CHLOROPHYLL TO CARBON RATIO DERIVED FROM AN ECOSYSTEM MODEL WITH EXPLICIT PHOTODAMAGE

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Phytoplankton

RESULTS

CONCLUSIONS

REcoM-2 and the role of photophysiology

REcoM-2 is an **ecosystem model** coupled to the MITgcm. It defines carbon, nitrogen and chlorophyll as state variables, allowing **variable stoichiometry**.

Physical forcing

MIT Global Circulation Model

Circulation Model Light Photosynthesis Photosynthesis Chlorophyll Chla synthesis Biosynthesis Respiration Nitrogen Uptake / Assimilation



Aggregation

Zooplankton



DOM



CONCLUSIONS

REcoM-2 and the role of photophysiology



Parameterization of chlorophyll non-reversible damage



Model accuracy: satellite chlorophyll and literature Chl:C

Chlorophyll Output annual Set k values model run Chl:C ratio climatology **Chlorophyll: satellite** Chl:C ratios: literature annual means at surface upper 200m; n ≤ 100 2000-2015 1990-2014 Year Atlantic Ocean Authors Jakobsen & 2016 Baltic Sea Authors Year Pacific Ocean Markager **OC-CCI** Li et al. 2010 California coastal curr. Buck et al. 1996 North Atlantic North & Equatorial P. Furuya 1990 Marañon Atlantic gyres 2005 Chang et al. 2003 East China Sea 2006 Perez et al. A. subtr. gyres 2003 Equatorial Pacific Brown et al. Caron et al. 1995 Sargaso Sea Cambell et al. 1994 Hawaii Goericke & 1998 Sargaso Sea Jones et al. 1996 Hawaii Welschmeyer 150 300 320 340 360 50 46 30 **REcoM** Compbelli 26 utitude. ranon 30 DOME 2003 -19 -58 -10 -160 -78 -86 80 120 140 180 200 220 240 260 280 320 ań 100 120 140 160 180 220 240 260 280 300 320 340 360 Longitude

Correlation with satellite chlorophyll and literature Chl:C



CONCLUSIONS

Analysis of patterns: Chl:C gradient in depth



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CONCLUSIONS

Analysis of patterns: seasonality at high latitudes



CONCLUSIONS

Analysis of patterns: Chl:C under the ice sheet



Daly (1990) L&O. Chl:C under the ice at 60°S

Summary and conclusions

Optimization of models with surface chlorophyll can be biased towards the description of high light conditions.

Modelling non-reversible damage to chlorophyll as a function of light intensity provides:

- accurate surface chlorophyll fields.
- realistic phytoplankton stoichiometry in conditions not seen by satellites.

Thanks!

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Details physical forcing and ecosystem model



by Schartau 2004.

by Schourup-Kristensen 2014.

Phytoplankton growth model



All phytoplankton

lim = Liebig's law (DIN,Fe) Pmax = Pcm * lim * Tfunc $Photosynth = Pmax*(1-exp((-\alpha * Chl:C * E) / Pmax))$ $N_assim = Vcm * Pmax * Qmax * Ni/(Ni + kdin) * lim(Qmax)$ $Chl_synth = N_assim * Chl:Nmax*(Phot/(Chl:C* \alpha * E)))$ Respiration = Rref * Tfunc +**Biosynth** $*N_assim$

dC = (Phot - Respiration - excretionC) * phyC dN =(N_assim* phyC) - (excretionN*phyN) dChl =(Chl_synth*phyC) - (damageCHL*phychl)

additional for diatoms

dSi=(Si_assim*phyC)-(excretionSi*phySi)

