# **Bulk Geochemistry of the Sand Fraction from CRP-3 (Victoria Land Basin, Antarctica): Evidence for Provenance and Milankovitch Climatic Fluctuations**

M. Sprovieri\*, A. Bellanca & R. Neri

Dipartimento di Chimica e Fisica della Terra ed Applicazioni alle Georisorse e ai Rischi Naturali (CFTA), Università di Palermo 90123 Palermo, Italy

Received 9 January 2001; accepted in revised form 6 November 2001

**Abstract** - A total of 167 samples distributed throughout the CRP-3 drillhole from 5.77 to 787.68 mbsf and representing fine to coarse sandstones have been analysed by X-ray fluorescence spectrometry (XRF). Bulk sample geochemistry (major and trace elements) indicates a dominant provenance of detritus from the Ferrar Supergroup in the uppermost 200 mbsf of the core. A markedly increased contribution from the Beacon sandstones is recognized below 200 mbsf and down to 600 mbsf. In the lower part of CRP-3, down to 787.68 mbsf, geochemical evidence for influxes of Ferrar materials is again recorded.



On the basis of preliminary magnetostratigraphic data reported for the lower 447

mbsf of the drillhole, we tentatively evaluated the main periodicities modulating the geochemical records. Our results identify a possible influence of the precession, obliquity and long-eccentricity astronomical components (21, 41, and 400 ky frequency bands) on the deposition mechanisms of the studied glaciomarine sediments.

### INTRODUCTION

This paper presents the results of major and trace element analyses performed on fine to coarse sandstones from the CRP-3 drillhole. These results have been discussed to assess the sediment provenance and processed by spectral analysis to test the presence of cyclic patterns in the geochemical records.

When biostratigraphic and magnetostratigraphic constrains are poorly defined, the recognition of statistically significant periodicities in sedimentary records related to astronomical/orbital forcing may be useful to calibrate the time interval of the studied succession (House, 1995).

By analysing physical properties of mudstone and fine-grained sandstone intervals in CRP-2/2A drillhole, Niessen et al. (2000), Cape Roberts Science Team (1999), and Claps et al. (2000) identified highfrequency periodicities which the authors correlated to the Milankovitch periodic orbital forcing. Interesting cyclostratigraphic results have also been obtained for the upper 200 mbsf of the CRP-3 core. Based on combined spectral methodologies, it has been suggested the presence of the three classic Milankovitch periodicities modulating the magnetic susceptibility in this part of the core (Cape Roberts Science Team, 2000).

These results stimulated us to use a similar approach for testing cyclicity in the geochemical

records from the lower part of the CRP-3. While a relatively reliable time framework for the upper part of the CRP-3 core is provided by a good biostratigraphic and magnetostratigraphic control, only approximate temporal constrains are available for the 350 to 789.77 mbsf interval (Cape Roberts Science Team, 2000). Then, the recognition of selected periodicities in the geochemical signals could represent a useful tool to calibrate the time record of the lower part of the CRP-3 core, ascribed to the early Oligocene. Furthermore, the recognition of specific frequency bands related to well-known climatic influences could give the opportunity to study modes and times of possible response of the Oligocene East Antarctic Ice Sheet.

#### SAMPLES AND PROCESSING

Analyses were performed on 167 sand-grained samples scattered throughout the drillhole from 5.77 to 787.68 mbsf, with a mean distribution of one sample every 5 m.

Si, Ti, Al, Fe, Mn, Mg, Ca, Na, K, P and Cr, Ba, La, Ce, V, Zr, Nb, Y, Sr, Rb, Ni were determined by X-ray fluorescence spectrometry (XRF) on pressed, boric-acid backed pellets of bulk rock. Data reduction was achieved using the method described by Franzini et al. (1975). Certified reference materials were used as monitors of data quality. Analytical errors were below 1% for Si, Al, Na; below 3% for Ti, K, Fe, Ca; and below 10% for Mg, Mn, P and trace elements. All samples were washed repeatedly in deionized water prior to analysis to avoid contamination resulting from drilling mud and seawater.

### **RESULTS AND DISCUSSION**

Major and minor element analyses of CRP-3 samples are given in tables 1 and 2 and in figures 1 and 2. The data plotted in the figures are normalized to 100% L.O.I. (Loss On Ignition)-free. In order to discuss the geochemical results in terms of provenance, we assume that, as for the previously studied CRP-1 and CRP-2/2A sequences, main sources for sediments in CRP-3 are the Transantarctic Mountains (TAM) with quartzose sandstones of the Devonian-Triassic Beacon Supergroup, Jurassic dolerites and Kirkpatrick basalts, coarse-grained plutonic rocks (Cambro-Ordovician Granite Harbour Intrusive Complex), and minor metamorphic rocks from the Upper Proterozoic basament.

### DEPTH PROFILES

The depth profile of SiO<sub>2</sub> (Fig. 1) exhibits a marked increase of the element below approximately 160 mbsf. Despite of wide fluctuations, SiO<sub>2</sub> concentrations remain high, generally greater than 80%, throughout a thick interval of the drillhole down to about 600 mbsf. Below 600 mbsf, SiO<sub>2</sub> clearly decreases. The overall very high contents of <sup>2</sup>SiO<sub>2</sub> are indicative of strong influxes of detritus from the Beacon sandstones throughout most of CRP-3. Higher



Fig. 1 - CRP-3 lithologic section with depth profiles for  $SiO_2$ , Ni, V, and Cr.

Tab. 1 - Major element concentrations (wt%) of CRP-3 samples. Data listed are normalised to 100% (hydrous basis).

