w)/w_t) + \tanh((\phi - \phi_e)/w_t))) \quad (1)$$

88 where  $h_0$  is the side of the smallest triangle, and  $w_t$  (in degrees) defines the width  
 89 of the transitional zone. There are some variations of this basic setup. The pa-  
 90 rameters of the meshes used in different runs are presented in Table 1.

91 The density depends linearly on the temperature,  $\rho - \rho_r = -\rho_r \alpha (T - T_r)$ , with  $\rho_r$   
 92 and  $T_r$  the constant reference values and  $\alpha = 2.5 \times 10^{-4} \text{ K}^{-1}$  the thermal expansion  
 93 coefficient. The initial temperature distribution is linear in the meridional direction  
 94 with the gradient  $T_{0y} = -0.5 \times 10^{-5} \text{ K/m}$  and also in the vertical direction with  
 95 the gradient  $T_{0z} = 8 \times 10^{-3} \text{ K/m}$  in the entire channel. There are buffer zones  
 96  $1.5^\circ$  wide adjacent to the northern and southern walls where the temperature  
 97 is relaxed to the initial one over the entire depth. The inverse relaxation time  
 98 scale varies linearly from  $(3 \text{ day})^{-1}$  at the wall to zero outside the  $1.5^\circ$  zones.  
 99 A small sinusoidal perturbation is applied to the temperature to speed up the  
 100 development of the baroclinic instability, which equilibrates in about half a year.  
 101 We only deal with short runs of several years (4 or 5) in duration and present  
 102 the results averaged over the entire period of integration excluding the first year.  
 103 While this is certainly insufficient to obtain stationary patterns of eddy variances,  
 104 it is sufficient to draw qualitative conclusions for our questions. The configuration  
 105 is schematically presented in the top panel of Fig. 1.

run	$rh_0$	$h_0$	$w_t$	$\phi_w$	$L$
A	1/3	1/12	1	7.5	40
A'	1/3	1/12	2.5	7.5	40
B	1/3	1/18	1.5	7.5	40
C	1/3	1/12	1.5	7.5	60
C'	1/3	1/12	4.5	10	60
D	1/6	1/24	1.5	7.5	40
E	1/9	1/36	1.5	10	40

Table 1: Geometrical parameters of meshes used, see Eq. (1).  $\phi_e$  is always symmetric to  $\phi_w$  with respect to the center of the mesh. The second and third column specify the coarse and fine resolution. All quantities are in degrees.

106       Runs A and A' use the mesh refinement factor  $r = 4$ , and differ in the width of  
107 transitional zone  $w_t = 1^\circ$  (A) and  $w_t = 2.5^\circ$  (A'). Run B is performed on a mesh  
108 with  $w_t = 1.5^\circ$ , and a resolution of  $1/18^\circ$  in the fine resolution domain ( $r = 6$ ).  
109 Run C is similar to A, but the channel is longer ( $60^\circ$ ), with the same length of  
110 the coarse resolution domain, but an extended fine resolution domain. In C' the  
111 transitional part is rather wide, and for that reason  $\phi_w$  ( $\phi_e$ ) is moved a bit to the  
112 east (west). Case D doubles the resolution of the mesh of case A making it eddy  
113 resolving everywhere (see below), while case E improves the resolution further. In  
114 this case  $\phi_w$  ( $\phi_e$ ) is also slightly shifted in order to make the length of the coarse  
115 part equal to the fine one (for both are eddy resolving).

116       The simulations are performed with a finite-volume ocean circulation model  
117 described in Danilov (2012). It uses a cell-vertex (quasi-B-grid) discretization.  
118 The runs are stabilized with a weak quadratic bottom drag (with  $C_d=0.001$ ) and  
119 biharmonic viscosity. The scalar advection is simulated with a variant of a gradi-  
120 ent reconstruction scheme which combines 3rd and 4th order estimates (weighted  
121 as 0.15/0.85), with the 3rd order part responsible for some upwind diffusion. On  
122 uniform meshes it is equivalent to a flow-oriented biharmonic operator. No explicit  
123 horizontal diffusion is used, and the Pacanowski–Philander scheme (Pacanowski  
124 and Philander (1981)) is applied for vertical mixing. The biharmonic viscosity  
125 coefficient includes contributions from the Smagorinsky, Leith and modified Leith  
126 parameterizations (see Fox-Kemper and Menemenlis (2008) for a review), multi-  
127 plied with the areas of mesh cells (to 'translate' them from the original harmonic  
128 to the biharmonic form). It is capped at  $A_{bh} = v_v S_c^{3/2}$ , where  $v_v = 0.02$  m/s and  
129  $S_c$  is the cell area. Additionally, because of the too large velocity space of the  
130 cell-vertex discretization, we apply a background 'biharmonic' filter as detailed in  
131 Danilov and Androsov (2015). It provides an efficient coupling of velocities at the  
132 nearest cells. It is equivalent to the biharmonic viscosity operator with the coef-  
133 ficient  $v_f h^3$ , where  $v_f = 0.007$  m/s, on a uniform equilateral mesh, but deviates  
134 from it on general meshes. No 'manual' tuning of dissipation is performed. The  
135 vector-invariant form of momentum advection is used in the runs listed in Table 1.  
136 It turned out that it may lead to a transient noisy pattern in the vertical velocity  
137 over the varying portion of mesh (see the next section), but it does not affect the

138 main result here.

139 By construction, the vertical shear is  $\Lambda = g\alpha T_{0y}/f$ , which introduces the inverse  
140 time scale and, multiplied with the fluid depth, the scale for the horizontal velocity  
141  $U = \Lambda H$ . There are two more inverse time scales set by the Coriolis frequency  
142  $f$  and the buoyancy frequency  $N^2 = g\alpha T_{0z}$ , which together with  $\Lambda$  would lead to  
143 two dimensionless parameters related to the evolution of baroclinic instability.

