Steps towards a global model of photochemical cycling of iron

Ying Ye and Christoph Völker

Alfred Wegener Institute, Helmholtz Centre for Polar and Marine Research

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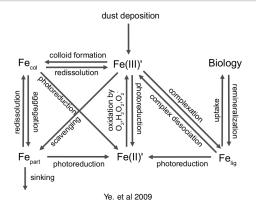
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Redox reactions implemented first in 1D models: Weber et al. (2005, 2007) and Ye et al. (2009)

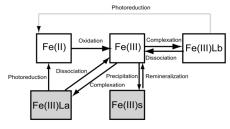


- photochemical production of O₂⁻: proportional to irradiance;
- O₂ and H₂O₂ concentration fixed;
- no direct link to CDOM

Ying Ye and Christoph Völker (AWI)

Implementation in global models

- Tagliabue et al. (2009): first order impact of light and temperature on Fe speciation
- Tagliabeu and Völker (2011): numerical problem solved for different time steps of reactions in the iron cycle
 - fast reactions in equillibrium: redox and organic complexation
 - slow reactions: scavenging, uptake and remineralisation
- oxidation by O₂ considered but not that by H₂O₂ and O₂⁻



Tagliabue et al. (2009)

First step: offline calculation of redox species

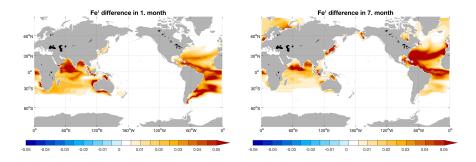
- redox species approach equilibrium;
- model output of DFe, total ligand and irradiance used as input;
- output species: Fe(III), Fe(II), FeL, L', O₂⁻
- rate constants derived first from measurements at 25°C

$$\frac{\partial}{\partial t} Fe(II) = k1_{red} \cdot Fe(III) + k2_{red} \cdot FeC + k3_{red} \cdot FeL -(k_{ox}^{O_2} + k_{ox}^{H_2O_2} \cdot H_2O_2 + k_{ox}^{O_2^-} \cdot O_2^-) \cdot Fe(II)$$

 \sim

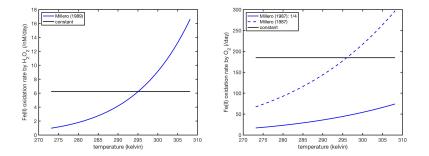
Effect of light on iron speciation: R_{ir0} and R_{const}

- two types of reactions depend on light: photoreduction of Fe(III), FeC and FeL, and production of O₂⁻
- photochemical reactions result in higher concentration of free Fe in tropical and subtropical Atlantic and Indian Ocean

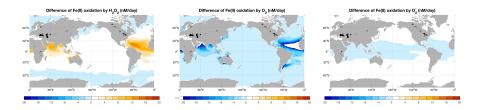


Temperature-dependent Fe(II) oxidation: R_{const} and R_{temp}

- functions fitted based on measurements at different temperatures (Millero and Sotolongo, 1989; Millero et al. 1987);
- k^{O2}_{OX} is assumed to be 1/4 of measured rates (Millero and Sotolongo, 1989; Moffet and Zika, 1987).

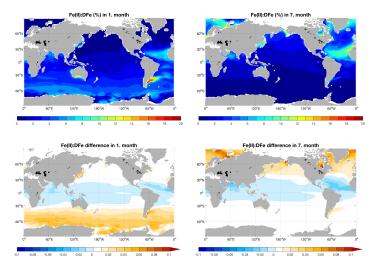


Effect on Fe(II) oxidation by H_2O_2 , O_2 and O_2^-



- oxidation by H₂O₂ dominates in the two runs R_{const} and R_{temp};
- oxidation by H₂O₂ increases at lower and decreases in higher latitudes
- oxidation by O₂ decreases, the stronger decrease at low latitudes is caused by the competition with H₂O₂ and lower O₂ saturation concentration;
- oxidation by O₂- decreases at lower latitudes and increases slightly in higher latitudes: oxidation by H₂O₂ and O₂ decreases in colder regions leading to more Fe(II) available for oxidation by O₂⁻

Total effect on Fe(II) fraction



 \rightarrow strong increase in summer: high photoreduction of Fe(III) + lower oxidation

Effect of temporal and spatial variability of O₂⁻: R_{cdom}

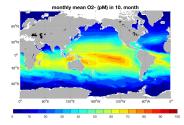
- R_{const}: related to irradiance;
- main process producing O₂⁻: CDOM photochemical degradation
- equation of CDOM degradation according to Dutkiewicz et al. (2015)

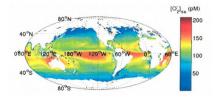
$$\frac{\partial}{\partial t}O_2^- = r_{phot}^{CDOM} \cdot MIN(\frac{PAR}{k_{phot}}, 1.0) \cdot CDOM$$

 r_{phot}^{CDOM} : photochemical degradation rate of CDOM; k_{phot} : light level for bleaching CDOM

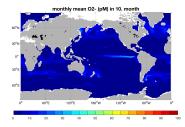
modelled spatial variability of O_2^-

R_{temp}



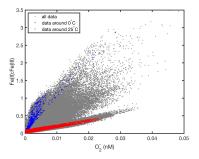






- Powers and Miller (2014): H₂O₂ production estimated from satellite data, dismutation and additional first-order sink of O₂⁻;
- our calculation: constant H₂O₂ of 100 nM; O₂⁻ production estimated from CDOM photochemical degradation; dismutation and redox reaction with Fe and Cu as sink;
- midday steady state concentration compared with monthly averaged concentration!

Fe(II):Fe(III) ratio as a function of O_2^- concentration



- Fe(II):Fe(III) increases with O₂⁻
- temperature controls the slope

Things that need to be discussed and/or tested in sensitivity runs

- role of H₂O₂ spatial variability
- role of Cu(I)/Cu(II) (so far constant total Cu of 1 nM used)
- O₂⁻ source from CDOM degradation
- uncertainties in assumptions of rate constants and their dependence on temperature and pH

After this: implementing into 3D global model! $\ddot{-}$