Phytoplankton growth model: processes dependent on light



Phytoplankton growth model: high vs low light

amonific magnitudion note is a linear function of the note of nitrate

Geider et al.(1998) L&O. A *Photosynthesis* 3 dynamic regulatory model. $= P_{max} \times \left(1 - e^{\frac{-\alpha\theta E}{P_{max}}}\right)$ 2 681 oth domamics **Photos** Nitrate Assimilation (g N [g C]⁻¹ d⁻¹) В Α Photosynthesis (d⁻¹) Variable Variable 0 200 400 800 1000 600 N:C N:C Variable Chl:C 3.0 Chla syntesis $= N \ assim \times Chl: N_{max} \times \frac{Phot}{\alpha \theta E}$ 2.5 Chla synthesis Ek Irradiance Nitrate 2.0 Chlorophyll Synthesis (g Chl [g C]⁻¹ d⁻¹) С D 1.5 Respiration (d⁻¹) 1.0 "Cost" of 0.5 Biosynthesis Variable $\frac{E_0}{E_k}$ 0.0 Maintenance 🕈 Respiration 0 200 400 800 1000 600 Nitrate Assimilation Nitrate Assimilation (a N [a C]⁻¹ d⁻¹) $(g N [g C]^{-1} d^{-1})$ 0.12 *Chla damage* = kFig. 1. Graphical summary of the model showing the depen-0.10dencies of photosynthesis, nitrate assimilation, Chl a synthesis, and respiration on environmental and physiological variables (see Table 0.08 2 for mathematical details and the tout for a fuller auriantian). A. Photosynthesis is a saturating 0.06 Light limitation nitial Light saturation slope increases with increasi rate 0.04 increases with increasing N: C. The light-saturation parameter (E_i) Phot is given by the irradiance at which the initial sloper intercepts the 0.02 light-saturated rate. B. The carbon-specific nitrate assimilation rate < 10.00 is a saturating function of nitrate concentration where the maximum uptake rate is downregulated at high values of $N: \mathbb{C}$. The rate of αθΕ 800 0 200 400 600 1000 Chl a synthesis is obligately coupled to protein synthesis and thus to nitrate assimilation. However, the magnitude of the coupling delight (W m-2) pends on the ratio of irradiance to the light-saturation parameter 15 (E_0/E_i) . At a given rate of nitrate assimilation the carbon-specific rate of Chl a synthesis declines as E_0/E_k increases. D. The carbon-

Phytoplankton growth model: steady state solutions



PAR

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Model accuracy: other metrics, annual NPP and export production



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Analysis of patterns: Chla:C under the ice sheet



References field data

Brown SL (2003) Microbial community abundance and biomass along a 180° transect in the equatorial Pacific during an El Niño-Southern Oscillation cold phase. J Geophys Res **108**:8139

Buck KR, Chavez FP, Campbell L (1996) Basin-wide distributions of living carbon components and the inverted trophic pyramid of the central gyre of the North Atlantic Ocean, summer 1993. Aquat Microb Ecol **10**:283–298

Campbell L, Nolla H a., Vaulot D (1994) The importance of Prochlorococcus to community structure in the central North Pacific Ocean. Limnol Oceanogr **39**:954–961

Caron DA, Dam HG, Kremer P, Lessard EJ, Madin LP, Malone TC, Napp JM, Peele ER, Roman MR, Youngbluth MJ (1995) The contribution of microorganisms to particulate carbon and nitrogen in surface waters of the Sargasso Sea near Bermuda. Deep Res Part I **42**:943–972

Chang J, Shiah FK, Gong GC, Chiang KP (2003) Cross-shelf variation in carbon-to-chlorophyll a ratios in the East China Sea, summer 1998. Deep Res Part II Top Stud Oceanogr **50**:1237–1247

Furuya K (1990) Subsurface chlorophyll maximum in the tropical and subtropical western Pacific Ocean: Vertical profiles of phytoplankton biomass and its relationship with chlorophyll a and particulate organic carbon. Mar Biol **107**:529–539

Goericke R (1998) Response of phytoplankton community structure and taxon-specific growth rates to seasonally varying physical forcing in the Sargasso Sea off Bermuda. Limnol Oceanogr **43**:921–935

Jones DR, Karl DM, Laws EA (1996) Growth rates and production of heterotrophic bacteria and phytoplankton in the North Pacific subtropical gyre. Deep Sea Res I **43**:1567–1580

Li QP, Franks PJS, Landry MR, Goericke R, Taylor AG (2010) Modeling phytoplankton growth rates and chlorophyll to carbon ratios in California coastal and pelagic ecosystems. J Geophys Res Biogeosciences **115**:1–12

Marañón E (2005) Phytoplankton growth rates in the Atlantic subtropical gyres. Limnol Oceanogr 50:299-310

Pérez V, Fernández E, Marañón E, Morán XAG, Zubkov M V (2006) Vertical distribution of phytoplankton biomass, production and growth in the Atlantic subtropical gyres. Deep Res I **53**:1616–1634

Jakobsen HH, Markager S (2016) Carbon-to-chlorophyll ratio for phytoplankton in temperate coastal waters: Seasonal patterns and relationship to nutrients. Limnol Oceanogr. **43**:679-694

Daly KL (1990) Overwintering development, growth, and feeding of larval Euphausia superba in the Antarctic marginal ice zone. 35:1564–1576

Sakshaug E, Holm-hansen O (1986) Photoadaptation in Antarctic phytopfankton: Variations in growth rate, chemical composition and P versus I curves. J Plankton Res 8:459–473

Wang X, Murtugudde R, Hackert E, Marañón E (2013) Phytoplankton carbon and chlorophyll distributions in the equatorial Pacific and Atlantic: A basin-scale comparative study. J Mar Syst **109**–110:138–148