144 The bottom drag ( $C_d$ ) affects the vertical profile, removing the symmetry be-  
145 tween the surface and the bottom, and thus influences the propagation speed of  
146 unstable baroclinic perturbations. The ratio between the largest (south) to small-  
147 est (north) Rossby radii is  $\sqrt{2}$ , so that coarse and fine meshes are either eddy  
148 permitting or eddy resolving for all latitudes, but there is no symmetry between  
149 the north and south. The strength of variability depends on the Reynolds numbers  
150 based on the scale of eddies and grid scale and on respective Peclet numbers. Vari-  
151 ations in these parameters will lead to quantitative changes, but are not expected  
152 to change our conclusions on a qualitative level.

### 153 3. Results

#### 154 3.1. Retarded turbulence development

155 The middle and bottom panels of Fig. 1 show a snapshot of temperature and  
156 relative vorticity at a depth of 100 m in case A, which is typical of the flow at other  
157 times. The two lines in the bottom panel are drawn at  $\phi = \phi_w$  and  $\phi = \phi_e$ . The  
158 temperature distribution is indicative of the presence of a strong eastward flow  
159 (with zonal velocities in excess of 1 m/s). The relative vorticity pattern illustrates  
160 the marked difference in eddy dynamics on the coarse and fine parts of the mesh  
161 and between the upstream and downstream parts of the refined domain. Indeed,  
162 eddies do not appear immediately as the mesh is refined, but develop downstream  
163 of the western edge of the fine-resolution section. The estimate  $1/\tau \sim 0.3\Lambda f/N$   
164 for the maximum growth rate of linear Eady instability problem gives for the  
165 turbulence development length  $L \sim U\tau \sim 3\pi L_R$ , which is approximately just the  
166 scale of the fastest growing waves (in this estimate  $U$  should be the amplitude  
167 of velocity at steering level, which is about half of the surface velocity, but we  
168 neglect this difference). For the linear stratification used by us,  $N = 4.5 \times 10^{-3} \text{ s}^{-1}$ ,  
169 resulting in  $L_R \approx 26 \text{ km}$  at the channel axis, and  $L$  about  $3^\circ$ . The cyclones forming  
170 around the longitude of  $10^\circ$  have the size of  $L/2$ , in agreement with this scaling. A  
171 much longer distance is needed for turbulence to equilibrate downstream, through  
172 the formation of new eddies and their straining into elongated vorticity filaments.

173 In order to see the turbulence 'retardation' effect, we present the pattern of the  
174 standard deviation (std) of the sea surface height (ssh) in Fig. 2. While longer  
175 time averaging is needed to make the pattern more uniform, we can nevertheless  
176 conclude that the turbulence is suppressed for about  $8\text{-}10^\circ$  into the refined domain,  
177 but overshoots past the downstream edge of the refined area. Here we measure  
178 the 'retardation' length as the distance where the std is still less than the median  
179 value between the coarse-mesh and fine-mesh values. The extent of suppression

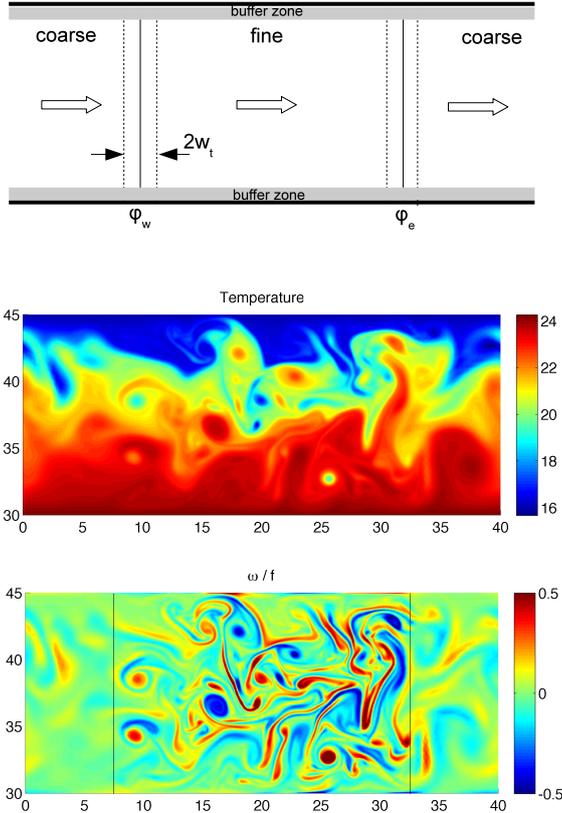


Figure 1: Top: Setup schematics. Large arrows indicate the flow direction. The solid meridional lines show the centers of transitional zones, the dashed lines mark the transitional zones, as described by Eq. (1). Middle and bottom: Snapshots of temperature ( $^{\circ}\text{C}$ ) and relative vorticity (normalized by the local value of the Coriolis parameter) at approximately 100 m depth in case A. While only the sharpness of temperature filaments reveals the presence of mesh refinement in the middle panel, the relative vorticity field shows the formation of eddies on the fine mesh and their decay on the coarse mesh.

180 or overshoot depends on the quantity being explored. In a quasi-geostrophic scal-  
 181 ing the spatial spectrum of elevation variance will be dominated by larger scales  
 182 compared to the spectra of horizontal velocity or relative vorticity. This implies  
 183 that the difference between the eddy-permitting (coarse) and eddy resolving (fine)  
 184 parts of the mesh is less expressed in the ssh variability, and it is only a factor of  
 185 about 2 in Fig. 2.