Depth	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K,O	P2O5	L.O.I
5.77	71.24	0.48	10.51	3.59	0.07	2.52	2.60	4.26	2.14	0.09	2.49
9.90	69.30	0.52	10.52	4.88	0.08	2.76	3.86	2.11	2.16	0.10	3.71
13.26	71.89	0.56	10.96	4.32	0.08	2.81	2.66	2.12	2.21 .	0.09	2.31
21.16	77.99	0.41	8.38	3.75	0.07	2.20	2.82	1.13	1.34	0.08	1.82
24.32	80.02	0.37	7.85	3.44	0.07	2.16	2.72	1.03	1.25	0.07	1.03
24.88	87.75	0.20	3.97	2.32	0.06	1.22	1.81	0.86	0.68	0.04	1.09
28.80	78.98	0.36	7.44	3.26	0.08	2.10	2.82	1.73	1.22	0.07	1.93
33.14	73.42	0.53	10.74	4.38	0.07	2.78	3.08	1.39	1.77	0.09	1.75
47.61	70.24	0.68	12.09	4.68	0.08	2.94	3.36	1.74	2.27	0.13	1.78
49.40	69.83	0.62	11.91	5.04	0.08	2.96	2.97	1.47	2.06	0.10	2.98
67.03	65.92	0.26	7.00	12.08	0.06	1.92	2.09	1.05	0.91	0.00	8.05
09.97	67.62	0.40	9.94	4.02	0.11	2.90	2.07	1.34	2.02	0.11	4.10
78.35	67.59	0.62	12.50	5.08	0.09	3.47	3.57	1.41	2.03	0.11	2.30
83 39	73.82	0.01	10.13	4 43	0.09	3.70	3.23	1.35	1 79	0.10	1 34
89.47	69.55	0.55	11.38	4.67	0.08	3.41	4.14	1.39	2.11	0.09	2.63
94.47	75.94	0.38	7.74	3.22	0.09	2.28	4.73	1.40	1.26	0.07	2.88
99.47	69.78	0.55	11.64	5,13	0.08	4.39	2.55	1.72	2.22	0.09	1.85
103.58	69.39	0.54	11.53	4.72	0.08	3.69	2.61	1.88	2.25	0.10	3.22
110.16	66.89	0.55	11.90	5.13	0.10	4.10	3.65	1.85	2.34	0.10	3.39
115.87	73.41	0.45	10.42	3.93	0.08	3.22	2.19	1.78	1.74	0.07	2.71
123.42	63.73	0.61	12.34	5.38	0.10	3.79	5.25	1.85	2.56	0.11	4.28
130.57	69.37	0.57	11.64	5.24	0.08	4.26	1.84	1.59	2.16	0.08	3.19
135.15	70.87	0.58	11.49	4.83	0.08	3.32	2.00	1.67	2.17	0.08	2.92
139.88	76.44	0.28	6.01	3.00	0.14	1.91	5.33	1.93	0.75	0.04	4.17
145.09	77.99	0.38	5.93	4.73	0.08	3.76	3.56	1.29	0.75	0.05	1.48
148.73	82.62	0.31	4.99	3.90	0.07	3,44	1.69	1.15	0.76	0.04	1.04
155.82	87.25	0.21	4.01	1./1	0.06	2.50	1.04	1.24	0.76	0.03	0.59
150.52	04.01 74.01	0.23	0.52	1.97	0.00	2.70	1.55	1.30	0.95	0.05	1.95
171 75	04.80	0.55	9.33	4.17	0.06	5.99	1.90	0.16	0.24	0.03	1 08
171.95	75 58	0.03	8.29	4.01	0.00	4 4 9	1.50	2.06	1 48	0.01	1.98
177.19	81.41	0.45	5.46	2.54	0.07	2.89	2 75	1 33	1.40	0.00	2.11
177.47	77.66	0.30	7.57	2.38	0.07	3.30	2.51	2.50	1.26	0.08	2.36
183.66	88.50	0.19	3.92	1.65	0.06	1.95	0.79	1.02	0.77	0.04	1.11
189.16	47.29	0.46	9.41	4.52	0.23	3.43	16.25	1.28	2.07	0.29	14.77
189.51	80.19	0.22	6.07	4.36	0.05	1.54	0.67	1.44	1.16	0.04	4.27
196.27	74.16	0.43	9.99	3.89	0.07	4.08	0.90	1.64	2.27	0.06	2.51
202,56	88.67	0.18	4.17	1.55	0.06	1.76	0.60	0.99	0.96	0.03	1.02
207.68	73.76	0.28	7.58	3.17	0.08	5.17	2.12	1.86	1.59	0.04	4.34
210.49	77.68	0.30	6.86	2.52	0.07	3.88	2.34	1.82	1.34	0.04	3.15
215.39	86.73	0.30	4.08	2.56	0.06	2.64	1.08	1.04	0.76	0.03	0.71
226.39	86.98	0.23	3.59	2.50	0.06	2.06	1.69	0.72	0.75	0.04	1.39
229.00	/0./1	0.47	11.38	4.64	0.07	3.84	1.03	1.69	2.48	0.06	3.64
230.42	00.82 84.10	0.23	2.07	1.70	0.13	1.52	13.57	1,22	1.12	0.05	8.39 2.46
233,47	86.55	0.17	5 3 5	2.10	0.08	1.95	4.92	0.80	0.79	0.02	1.07
239.80	90.43	0.23	3.35	1.91	0.05	1.65	0.54	0.61	0.59	0.03	1.07
250.24	92.89	0.17	2.74	1.36	0.05	0.70	0.35	0.50	0.59	0.02	0.62
256.63	83.78	0.28	6.10	2.04	0.07	1.04	2.34	0.81	1.67	0.04	1.83
259.30	85.25	0.28	6.35	1.74	0.06	0.83	1.15	0.75	1.66	0.04	1.90
270.94	93.05	0.13	2.70	0.92	0.05	0.35	0.42	0.66	0.61	0.02	1.09
273.85	82.69	0.22	4.52	1.62	0.08	0.54	4.77	0.71	1.07	0.03	3.75
279.65	77.56	0.27	4.87	2.27	0.08	1.10	7.31	0.84	1.06	0.03	4.60
283.21	83.99	0.33	5.85	2.62	0.06	1.40	1.63	1.19	1.18	0.03	1.72
286.19	80.75	0.33	6.83	3.25	0.06	2.59	1.80	1.15	1.55	0.03	1.65
289.17	87.04	0.21	3.23	2.11	0.07	0.73	3.28	0.62	0.56	0.03	2.13
291.40	70.55	0.36	4.28	3.59	0.11	2.10	11.51	0.75	0.68	0.03	6.03
298.96	77.55	0.22	3.42	4.92	0.10	3.21	8.74	0.68	0.68	0.03	0.44
301.64	86.63	0.22	4.23	3.65	0.06	1.28	0.83	0.69	0.94	0.03	1.45
303.71	09.10 76 01	0.18	3.38 8 77	1.85	0.06	0.82	1.24	0.59	0.82	0.03	1.14
315.50	70.84 70.82	0.30	0.72	5.50	0.07	3./1 5 70	1.22	1.30	2,48	0.04	1.83
313.74	70.02 67.82	0.50	10.30	J.21 5 61	0.08	5.28 5.11	1.15	1.72	2.24	0.07	2.03
326.60	91.82	0.16	3.02	1 23	0.06	1 42	0.21	0.47	0.85	0.07	2.70
333.16	69.35	0.16	2.90	1.2.5	0.00	1.42 2.28	1 34	0.47	0.63	0.02	20.67
335.93	84.64	0.18	3.28	2.11	0.07	3.07	3.07	0.64	0.75	0.02	2.17
341.11	79.64	0.28	5.01	2.84	0.07	2.57	4.30	0.68	1.41	0.04	3.16
345.73	87.79	0.27	5.16	3.00	0.06	0.68	0.75	0.77	0.99	0.05	0.49
348.67	92.71	0.12	2.30	1.24	0.05	0.36	1.38	0.31	0.46	0.02	1.04

Tab.	1	-	Continued.