186 The variability of other fields, likewise, confirms the presence of 'retardation'.  
 187 We use the meridionally averaged variance of three-dimensional fields to further  
 188 demonstrate it. Figure 3 shows, from top to bottom, the mesh resolution  $h/h_0$  as  
 189 given by Eq. (1) (1 corresponds to  $1/12^{\circ}$ ), std for temperature, relative vorticity  
 190 and vertical velocity, and the pattern of the eddy kinetic energy. All patterns  
 191 of variability convey the same message and show, similar to Fig. 2 above, that  
 192 the turbulence saturation is delayed some distance downstream into the refined

193 domain. Note that the colorbar does not drop to zero for the temperature vari-  
 194 ability, and consistent with the behavior of the ssh, the std of temperature varies  
 195 only within a factor of 2. Clearly, this ratio would be larger if the coarse mesh  
 196 were coarser. The changes seen in the vertical velocity and relative vorticity are  
 197 more dramatic. The new aspect of patterns in Fig. 3 (compared to Fig. 2) is  
 198 the modulation of variability with depth showing that the turbulent flow did not  
 199 reach a fully saturated level even at the end of the refined zone. The variability  
 200 gradually propagates into deeper layer, a process continuing downstream to longi-  
 201 tude of about  $30^\circ$ , although this adjustment is not as strong as the change at the  
 leading part of the refined domain.

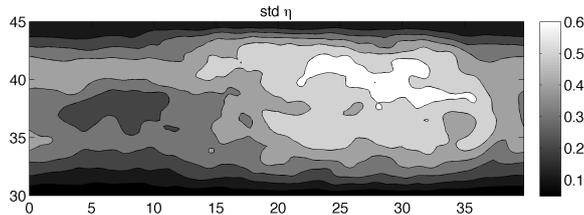


Figure 2: Standard deviation of sea surface height (m) in case A (contours are drawn in 0.1 m intervals). The centers of transitional zones are at  $7.5^\circ$  and  $32.5^\circ$ . The 'quasi-equilibrium' behavior is only reached  $8-10^\circ$  downstream into the refined domain, beginning from the longitude of  $18-20^\circ$ . In contrast, the variability is stronger than would be maintained on the coarse mesh for about  $5^\circ$  downstream the fine-coarse transition.

202  
 203 As the mesh becomes coarser (past  $32.5^\circ$ ), the turbulence decays. Among the  
 204 fields shown in Fig. 3, the temperature variability survives the furthest. The  
 205 relative vorticity variability drops down almost within the mesh transition zone.  
 206 This is linked to the fact that the relative vorticity variance is contributed by the  
 207 smaller scales of the flow compared to the temperature or velocity.

208 The amount of the available potential energy released by eddies depends on  
 209 their strength, so that the patterns presented above should correlate with the  
 210 pattern of the conversion rate of the available potential energy to the kinetic energy.  
 211 In Fig. 4 we present the time and meridional mean of the distribution of the  
 212 conversion rate  $R = -g\rho w$ , where  $\rho$  is the density perturbation and  $w$  the vertical  
 213 velocity, in run A<sup>1</sup>. The temperature relaxation zones adjacent to the walls are  
 214 excluded from averaging, so the quantity shown is mostly contributed by eddy  
 215 perturbations. The distribution of  $R$  as a function of horizontal coordinate remains  
 216 patchy for the available duration of experiments, but reveals a consistent pattern  
 217 after meridional averaging. The negative contributions in Fig. 4 originate from  
 218 the vicinity of the southern wall. The modulation seen in the pattern is linked  
 219 to the lack of zonal symmetry in the variability, as suggested by the asymmetry  
 220 between the northern and southern part of the channel in Fig. 2 (the mean flow  
 221 also contains a non-zonal pattern). Despite the modulation, the coarse part of the

<sup>1</sup>Note that the released energy is redistributed, so it does not coincide with the pattern of pressure work.

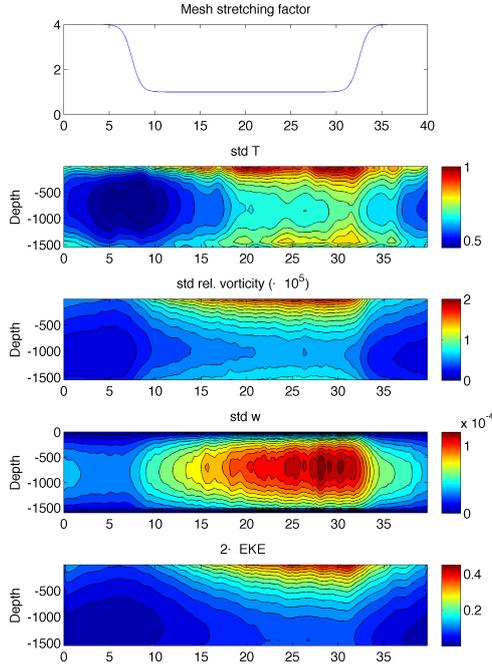


Figure 3: Mesh stretching factor (top) and the meridional mean patterns of standard deviation for temperature ( $^{\circ}\text{K}$ ), relative vorticity ( $\text{s}^{-1}$ ), vertical velocity ( $\text{m/s}$ ) and eddy kinetic energy ( $\text{m}^2/\text{s}^2$ ) in case A.

222 channel and the initial part of the fine mesh are characterized by lower values of  
 223  $R$ , and the areas of increased  $R$  gradually go deeper until  $15^{\circ}$  E.

### 224 3.2. Smooth versus sharp mesh transitions and further mesh refinement

225 Case A' and case B show very similar behavior to case A and are not displayed.  
 226 For A' the underlying reason is rather simple. The resolution of  $1/3^{\circ}$  is still coarse  
 227 ( $h \sim L_R$ ) and supports only weak transient motions for the selected viscosity in the  
 228 flow entering the fine-resolution part. So as far as the geometrical transition zone  
 229 remains narrower than the physical transition zone needed for the turbulent flow  
 230 to saturate, its width is of little relevance (but see further). Case B is characterized  
 231 by the finer resolution and hence smaller dissipation. One might expect that the  
 232 turbulent flow will evolve faster into a saturated regime, which is, however, not  
 233 observed. This signals that the subgrid dissipation on the  $1/12^{\circ}$  mesh of case A is  
 234 already sufficiently small, so that further refinement and decrease in viscosity and  
 235 implicit diffusivity only leads to the formation of smaller scales leaving the larger-  
 236 scale part of the spectrum unmodified. Understanding all the detail requires a  
 237 separate study, which is not pursued here. Although we have not performed runs  
 238 with even larger refinement factors, we would expect that the same 'retarded'  
 239 behavior will be observed even then.