Depth	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	$P_2O_5$	L.O.1
351.85	87.20	0.29	4.81	0.02	0.06	3.09	0.74	1.03	0.99	0.03	1.74
351.85	83.61	0.27	4.70	3.98	0.05	3.03	0.70	0.96	0.92	0.03	1.74
356.15	84.46	0.21	5.82	2.69	0.06	2.41	0.84	1.00	1.45	0.03	1.03
358.92	83.88	0.24	5.01	2.90	0.06	2.72	1.43	0.89	1.29	0.04	1.54
365.84	80.23	0.15	2.68	1.33	0.08	1.12	7.86	0.38	0.71	0.02	5.42
375 29	82.83	0.12	2.12	1.37	0.08	0.54	7.13	0.33	0.44	0.02	5.05 4.19
381.97	88.46	0.13	3.95	2.75	0.05	0.66	1.07	0.52	0.77	0.02	1.51
385.90	83.77	0.19	3.62	2.13	0.07	0.63	5.08	0.60	0.74	0.03	3.14
390.75	86.11	0.19	3.78	3.32	0.05	2.81	0.59	0.73	1.00	0.02	1.39
392.54	71.12	0.22	3.70	3.90	0.11	3.73	9.30	0.80	0.81	0.03	6.27
396.55	79.98	0.18	2.77	2.24	0.08	1.17	7.91	0.47	0.70	0.04	4.47
401.51 405.70	85.62 73.45	0.17	3,90 6 34	2./1	0.05	2.80	1.08	0.74	0.98	0.03	1.87
412.50	79.11	0.30	5.95	2.95	0.06	4.63	1.51	1.16	1.77	0.03	2.50
413.06	81.85	0.29	5.18	2.99	0.06	4.40	1.05	1.11	1.47	0.03	1.58
419.00	71.68	0.09	1.29	1.81	0.15	0.98	13.53	0.33	0.21	0.07	9.87
422.78	80.87	0.12	1.41	1.68	0.10	1.34	9.41	0.24	0.38	0.02	4.43
426.13	92.45	0.10	2.17	1.37	0.05	1.60	0.52	0.36	0.67	0.02	0.67
426.74	87.30	0.11	2.11	1.39	0.06	1.32	3.66	0.46	0.68	0.02	2.88
432.02	80.01 78.67	0.16	5.31	2.27	0.06	4.09	1.12	0.71	1.04	0.02	1.20
440.24	80.64	0.28	4.76	3.88	0.00	5.06	0.79	1.10	1.67	0.03	1.72
445.20	93.07	0.09	1.38	0.63	0.06	0.27	2.21	0.21	0.28	0.02	1.78
449.70	63.84	0.07	0.90	0.59	0.19	0.27	21.84	0.16	0.20	0.03	11.90
450.21	91.38	0.08	1.27	1.19	0.07	0.32	2.78	0.20	0.29	0.01	2.42
455.84	85.67	0.19	4.23	2.34	0.07	1.33	1.94	0.64	1.07	0.03	2.50
460.10	82.60	0.19	3.55	2.01	0.07	2.99	3.51	0.65	1.30	0.02	3.10
473.58	88.03	0.17	3.01	2 30	0.00	3.05	0.46	0.50	0.90	0.02	1.17
475.31	85.81	0.17	3.96	2.18	0.06	3.07	0.60	0.75	1.47	0.04	1.89
480.63	79.28	0.25	4.58	3.82	0.08	2.55	4.05	0.82	1.47	0.04	3.07
481.19	86.41	0.18	3.04	2.90	0.07	2.15	2.18	0.52	0.65	0.02	1.87
486.00	92.69	0.11	1.71	1.13	0.06	0.33	1.62	0.17	0.42	0.02	1.73
490.91	90.70	0.09	1.48	1.35	0.07	0.39	2.86	0.23	0.37	0.02	2.43
495.12 500.18	90.55	0.03	0.90	0.55	0.05	0.25	0.51	0.13	0.24	0.00	0.96
505.30	97.08	0.04	0.65	0.28	0.05	0.23	0.62	0.12	0.14	0.01	0.73
509.33	85.61	0.12	2.03	1.99	0.09	0.39	4.95	0.23	0.65	0.02	3.92
513.14	91.33	0.11	1.85	0.97	0.07	0.31	2.75	0.24	0.66	0.02	1.68
518.97	80.33	0.14	2.75	2.37	0.08	0.56	7.32	0.40	0.72	0.02	5.31
525.33	83.59	0.08	1.71	1.15	0.07	0.37	7.26	0.16	0.31	0.01	5.30
533.19	95.05 88.98	0.08	1.58	132	0.03	0.28	0.90	0.15	0.23	0.01	3.34
540.07	84.51	0.22	3.17	3.23	0.07	3.23	2.22	0.15	0.89	0.01	1.74
543.82	82.26	0.21	3.72	3.58	0.06	6.12	0.51	1.17	0.99	0.02	1.35
546.10	86.90	0.16	2.89	2.63	0.06	4.10	0.84	0.77	0.76	0.02	0.87
550.00	81.49	0.18	3.64	3.37	0.06	5.39	1.59	0.95	0.85	0.02	2.45
556.17	82.50	0.29	4.78	3.69	0.06	4.09	1.07	1.12	1.04	0.02	1.33
566.07	96.41	0.13	1 44	0.70	0.05	0.43	0.69	0.44	0.00	0.01	0.00
577.10	70.88	0.12	1.47	1.98	0.14	0.28	13.04	0.20	0.33	0.03	11.47
581.97	71.66	0.52	9.36	6.47	0.10	3.34	3.25	1.21	1.54	0.06	2.50
586.78	90.03	0.18	2.62	1.94	0.05	1.42	1.32	0.46	0.59	0.02	1.37
586.78	90.02	0.18	2.58	1.95	0.05	1.44	1.33	0.47	0.59	0.02	1.37
594.21	84.20	0.16	3.38	2.82	0.06	1.88	3.20	0.64	0.92	0.02	2.71
606.51	86.81	0.17	4.42	1.58	0.05	1.94	0.56	0.80	0.53	0.02	3.57
613.25	74.57	0.16	3.30	3.26	0.07	1.60	8.83	0.57	0.79	0.01	6.82
615.34	66.55	0.64	12.99	5.99	0.08	3.17	3.44	1.42	2.36	0.12	3.23
620.43	83.43	0.20	4.96	2.38	0.06	2.43	2.19	1.03	1.13	0.03	2.16
626.15	81.25	0.11	2.34	1.67	0.07	1.17	7.33	0.40	0.63	0.02	5.00
630.35	85.12	0.17	2.91	2.76	0.07	3.50	2.34	0.67	0.54	0.02	1.91
034.15 638.27	81.34 81.47	0.23	4.28	3.88 3.17	0.06	5.67	1.12	1.21	0.69	0.02	1.49
643.08	82.26	0.22	3.56	3.17	0.00	5.00 6.21	0.77	1.30	0.92	0.03	1.59
649.11	76.02	0.27	5.06	4.58	0.06	8.69	1.22	1.44	0.74	0.02	1.90
655.83	73.68	0.31	6.70	5.17	0.07	8.83	0.81	1.56	1.04	0.03	1.80
659.10	76.12	0.24	7.23	4.01	0.07	7.37	0.70	1.67	1.12	0.02	1.44
664.00	84.76	0.22	5.56	2.04	0.06	2.99	0.67	1.05	1,11	0.02	1.52