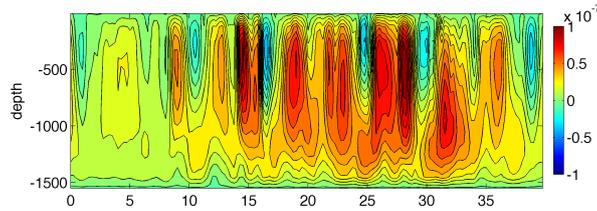


Figure 4: The meridionally averaged distribution of the APE conversion rate (in  $\text{W}/\text{m}^3$ ) in case A. The narrow forcing bands in the vicinity of the northern and southern walls are excluded. The black areas correspond to the negative conversion rate, and it is positive outside them. The alternating character of the distribution is linked to the presence of non-zonal component in the mean flow caused by the change in mesh resolution.

240 Since the length of the fine part in case A seems to be insufficient for reaching  
 241 full equilibration, simulations have been repeated on a mesh with the fine part  
 242 approximately twice as long (C and C'). The temperature variability in case C, as  
 243 shown in Fig. 5, is close to equilibrium east of  $33^\circ$ , indicating that the channel of  
 244 cases A and B is indeed too short. For the longer channel, according to the pattern  
 245 of temperature variability, the initial evolution of turbulence (between  $10^\circ$ – $17^\circ$ ) is  
 246 followed by the region between the longitudes of  $17^\circ$  and  $33^\circ$  where the strength  
 247 of turbulence is still under adjustment, although at a slower rate. It is close to  
 248 equilibrium on the remaining part of the fine mesh. Similar behavior is seen for  
 249 the vertical velocity, relative vorticity and kinetic energy, yet it is also clear that  
 250 full equilibrium is not reached even in the long channel. The variability of relative  
 251 vorticity and the eddy kinetic energy continue to propagate to deeper layers all  
 252 the way to  $\phi_e$ . In order to characterize this behavior, in Fig. 6 we present the std  
 253 of relative vorticity (thick black curve) averaged both meridionally and vertically.  
 254 The thin black curve represents an exponential fit,  $F(x) = a + b \exp(-(x - \phi_w)/L_s)$ ,  
 255 with the  $e$ -folding length  $L_s = 10^\circ$ , and parameters  $a$  and  $b$  set by the std values at  
 256  $x = \phi_w$  and the end of the fine mesh section. We will refer to  $L_s$  further as the  
 257 saturation length. Fitting the variability of other fields suggests  $L_s$  between 10  
 258 and  $13^\circ$ .

259 Working with the longer channel gives us the possibility to explore the effect  
 260 of very gradual transition in mesh resolution. The right panel in Fig. 5 shows  
 261 the statistics for case C' in which the transitional zones are approximately of the  
 262 length of the coarse part of the channel, and are also comparable to the distance  
 263 it takes to reach saturation in cases A and C. Although there are some differences,  
 264 the central, equilibrated parts between  $25^\circ$  and  $45^\circ$ , where the resolution is fine on  
 265 both meshes, are rather similar. For the 'coarse-fine' transition an offset to the east  
 266 is observed in patterns of relative vorticity and kinetic energy in case C', which is  
 267 explained by the larger  $\phi_w$  (see Table 1). The decay becomes more gradual on the  
 268 'fine-coarse' transition in case C'. We therefore conclude that smooth transition  
 269 does not hinder reaching the equilibrium even for transition zones comparable in  
 270 size to the physical length needed to reach saturation in cases A and C. To facilitate  
 271 the comparison, Fig. 6 presents also the meridionally and vertically mean std of

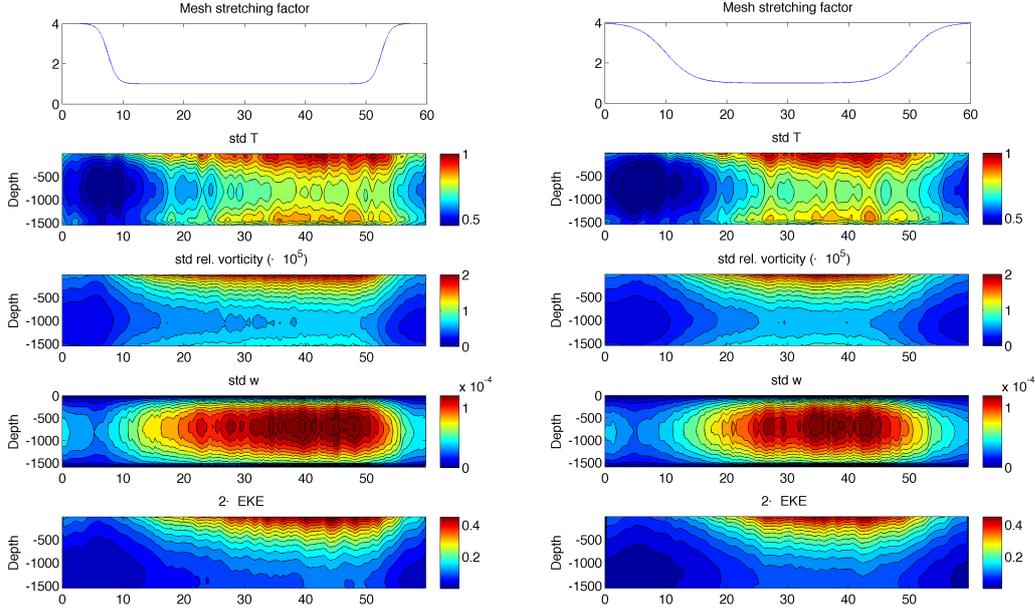


Figure 5: Left: Same as in Fig. 3, but in a longer channel (case C). Right: Case  $C'$  where the transition between coarse and fine mesh is made more gradual.