Tab.	1	~	Continued.

Depth	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	L.O.I
669.63	87.06	0.17	3.31	3.02	0.06	3.93	0.47	0.80	0.80	0.02	0.37
673.25	76.76	0.22	5.04	4.83	0.07	8.47	1.00	1.54	0.59	0.02	1.46
676.29	75.97	0.38	8.37	3.84	0.06	4.81	0.94	1.46	1.42	0.04	2.71
680.99	79.04	0.23	4.95	4.37	0.06	6.79	1.01	1.27	0.79	0.03	1.45
685.33	83.89	0.19	3.73	3.15	0.06	5.32	0.72	0.97	0.64	0.02	1.31
686.49	81.80	0.18	4.20	3.20	0.06	6.58	0.70	1.07	0.93	0.02	1.25
691.11	78.11	0.25	4.81	4.87	0.07	7.78	1.14	1.29	0.65	0.03	0.98
696.37	66.34	0.37	8.35	6.30	0.08	11.50	1.96	1.89	1.12	0.03	2.06
696.74	72.38	0.31	7.78	4.53	0.07	8.78	1.67	1.80	1.19	0.04	1.44
704.84	74.12	0.29	7.36	4.27	0.07	7.81	1.41	1.73	1.22	0.03	1.69
708.45	75.18	0.31	6.71	4.68	0.07	7.55	1.41	1.69	0.89	0.03	1.48
710.63	74.33	0.34	7.13	5.78	0.07	6.53	1.25	1.31	1.05	0.03	2.18
715.50	77.43	0.29	4.73	5.12	0.08	7.24	1.40	1.31	0.65	0.03	1.73
720.87	71.82	0.32	7,31	5.43	0.07	9.16	2.02	1.78	0.93	0.03	1.12
724.05	76.90	0.24	5.37	4.63	0.07	7.91	1.37	1.36	0.73	0.03	1.38
730.84	73.98	0.32	6.30	4.98	0.08	8.93	1.20	1.50	1.07	0.03	1.60
735.36	74.28	0.29	6.34	4.92	0.07	8.50	1.41	1.42	0.93	0.02	1.81
737.32	72.03	0.32	6.55	4.97	0.08	9.21	1.93	1.48	0.86	0.03	2,54
740.86	79.50	0.29	4.56	4.82	0.07	6.46	1.18	1.20	0.60	0.05	1.29
744.85	75.92	0.29	6.74	3.83	0.07	7.30	1.50	1.69	1.06	0.04	1.56
750.04	74.88	0.33	7.54	3.58	0.07	7.08	1.44	1.60	1.40	0.03	2.05
754.84	77.83	0.23	5.87	3.73	0.06	6.14	1.25	1.37	0.88	0.03	2.60
759.71	71.86	0.33	8.12	4.79	0.07	7.43	1.49	1.56	1.27	0.03	3.05
765.03	73.55	0.56	6.98	5.14	0.07	7.68	1.32	1.50	1.26	0.03	1.90
770.23	71.16	0.35	7.47	5.03	0.07	7.21	2.25	1.50	1.04	0.03	3.89
776.58	85.30	0.29	3.57	3.31	0.06	3.32	1.25	0.68	0.63	0.03	1.56
777.67	79.37	0.31	5.35	4.12	0.07	4.64	2.48	1.00	0.97	0.04	1.64
780.51	84.44	0.12	2.19	1.39	0.09	1.47	5.35	0.37	0.50	0.02	4.05
787.68	72.66	0.31	6.95	5.75	0.07	7.45	1.72	1.44	1.44	0.05	2.14

SiO<sub>2</sub> values in the interval between approximately 200 and  $^{2}600$  mbsf account for increased proportions of quartz grains suggested on the basis of sand grain compositional modes (Cape Roberts Science Team, 2000).

Ni and Cr curves closely mirror the SiO<sub>2</sub> profile (Fig. 1). Above 160 mbsf and below 600 mbsf, values mainly fluctuate between 20 and 40 ppm and between 30 and 60 ppm for Ni and Cr, respectively. Concentrations of these elements are markedly lower through the interval between 160 and 600 mbsf. Given that modal investigations in CRP-3 have revealed the absence of alkaline volcanic lithic grains and ferromagnesian minerals typical of the McMurdo Volcanic Group (Smellie, this volume), the higher contents of Ni and Cr in the sediments above 160 mbsf and below 600 mbsf should be derived from the Ferrar dolerite. These elements average 61 and 96 ppm in the Ferrar dolerite, whereas they are very depleted (6 and 18 ppm) in the Beacon Sandstone (Roser & Pyne, 1989).