272 relative vorticity in run  $C'$  (thick gray curve), with the exponential fit based on  $L_s$   
 273 found for case C. There is a clear offset of  $2.5^\circ$  between the black and gray curves  
 274 on the west of the refined mesh, but both approach their quasi-equilibrium further  
 275 downstream approximately by the same exponential law despite the resolution is  
 276 coarser in case  $C'$  over a distance of about  $L_s$ .

277 In practice one would like to reach a goal resolution in a predefined domain,  
 278 and the question is how this resolution has to be matched to the coarse one outside.  
 279 The comparison presented in Fig. 5 and 6 can be viewed from this perspective.  
 280 Let us arbitrarily define the boundary of the refined domain to be where  $r = 1.1$ ,  
 281 which is at approximately  $9.2^\circ$  E for case C and  $17.6^\circ$  E for case  $C'$  on the west  
 282 side (indicated by arrows in Fig. 6). Considering the longitude of  $25^\circ$  as the  
 283 place where the turbulence becomes saturated in both cases, we see that in case  
 284  $C'$  one would sacrifice less of the fine-resolution domain than in case C. Viewed  
 285 from this standpoint, smooth transitions should be preferred, and the size of the  
 286 transitional zone should be comparable to the length needed for turbulence to  
 287 saturate. In designing a mesh, the transition zone should start sufficiently far  
 288 outside the region of interest. This length may also depend on the resolution of  
 289 the coarse part of the mesh.

290 One does not expect large changes to the behavior presented above if the coarse  
 291 mesh is made even coarser except for even further suppressed variability over the  
 292 coarse part of the domain and perhaps somewhat longer equilibration zone. Indeed,  
 293 the transient features that serve as seeds for the baroclinic instability over the fine

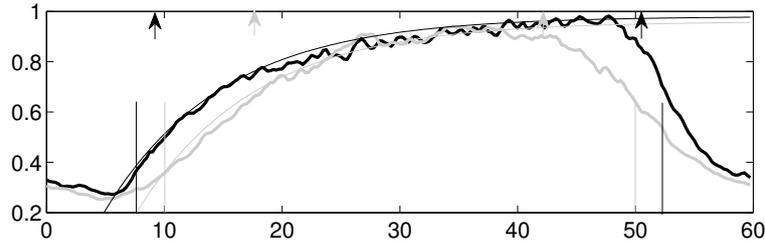


Figure 6: The std of relative vorticity averaged meridionally and vertically, in  $10^{-5} \text{ s}^{-1}$ , in cases C (thick black curve) and C' (thick gray curve). Thin lines correspond to exponential fit with the e-folding (saturation) scale  $L_s = 10^\circ$ . The vertical thin lines indicate positions of the centers of the mesh transition zones. The arrows show the boundaries of fine mesh where  $r = 1.1$ .

294 mesh are already not strong enough on the eddy-permitting coarse mesh in cases  
 295 A-C, so that further coarsening would not change the overall picture. In contrast,  
 296 refining the coarse mesh so that it becomes eddy resolving may have an impact on  
 297 the turbulence 'retardation', as indicated by the results of cases D and E shown  
 298 in Fig. 7. Note that the transition zones are centered at  $10^\circ$  and  $30^\circ$  in case E so  
 299 that the fine part occupies exactly a half of the channel.

300 In case D the mesh is twice as fine as in case A, so that the Rossby radius  $L_R$   
 301 is approximately resolved by two triangles on the coarse mesh. While some delay  
 302 in reaching saturation downstream the 'coarse-fine' transition zone is still present,  
 303 the temperature and relative vorticity patterns now change much more sharply  
 304 (much smaller distance is needed to reach saturation) than in case A. There is  
 305 much more uniformity in the patterns of vertical velocity and EKE.

306 In case E the mesh is further refined, and now the coarse mesh resolution  
 307 approximately corresponds to three elements per the Rossby radius. At this reso-  
 308 lution, there is little difference in the variability of temperature and eddy kinetic  
 309 energy between the coarse and fine parts of the mesh, but there are still differences  
 310 in the relative vorticity and vertical velocity fields. These fields are contributed by  
 311 small scales of the flow, so they show less saturation than the variability of temper-  
 312 ature and velocity as the mesh is refined. The 'coarse-fine' transition is now sharp  
 313 for all fields, with no apparent 'retardation' (there is still some delay in case D).  
 314 We conclude that the resolution of about two mesh elements per Rossby radius is  
 315 critical for representing eddies, similar to the conclusion in Hallberg (2013). At  
 316 finer resolution the large-scale part of the flow is already faithfully modeled, and  
 317 we may guess that the APE to KE conversion is close to saturation everywhere in  
 318 the domain.

### 319 3.3. Vertical velocity in transitional zones

320 Thus far we have dealt with the retardation of the turbulence development  
 321 related to the lack of sufficiently strong perturbations in the flow upstream of the  
 322 fine-resolution area. We concluded that smooth transition should be preferred. We  
 323 discuss now some numerical aspects related to the variable resolution as applied

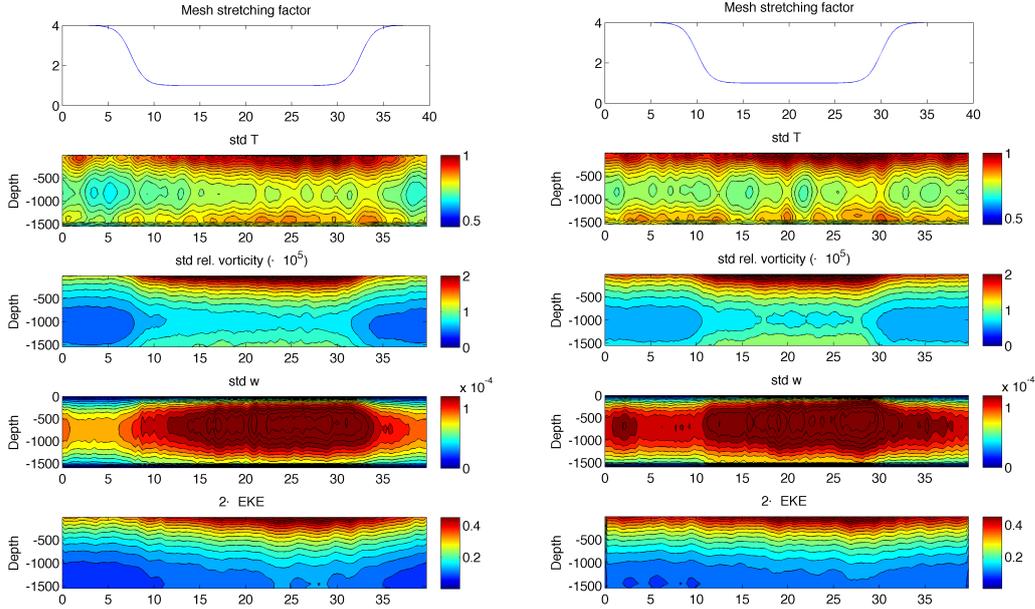


Figure 7: Same as in Fig. 3, but on a finer mesh, for case D (left) and case E (right). In both cases the ‘coarse’ part is eddy-resolving. In case E the fine part is slightly shorter than in case D, see Table 1, so that  $7.5^\circ$  E in the left panel should be related to  $10^\circ$  E in the right panel for the coarse-fine transition.

324 to simulating eddy dynamics.

325 Any unstructured mesh with variable resolution contains geometrical irregu-  
 326 larities in places where its resolution is changing, and local errors in representing  
 327 numerical operators will more pronounced. Controlling these errors requires a cer-  
 328 tain level of mesh smoothness and the availability of dissipative operators that can  
 329 handle the irregularities on the grid scale. For eddying flows, viscous dissipation  
 330 is generally tuned so as to eliminate the cascade of variance at grid scales. On  
 331 variable-resolution meshes this has an additional implication, for local dissipation  
 332 has to eliminate the variance also in places where eddies are advected from the  
 333 fine to coarse mesh, supplying variance at scales and with levels that do not match  
 334 those of local dynamics. In practice this means that numerical stability and regu-  
 335 lar behavior of modeled fields need to be maintained, which of course depends on  
 336 the discretization and simulated dynamics.

337 In all simulations reported here the dissipation (viscosity) is selected so as  
 338 to maintain the grid-scale Reynolds number at a certain level when the mesh  
 339 resolution is varied ( $h^3$  scaling for biharmonic viscosities), and this turns out to  
 340 be sufficient for preserving smooth behavior of temperature, velocity and relative  
 341 vorticity, as illustrated by the snapshots in Fig. 1, and similar patterns in other  
 342 simulations. With the exception of cases  $A'$  and  $C'$ , the mesh transition occupies a  
 343 zone which is just the size of a typical eddy, so that sharper transitions are hardly of

344 practical interest. The ability of biharmonic operators to control the flow on such  
345 meshes is a very encouraging and important message. However, as is common in  
346 such situations, the full story is incomplete without analyzing the behavior of the  
347 vertical velocity ( $w$ ). Inspection of  $w$  in case A (not shown) already reveals a noisy  
348 pattern at the fine-coarse transition (the coarse-fine transition remains virtually  
349 noise-free). Its presence indicates that either the available dissipation still fails to  
350 control all the details of solutions, or that some specific numerical issues come into  
351 play. The noise becomes stronger if the mesh resolution is refined. The upper panel  
352 of Fig. 8 presents a snapshot of vertical velocity from case E at approximately  
353 400 m depth, where the problem of noise is much more apparent than in case A.  
354 In case E noise is seen in both transition zones, but is also present on the coarse  
355 mesh (around  $35^\circ$  E). We were not able to suppress the noise by increasing the  
356 viscosity threshold three-fold in the zones where the resolution is varied. However,  
357 we did not see any immediate effect of this noise on the model stability. There  
358 is no apparent increase in  $w$ -variance in places where it is present (it looks rather  
359 like scattering on mesh inhomogeneities).

360 The fact that noise is seen only in  $w$ , and not in the relative vorticity hints at  
361 the Hollingsworth instability (Hollingsworth et al. (1983)) which sometimes ham-  
362 pers the performance of codes based on the vector-invariant form of momentum  
363 advection and was explored for C-grid discretization. It should be recalled that  
364 the instability occurs because the two terms on the right hand side of the equality  
365  $(\mathbf{u} \cdot \nabla) \mathbf{u} = \omega \mathbf{e}_z \times \mathbf{u} + \nabla \mathbf{u}^2 / 2$ , where  $\mathbf{u}$  is the horizontal velocity and  $\omega = (\nabla \times \mathbf{u}) \cdot \mathbf{e}_z$   
366 the relative vorticity, do not necessarily give the left hand side in the discrete for-  
367 mulation, but may contain an error that projects on the horizontal divergence and  
368 leads to instability. The problem does not occur for the relative vorticity because  
369 discretizations commonly maintain the property that the discrete curl operator  
370 gives exactly zero when applied to the discrete gradient. An analysis similar to  
371 the simplified analysis of Hollingsworth et al. (1983), however, shows that the  
372 cell-vertex (quasi-B-grid) discretization of the vector-invariant form of momentum  
373 employed by us is stable on uniform meshes, so that the problem can only be asso-  
374 ciated with some non-compensation on a variable mesh. While a rigorous analysis  
375 is beyond the scope of this paper (the eigenvalue analysis of Hollingsworth et al.  
376 (1983) can in principle be repeated on a limited patch of a non-uniform mesh),  
377 there are additional arguments in favor of this viewpoint.