The vanadium profile (Fig. 1) is similar to those of Ni and Cr and supports the suggestion of an influence of Ferrar detritus in the upper and lower part of the CRP-3 core and a dominant input of Beacon materials in the rest of the drillhole. In fact, V is tipically enriched in the Ferrar dolerite (283 ppm) and depleted in the Lower Beacon (25 ppm; Roser & Pyne, 1989).

Also Rb/Sr ratios, that are lower (on average 0.42) above approximately 200 and below 600 mbsf

(Fig. 2) are compatible with a dominant Ferrar (Rb/Sr=0.34) provenance of the detritus in these stratigraphic intervals. Throughout the rest of the drillhole, the Rb/Sr ratios widely fluctuate with averagely higher values (0.88) reflecting a striking influence of the Beacon sandstone (Rb/Sr=1.17; Roser & Pyne, 1989).

Our suggestions on the sediment provenance based on geochemical proxies from the CRP-3 drillhole are consistent with the results from sandstone detrital modes (Smellie, this volume) and clay mineral assemblages (Ehrmann, this volume).

Concentrations of zirconium (Fig. 2) exhibit high variability in CRP-3 sediments and the stratigraphic profile of this element does not correlate with those of SiO<sub>2</sub>, Ni, Cr, V, and Rb/Sr. A reason for this could be the similar contents of Zr in the two main sources, Beacon and Ferrar (142 ppm for both sources; Roser & Pyne, 1989).

Ca values in CRP-3 (Fig. 2) are mostly about 2% but many anomalously high concentrations, measured between 200 and 600 mbsf, reveal the occurrence of diagenetic carbonate.

### SPECTRAL ANALYSIS AND TIME FRAMEWORK

An ensemble of geochemical signals obtained from the CRP-3 core were processed by spectral analysis. The original adopted sampling rate (on average 1 sample/5 m) did not allow us to extract from the geochemical signals of the upper part of the

# M. Sprovieri et al.

Tab,	2 -	Trace	element	concentrations	(ppm)	of	CRP-3	samples.
------	-----	-------	---------	----------------	-------	----	-------	----------

		·										
Depth (mbsf)	Rb	Sr	Zr	Cr	v	Ni	Ba	La	Ce	Y	Nb	Rb/Sr
										-		
5.77	47	74	72	43	96	20	328	19	32	0	15	0.626
9.90	40	76	76	45	102	21	324	21	46	0	16	0.523
1.3.20	49	76	8,5	53 26	112	21	361	21	37	0	17	0,644
24.30	40	85	212	30	97	22	208	10	24	22	13	0.527
24.32	18	30	107	24	57	14	150	6	24	10	8	0.323
28.80	34	63	139	34	88	19	232	13	22	19	12	0.529
33.14	50	83	164	47	112	23	317	16	24	35	15	0.594
47.61	63	123	154	52	124	33	379	26	60	0	19	0.510
49.40	60	90	144	59	130	29	373	30	57	0	17	0.663
67.03	27	61	94	25	65	18	176	15	20	11	10	0.447
69.97	52	95	136	47	115	23	318	27	29	31	13	0.553
73.40	44	71	37	59	142	34	377	24	59	11	17	0.625
78.35	70	136	193	58	141	36	381	16	34	31	15	0.518
83.39	51	124	168	52	115	34	353	18	45	24	12	0,414
89.47	55	140	154	57	121	32	371	21	56	3	13	0.390
94.47	36	189	217	38	98	21	230	14	30	26	11	0.188
99.47	85	244	231	52	121	32	391	20	40	37	15	0.346
103.58	73	207	178	49	115	27	393	25	53	31	13	0.353
110.16	76	241	144	51	115	27	408	27	63	33	14	0.315
115.87	56	161	185	42	96	22	340	15	39	24	13	0.345
123.42	09	198	70	58	122	31	451	36	69	0	18	0.348
135.15	70	208	212	50	131	20	370	24	44	27	15	0.371
139.88	19	208	140	36	75	20	505	23	45	27	15	0.555
145.09	28	122	202	30	124	20	177	2		21	q	0.075
148.73	24	123	240	40	110	23	155	3	3	18	10	0.199
155.82	22	120	186	25	57	13	152	8	4	13	10	0.185
156.52	24	186	84	28	69	16	185	7	7	0	10	0,132
168.94	61	171	205	37	120	23	321	14	29	25	13	0,359
171.75	8	12	70	6	14	4	63	2	0	7	8	0,690
171.95	50	152	219	40	106	22	252	18	28	22	13	0.326
177.19	38	320	146	22	60	17	211	14	15	27	7	0.120
177.47	50	123	279	35	72	17	254	13	17	24	12	0,410
183.66	8	77	17	13	45	11	145	4	1	7	14	0.102
189.16	99	226	155	44	95	24	351	36	55	50	13	0,439
189.51	31	319	85	25	54	13	165	8	6	8	7	0.097
196.27	74	202	177	41	96	21	326	18	30	26	15	0.369
202.56	20	104	74	21	44	13	184	6	0	0	11	0.187
207.08	34	45	98	34	91	20	207	/	18	0	14	0.744
210.49	49	61 82	273	33	11	17	224	31	13	29	11	0.005
226.39	30	37	121	24	71	15	134	5	10	13	10	0.235
229.00	92	69	121	44	119	31	348	14	32	18	13	1331
230.42	29	136	149	41	63	19	194	12	3	25	11	0.215
235.47	22	94	123	29	50	15	405	6	4	23	8	0.235
239.86	35	34	104	30	78	17	171	6	0	35	11	1.030
244.83	20	32	117	18	54	11	118	0	0	10	10	0.620
250.24	23	30	161	23	58	10	117	3	6	14	9	0.780
256.63	63	68	291	29	68	14	212	4	9	15	11	0.924
259.30	56	47	161	27	66	16	222	9	10	15	12	1.191
270.94	14	51	81	14	39	10	138	2	0	16	9	0.269
273.85	39	89	149	22	55	11	177	8	4	16	10	0.439
279.65	40	87	228	42	88	17	180	7	10	14	9	0.452
283.21	24	93	/0	48	117	19	412	/	2	0	14	0.263
280.17	18	43	163	47	65	25	215	4	2	9	10	0.589
201.40	21	45	178	2J 54	152	20	125	3	10	47	12	0.419
298.96	19	38	81	27	80	16	132	7	12	4	12	0.472
301.64	22	29	91	24	100	10	144	4	1	0	13	0.743
305.71	32	39	135	21	64	7	144	2	0	9	9	0.817
312.35	92	62	200	38	100	20	281	9	12	18	11	1.476
315.74	94	70	181	46	126	25	324	21	37	23	14	1.339
320.60	77	53	72	50	134	30	390	20	48	0	18	1.460
326.64	17	14	18	16	33	8	126	2	0	6	12	1.169
333.16	21	48	132	16	57	16	174	5	2	23	· 11	0.439
335.93	28	43	257	24	58	12	126	3	1	15	10	0.646
341.11	62	55	211	31	73	18	208	7	4	22	11	1.123
345.73	15	36	46	29	79	13	158	9	0	8	14	0.414
348.67	9	17	47	11	31	9	108	7	0	0	13	0.531
351.85	29	89	105	45	108	23	281	12	9	8	8	0.330
350.15	63	80	131	33	68	17	245	4	11	10	9	0.796
330.92	20	24	152	35	83	25	199	2	11	11	9	1.038
369.61	12	32 28	1/9	12	30 20	11	132	د ء	8 12	10	10	0.780
375.29	12	20 26	200 20	10 70	20 52	10	90	כ ד	12	10	0	0.441
381.97	18	20	90	47 36	55 70	∠3 12	99 173	0	0	10	9 11	0.437
385.90	24	41	110	37	73	14	123	7 5	10	12	0	0.730
390.75	39	41	142	39	84	17	158	3	0	10	10	0.964
392.54	33	46	83	41	91	25	152	14	24	44	10	0.719
396.55	22	32	102	29	58	19	118	5	1	18	10	0.697
401.51	33	40	101	31	76	17	160	6	7	14	9	0.817
405.70	91	68	195	39	119	28	264	9	10	15	10	1.343
412.50	83	66	277	34	87	19	246	5	7	18	12	1.248
413.06	58	51	194	29	84	17	219	9	15	19	14	1.142
419.00	3	42	45	371	30	11	80	4	0	0	11	0.082

## Tab. 2 - Continued.

Depth (mbsf)	Rb	Sr	Zr	Cr	v	Ni	Ba	La	Ce	Y	Nb	Rb/Sr
422.78	8	17	14)	27	52	18	82	6	3	30	10	0.470
426.13	23	22	98	17	40	10	113	7	0	9	9	1.056
426.74	20	18	72	18	42	11	108	5	0	37	11	1.102
432.02	38	24	128	27	73	20	140	3	0	14	10	1.603
437.09	102	44	246	35	104	21	210	8	5	9	11	2.330
440.24	66	34	176	36	99	20	191	5	8	30	12	1.951
445.20	9	19	98	17	34	8	73	0	-8	13	8	0.465
449.70	9	34	65	21	25	8	59	0	3	24	8	0.254
450.21	9	18	64	13	28	9	91	1	2	10	10	0.494
455.84	28	29	85	27	64	16	154	2	10	0	11	0.973
460.10	55	37	171	27	63	17	177	7	9	11	9	1.511
466.34	33	26	154	21	50	11	148	3	3	21	9	1.270
473.58	25	21	119	11	43	10	102	13	8	10	10	1.167
475.31	64	41	204	25	66	22	209	5	0	10	10	1.575
480.63	69	37	116	47	100	39	172	9	7	12	9	1.847
481.19	22	18	116	24	77	26	100	5	-8	22	12	1.198
486.00	12	9	70	15	37	10	74	0	0	12	9	1.236
490.91	15	15	80	17	35	4	81	/	3	2	9	1161
495.12	, c	0	70	4	0	4	30	1	0	5	8	0.878
500.15	6	0	55	4	0	<i>4</i>	39	0	0	7	8	0.878
500.30	25	3 77	152	20	14	0	49	4	3	9	9	1143
513.14	18	12	74	16	30	24	93	3	5	Ó	ú	1.583
518.97	25	29	114	30	52	12	132	5	1	23	11	0.876
525.33	16	27	76	17	45	8	61	2	2	6	8	0.582
531.78	7	7	71	12	19	5	61	3	6	9	9	1.013
533.19	12	й П	47	11	21	7	84	1	-9	9	9	1.073
540,07	39	31	121	45	100	25	110	2	0	15	9	1.260
543.82	34	26	114	59	104	26	141	1	4	27	11	1.302
546.10	23	22	94	38	73	23	118	4	3	0	12	1.064
550.00	30	34	117	47	93	30	154	6	2	29	11	0.875
556.17	28	39	107	65	133	33	172	0	0	0	12	0.708
560.89	17	24	99	20	45	7	110	3	0	5	11	0.732
566.07	15	12	77	13	23	11	155	1	0	12	9	1.325
577.10	6	31	100	383	33	10	83	_ 3	0	0	13	0.184
581.97	54	67	104	114	193	46	317	5	14	16	11	0.796
586.78	25	29	134	32	70	17	114	8	0	12	9	0.863
594.21	31	39	87	29	69	13	160	5	0	15	12	0.803
602.09	44	37	120	26	66	23	206	6	1	18	10	1.174
606.51	12	26	68	13	26	6	83	0	/	9	12	0.471
613.25	35	63	155	33	61	15	170	4	12	1.5	8	0.551
615.34	39	54	160	29	63	30	177	11	14	14	9	0.724
620.43	40	40	152	34	34	20	203	0	10	12	9	0.087
620.15	13	49	111	18	30	13	122	2	0	15	12	0.412
634.73	20	48	143	55	117	37	161	4	13	13	10	0.415
638.27	20	40 60	142	40	95	24	196	9	5	16	10	0.480
643.08	20	45	169	52	109	31	136	4	7	16	9	0.441
649.11	31	67	181	84	146	44	181	7	5	21	10	0.455
655.83	31	59	156	91	165	43	180	10	0	9	12	0.515
659.10	46	83	170	54	111	31	204	7	10	15	10	0.554
664.00	35	68	140	36	72	22	206	10	8	28	11	0.517
669.63	16	30	137	45	85	27	132	5	0	27	11	0.525
673.25	18	50	115	58	131	33	123	5	9	0	11	0.359
676.29	57	84	226	41	111	21	247	7	16	26	12	0.686
680.99	24	62	129	73	131	40	184	3	12	32	12	0.388
685.33	21	46	111	40	90	26	144	7	0	34	11	0.447
686.49	16	36	118	37	90	24	144	3	0	42	12	0.450
691.11	23	62	190	70	147	41	165	9	11	16	10	0.308
696.37	30	89 52	1.54	/0	181	44	200	7	14	29	15	0.405
090.74	22	32	33	52	140	33	251	10	17	21	0	0.440
704.64	27	73	146	51	140	36	202	6	13	28	13	0.363
710.63	42	01	175	53	120	30	211	1	14	23	10	0.475
715.50	23	70	185	74	150	41	165	8	8	18	9	0.325
720.87	36	104	179	72	172	42	201	10	22	20	10	0.341
724.05	30	86	176	69	147	39	181	5	0	16	9	0.349
730.84	25	66	150	84	170	40	188	12	16	28	11	0.377
735.36	33	92	235	69	156	34	183	12	17	17	9	0.362
737.32	22	65	125	86	168	36	185	9	13	0	12	0.343
740.86	20	66	217	56	133	31	147	15	10	19	10	0.302
744.85	26	71	113	51	123	30	233	11	12	0	14	0.362
750.04	40	87	167	46	113	26	254	10	5	34	11	0.457
754.84	23	60	86	38	102	25	189	8	I	0	12	0.382
759.71	51	107	153	64	136	33	233	6	0	16	10	0.479
765.03	34	68	186	85	183	32	211	6	15	0	15	0.491
770.23	35	87	134	56	135	29	216	6	14	52	12	0.403
776.58	21	49	130	27	85	21	135	8	23	18	11	0.424
777.67	27	53	112	47	111	26	162	9	11	5	12	0.512
780.51	9	26	73	16	38	11	86	1	0	43	9	0.358
/87.68	51	81	133	/1	131	ا ز	181	<u>ح</u>	21	14	10	0.024