378 In the upper panel of Fig. 8 there is no noise over the fine part, where the mesh  
379 is really uniform. Due to the mesh generation procedure, the coarse part is only  
380 quasi-uniform. This observation confirms that the noise is associated with mesh  
381 irregularity. Further, and more conclusive evidence is provided by the bottom  
382 panel of Fig. 8, showing a snapshot from simulations configured as case E except  
383 for the flux form of momentum advection. The noise is absent everywhere. The  
384 flux implementation of momentum advection is described by Danilov (2012). We  
385 first compute the flux divergence on scalar control volumes, and then average the  
386 result to vector points (cell centroids). Since the continuity is also formulated on  
387 scalar control volumes, this flux form is consistent with it.

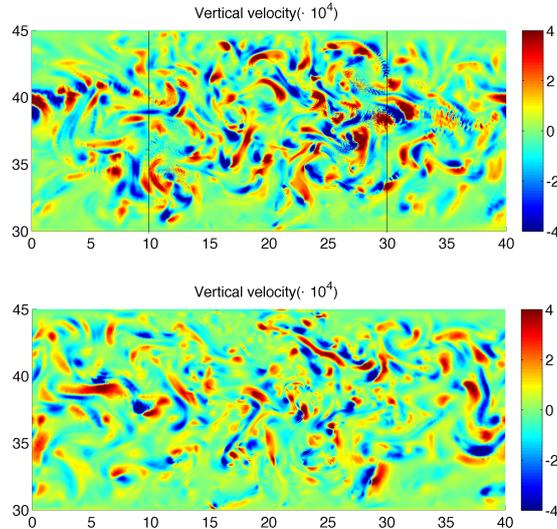


Figure 8: Snapshots of vertical velocity in case E at 400 m depth in simulations using the vector-invariant form (top) and the flux form (bottom) of momentum advection. Note the absence of noise in the second case and its presence in the first. The narrow transition zones are centered at the longitudes of  $10^\circ$  and  $30^\circ$  (indicated by black lines).

388       The flux form of momentum advection, eliminating the noise in  $w$ , also slightly  
 389       modifies statistics as compared to the vector-invariant form, leading to lower values  
 390       of variability for vertical velocity and relative vorticity. Since it does not change  
 391       the response to the varied resolution we keep the simulations carried out with the  
 392       vector-invariant form. The flux form of momentum advection can be discretized  
 393       in many ways based on high-order transport schemes, but it remains to see how  
 394       they behave in situations studied here.

#### 395   4. Discussion

396       The question of how to refine the mesh resolution is a difficult one, and here  
 397       we qualitatively explored only one of its aspects. The effect of retarded turbu-  
 398       lence development has an implication that the area of refinement has to be suffi-  
 399       ciently large or be connected through sufficiently wide transitional zones in order  
 400       to "achieve" the goal resolution if the transition is made from a non-eddy-resolving  
 401       mesh. Clearly, the examples considered above involve a particular flow and may  
 402       to a degree overemphasize the effect. The dynamics are dominated by the zonal  
 403       flow which advects transient features downstream leading to overshoots in variabil-  
 404       ity as the flow passes from the fine to the coarse mesh, and retarded turbulence  
 405       development when the flow passes from the coarse mesh to the fine mesh. The  
 406       effect might be weaker for recirculation zones or marginal seas where the role of  
 407       upstream 'seed' in triggering the development of turbulence will be less significant.  
 408       It may also be reduced in the presence of small-scale topography triggering eddy

409 formation. It remains to see whether this is so.

410 According to Hollingsworth et al. (1983) the noise in vertical velocity associ-  
411 ated with the Hollingsworth instability depends on the detail of the implementation  
412 of the vector-invariant form of momentum advection. In our case, as mentioned  
413 above, the instability is not obtained on uniform meshes, and it is of interest to  
414 learn how it is linked to the mesh non-uniformity. It remains to see whether the  
415 vector-invariant form of momentum advection can be adjusted so as to eliminate  
416 the noise in the vertical velocity on highly-variable meshes, or whether it should be  
417 abandoned altogether on such meshes for the cell-vertex discretization. Although  
418 the behavior seen here is not necessarily characteristic of other unstructured-mesh  
419 discretizations, we think that the observation that the numerical implementation  
420 of momentum advection is of primary importance for maintaining smooth behavior  
421 of simulated fields on meshes with variable resolution is of general importance and  
422 deserves special attention on its own. We mention in this respect the analysis by  
423 Gassmann (2013) carried out for the quasi-hexagonal C-grid discretization which  
424 illustrates implications of the Hollingsworth instability and proposes measures to  
425 eliminate them on uniform meshes.

426 There are many related questions. Is the strategy to uniformly resolve the  
427 Rossby deformation radius in realistic applications a beneficial one? What is the  
428 optimal choice of switching from parameterized to resolved eddies on meshes with  
429 a strong change in resolution? The study of Hallberg (2013) suggests to sharply  
430 switch on/off the eddy parameterization where  $L_R/h = 2$ , and this may still be  
431 a good solution even in the presence of the retardation effect. Any continuation  
432 of thickness diffusion into the fine domain will further damp eddy motions there.  
433 And yet, the consequences of implementing this recommendation on meshes with  
434 strongly varying resolution remain to be explored. In addition, a local mesh re-  
435 finement may modify the mean flow by virtue of the mean divergence of eddy  
436 Reynolds stresses. In simulations here the effects of this type were seen in de-  
437 viation of the time-mean flow from strict zonality, induced by the mere change  
438 in the mesh resolution. Answering these questions would be of general interest.  
439 Indeed, since the Rossby radius varies substantially and becomes rather small at  
440 high latitudes, even current high-resolution ( $1/10^\circ - 1/12^\circ$ ) models are on the edge  
441 between eddy-resolving and eddy-permitting over certain parts of the global ocean,  
442 facing similar questions. The impact of mesh refinement on the representation of  
443 eddy-topography interactions is yet another research topic, because many jets in  
444 the ocean are located in the vicinity of shelf break, where the Rossby radius varies  
445 substantially.