Fig. 2 - CRP-3 lithologic section with depth profiles for Rb/Sr, Zr, and Ca. Symbols as in figure 1.

core (0-340 mbsf) the periodicity previously recognized in different sedimentary physical properties by Cape Roberts Team (2000). These scientists estimated that the short-eccentricity orbital component can be recorded in the magnetic susceptibility and porosity signals of the upper part of CRP-3 and calculated a wavelength of about 5.8-9.7m for this orbital periodicity band. Such a result highlights that the sampling rate adopted for the geochemical analysis of the whole core (average of 1 sample/5 m) is lower than the Nyquist frequency and that all the classic Milankovitch periodicites are not identifiable in the geochemical records.

For the lower part of the CRP-3 core (below 340 mbsf), poorly constrained biostratigraphic and palaeomagnetic data, together with the presence of possible sedimentary hiatuses and abrupt variations of the sedimentation rate, limit the interpretation of the results of the spectral analyses. However, a preliminary study of the dominant periodicities characterizing selected geochemical signals is here proposed taking into account that the obtained results have to be considered only as pilot tools for further researches.

A two steps strategy has been adopted to interpret

the results of spectral analysis.

Power spectra have been estimated in the space domain on selected geochemical signals. Hierarchical patterns among the dominant periodicity bands recorded in the spectra have been compared with the ratios measured among the three Milankovitch frequencies (related to the precession, obliquity and eccentricity cycles) during the Cenozoic in order to preliminarily verify a hypothetical interaction between the studied cyclical sedimentary system and the orbital forcing.

Then, an average sedimentation rate has been estimated for the interval 340.8-627.3 mbsf, characterized by the palaeomagnetic chron C13n (Florindo et al., this volume) which covers the time interval between 33.058 and 33.545 Ma (Cande and Kent, 1995). The estimated average sedimentation rate allowed the transformation of the main frequency peaks present in the spectra (space domain) in time periodicities (time domain).

### NUMERICAL METHODOLOGY

Data have been re-sampled with a 5 m constant rate using a cubic spline interpolation method to avoid high frequency bias on the original data.

Long-term trends have been previously subtracted to the original record by means of a 5 points Gaussian filter.

Welch's averaged periodogram method (Welch, 1967) was used for the estimation of the Fourier Power Spectra.

# POWER SPECTRA, NUMERICAL FILTERING AND ORBITALLY-CONTROLLED SEDIMENTARY PROCESSES

Data of SiO<sub>2</sub>, Rb/Sr, and Ni for the CRP-3 drillhole have been processed by spectral analyses. These three geochemical signals have been chosen because they appear to be good tracers of sediment provenance for the studied area (see above). In particular, higher SiO<sub>2</sub> values and Rb/Sr ratios are thought to indicate a dominant Beacon Supergroup influence, whereas higher Ni contents should represent a more significant detrital input from the Ferrar Supergroup.

In the power spectra (Fig. 3), the y-axis indicates the variance associated to each harmonic component, while the x-axis refers to frequencies in cycles/meter. Peaks that are statistically significant (passing the 95% confidence level) are labelled with length of periodicity in meters.

The power spectra estimated for the three signals, in the 340-789.77 mbsf interval, show a very similar frequency structure and the presence of three main peaks at about 0.003, 0.029, and 0.052 cycles/meter (Fig. 3). The high variance and quite narrow frequency bands present in the spectra suggest a clear cyclic forcing process, which might have modulated the sedimentary record and the relative geochemical records. Moreover, the spectral results demonstrate the existence of a regular cyclic pattern in the studied signals and allow us to discard a stochastic mechanism as a controlling factor of their fluctuations.

If we normalize the above-mentioned three frequency bands to the highest one, we obtain a hierarchical pattern (0.058-0.047, 0.36-0.55, 1) similar to that calculated for the classic Milankovitch frequency bands (1/400 cycle/ky for the long-term eccentricity, 1/41-1/54 cycle/ky for the obliquity and 1/19-1/23 cycle/ky for the precession) as indicated by Berger and Loutre (1994) for the Cenozoic.

Such a preliminary result, highlighting a good correspondence of the hierarchical patterns between the two sets of periodicity bands, suggests a possible direct link between astronomical forcing and the sedimentary response in the glaciomarine environment at the margin of the Antarctic Ice Sheet during the early Oligocene.

Considering the average 0.59 m/ky sedimentation rate estimated within the interval characterized by the presence of the chron C13n and assuming it approximately constant along all the lower part of the core (from 340.8 to 789.77 mbsf), we could transform the three frequency peaks (corresponding to the wavelengths of 330, 34, and 19 meters reported in Fig. 3) in periodicities of 560 ky, 57 ky, and 32 ky, respectively.We can hypothesize that the recorded cyclicities are essentially governed by external forcing not strictly related to the Milankovitch astronomical modulation. However, such a hypothesis seems to be discarded considering the good match between the hierarchical patterns calculated among the orbital parameters and the set of periodicities modulating the studied sedimentary record. Alternatively, we could consider the main frequency peaks present in the power spectra as truly corresponding to the orbital periodicities of precession, obliquity and longeccentricity, respectively. Thus, the observed



*Fig. 3* - Power Spectral Density of three selected geochemical signals (SiO<sub>2</sub>, Rb/Sr and Ni) in the 350 to 787 mbsf interval of the CRP-3 core. Main peaks are labelled with the calculated periodicities in metres. B.W. and C.I. represent the Band-Width and the Confidence Interval of the power spectra.

discrepancies between the couples of periodicity bands (in the orbital and sedimentary record) could be explained by hypothesizing that short sedimentary intervals (and, then, time records) are missing. The fault recognized at 539 mbsf (Cape Roberts Science Team, 2000) represents one of the points, along the lower CRP-3 core, where the space/time record could be lost. However, abrupt variations in the sedimentation rate and/or other short intervals of interruption in the sedimentation (not evident by macroscopic analysis of the core) could justify the lack of the records.

In particular, considering the periodicity band of the obliquity and precession cycles and the time span of the palaeomagnetic chron C13n, we can estimate that about 3 obliquity cycles and 8 precession cycles are missing in this segment. The deficiency of other stratigraphic constrains in the lowermost part of the core limits the possibility to evaluate the amount of time missing in this part of the CRP-3 record.

#### CONCLUSIONS

Distribution of major and trace elements suggests changes in the predominant source of detritus throughout the CRP-3 sequence. The upper 200 mbsf of sediments are characterized by a strong influence of Ferrar Supergroup whereas, between 200 and 600 mbsf, provenance from Devonian Beacon Supergroup detritus becomes dominant. From ~600 down to ~788 mbsf, geochemistry of CRP-3 sediments appears to be again controlled by Ferrar detritus input.

Preliminary results obtained by spectral analyses of three selected geochemical signals (SiO<sub>2</sub>, Rb/Sr and Ni) from the lower 447 mbsf of CRP-3 core suggest that the studied sedimentary system has been mainly forced by the well-known Milankovitch astronomical components of long-term eccentricity, obliquity, and precession. Sediments appear to be influenced by cyclic alternation of Devonian Beacon Supergroup detritus and Ferrar Supergroup materials with consequent modulation of the three long intervals characterized by dominant inputs of the different source rocks. Further detailed researches, supported by a multidisciplinary approach, will enable the identification of the sedimentary mechanisms, which drove the deposition of sediments with periodic alternations of geochemical tracers.

ACKNOWLEDGEMENTS - We are grateful to Massimo Pompilio for his help for obtaining CRP-3 samples. We would like to thank L. Krissek and M. Claps for their reviews of an earlier version of the paper. This research was supported by the Italian *Programma Nazionale di Ricerche in Antartide* (PNRA).

#### REFERENCES

- Berger A., 1984. Accuracy and frequency stability of the Earth's orbital elements during the Quaternary. In: Berger A.L., Imbric J., Hays J., Kukla G. & Saltzman B. (eds.), *Milankovitch and Climate, Part 1*, Reidel Publ. Co., Dordrecht, 3-39.
- Berger A. & Loutre M.F., 1994. Astronomical forcing through geological time. In: de Boer P.L. & Smith D.G. (eds.), Orbital forcing and cyclic sequences, IAS Special Publication, 19, 15-24.
- Cande S. & Kent D., 1995. Revised calibration of the geomagnetic polarity time scale for the late Cretaceous and Cenozoic. J. Geophys. Res., 97, 13917-13951.
- Cape Roberts Science Team, 1999. Studies from the Cape Roberts Project, Ross Sea, Antarctica. Initial Report on CRP-2/2A. *Terra Antartica*, 6, 1-173.
- Cape Roberts Science Team, 2000. Studies from the Cape Roberts Project, Ross Sea, Antarctica. Initial Report on CRP-3. *Terra Antartica*, **7**, 1-209.
- Claps M., Niessen F. & Florindo F., 2000. High-frequency analysis of physical properties from CRP-2/2a and implication for sedimentation rate. *Terra Antartica*, **7**, 379-388.
- Ehrmann W., 2001. Variations in smectite content and crystallinity in sediments from CRP-3, Victoria Land Basin, Antarctica. This volume.
- Florindo F., Wilson G.S., Roberts A.P., Sagnotti L. & Verosub K.L., 2001. Magnetostratigraphy of late Eocene - early Oligocene strata from the CRP-3 core, Victoria Land Basin, Antarctica. This volume.
- Franzini M., Leoni L. & Saitta M., 1975. Revisione di una metodologia analitica per fluorescenza X basata sulla correzione completa degli effetti di matrice. *Rend. Soc. Ital. Miner. Petrol.*, 21, 99-108.
- House M.R., 1995. Orbital forcing timescales. In: House M.R. & Gale A.S. (eds.), *Orbital Forcing Timescales and Cyclostratigraphy*, The Geological Society of London, London, 1-18.
- Niessen F., Kopsch K. & Polozek K., 2000. Velocity and porosity from CRP-2/2A Core Logs, Victoria Land Basin, Antarctica. *Terra Antartica*, 7, 241-253.
- Roser B.P. & Pyne A.R., 1989. Wholerock geochemistry. In: P.J. Barrett (ed.). Antarctic Cenozoic History from the CIROS-1 Drillhole, McMurdo Sound. DSIR Bull., 245, 175-184.
- Smellie J.L., 2001. History of Oligocene erosion, uplift and unroofing of the Transantarctic Mountains deduced from sandstone detrital modes in CRP-3 drillcore, Victoria Land Basin, Antarctica. This volume.
- Welch P.D., 1967, The use of Fast Fourier Transform for the Estimation of Power Spectra: A method based on time averaging over short, modified periodograms. IEEE trans. Audio and Electroacoustics, vol. AU-15, no. 2.