446 Although our study relies on unstructured meshes, situations where the mesh  
447 resolution is varied sharply occur in setups with nesting as well as on orthogonal  
448 curvilinear meshes where poles are taken close to each other to allow refinement in  
449 a particular region. The findings of this study should be in equal degree relevant  
450 in those cases too.

## 451 **5. Conclusions**

452 We show that changing the mesh resolution from coarse to eddy-resolving is  
453 accompanied by retardation in the turbulence saturation because of the absence of  
454 sufficiently strong perturbations in the flow upstream, which makes the effective  
455 geometrical resolution coarser. The effect is noticeable in zones about  $10^\circ$  wide  
456 in the test cases reported here when the coarse mesh is only eddy permitting. It  
457 becomes less significant if the coarse mesh is itself eddy resolving. The resolution  
458 of two mesh intervals per Rossby radius seems to define the boundary.

459 These statements have a qualitative character as the details may depend on  
460 applications. The presence of topography, details of domain geometry, or reduced  
461 velocity shear may modify the manifestation of the 'retardation' effect.

462 We also show that biharmonic viscosity operators with commonly used magni-  
463 tude of biharmonic viscosity, scaled with the cube of the mesh size, are sufficient  
464 to ensure smoothness in the fields of temperature, horizontal velocity and relative  
465 vorticity even for sharp changes in the mesh resolution. However, changing the  
466 mesh resolution may lead to noise in the vertical velocity in the transition zones,  
467 which is linked to details of the vector-invariant momentum advection scheme, and  
468 is not present for the flux form of momentum advection.

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## 471 **References**

- 472 Chanut, J., Barnier, B., Large, W., Debreu, L., Pendluff, T., Molines, J. M.,  
473 Mathiot, P., 2008. Mesoscale eddies in the Labrador Sea and their contribution  
474 to convection and restratification, *J. Phys. Oceanogr.*, 38, 1617–1643.
- 475 Danilov, S., 2012. Two finite-volume unstructured mesh mod-  
476 els for large-scale ocean modeling. *Ocean Model.* 47, 1425.  
477 <http://dx.doi.org/10.1016/j.ocemod.2012.01.004>.
- 478 Danilov, S., Androsov, A., 2015. Cell-vertex discretization of shallow water  
479 equations on mixed unstructured meshes. *Ocean Dynamics*, 65, 33–47. DOI  
480 10.1007/s10236-014-0790-x
- 481 Debreu, L., Blayo, E., 2008. Two-way embedding algorithms: a review. *Ocean*  
482 *Dynamics*, 58, 415–428.
- 483 Durgadoo, J. V., B. R. Loveday, C. J. C. Reason, P. Penven, and A. Biastoch, 2013.  
484 Agulhas leakage predominantly responds to the Southern Hemisphere westerlies.  
485 *J. Phys. Oceanogr.*, 43, 2113–2131.

- 486 Fox-Kemper, B., Menemenlis, D., 2008. Can large eddy simulation techniques  
487 improve mesoscale rich ocean models? In: Hecht, M.W., Hasumi, H. (Eds.), 1679  
488 Ocean Modeling in an Eddying Regime, Geophysical Monograph 177, AGU, pp.  
489 1680 319337.
- 490 Gassmann, A., 2013. A global hexagonal C-grid non-hydrostatic dynamical core  
491 (ICON-IAP) designed for energetic consistency. Q.J.R. Meteorol. Soc. 139, 152-  
492 175.
- 493 Greenberg, D.A., Dupont, F., Lyard, F.H., Lynch, D.R., Werner, F.E., 2007.  
494 Resolution issues in numerical models of oceanic and coastal circulation. Cont.  
495 Shelf Res., 27, 13171343.
- 496 Hallberg, R., 2013. Using a resolution function to regulate parameterizations of  
497 oceanic mesoscale eddy effects. Ocean Modelling, 72, 92–103.
- 498 Hollingsworth, A., Kållberg, P., Renner, V. Burridge, D. M., 1983. An internal  
499 symmetric computational instability. Quart. J. R. Met. Soc., 109, 417–428.
- 500 Jansen, M. F., Held, I. M., 2014. Parameterizing subgrid-scale eddy effects using  
501 energetically consistent backscatter. Ocean Modelling, 80, 36–48.
- 502 Mertens, C., Rhein, M., Walter, M., Böning, C. W., Behrens, E., Kieke, D., Ste-  
503 infeldt, R., Stöber, U., 2014. Circulation and transports in the Newfoundland  
504 Basin, western subpolar North Atlantic. J. Geophys. Res. - Oceans, 119, 7772–  
505 7793.
- 506 Pacanowski, R. C. and Philander, S. G. H., 1981. Parameterization of vertical  
507 mixing in numerical-models of tropical oceans. J. Phys. Oceanogr., 11, 14431451.
- 508 Ringler, T., Petersen, M., Higdon, R., Jacobsen, D., Maltrud, M., Jones, P.W.,  
509 2013. A multi-resolution approach to global ocean modelling. Ocean Model.  
510 Ocean Modelling, 69, 211–232.
- 511 Timmermann, R., Hellmer, H. H., 2013. Southern Ocean warming and in-  
512 creased ice shelf basal melting in the 21st and 22nd centuries based on cou-  
513 pled ice-ocean finite-element modelling, Ocean Dynamics, 63(9), 1011-1026,  
514 doi:10.1007/s10236-013-0642-0.
- 515 Wekerle, C., Wang, Q., Danilov, S., Jung, T., Schroter, J., 2013, The Canadian  
516 Arctic Archipelago throughflow in a multiresolution global model: Model assess-  
517 ment and the driving mechanism of interannual variability, J. Geophys. Res.:  
518 Oceans, doi:10.1002/jgrc.20330.