

Paleo-ice sheet reconstructions constrained by glacial isostatic adjustment and geological data



GERMAN CLIMATE MODELING INITIATIVE

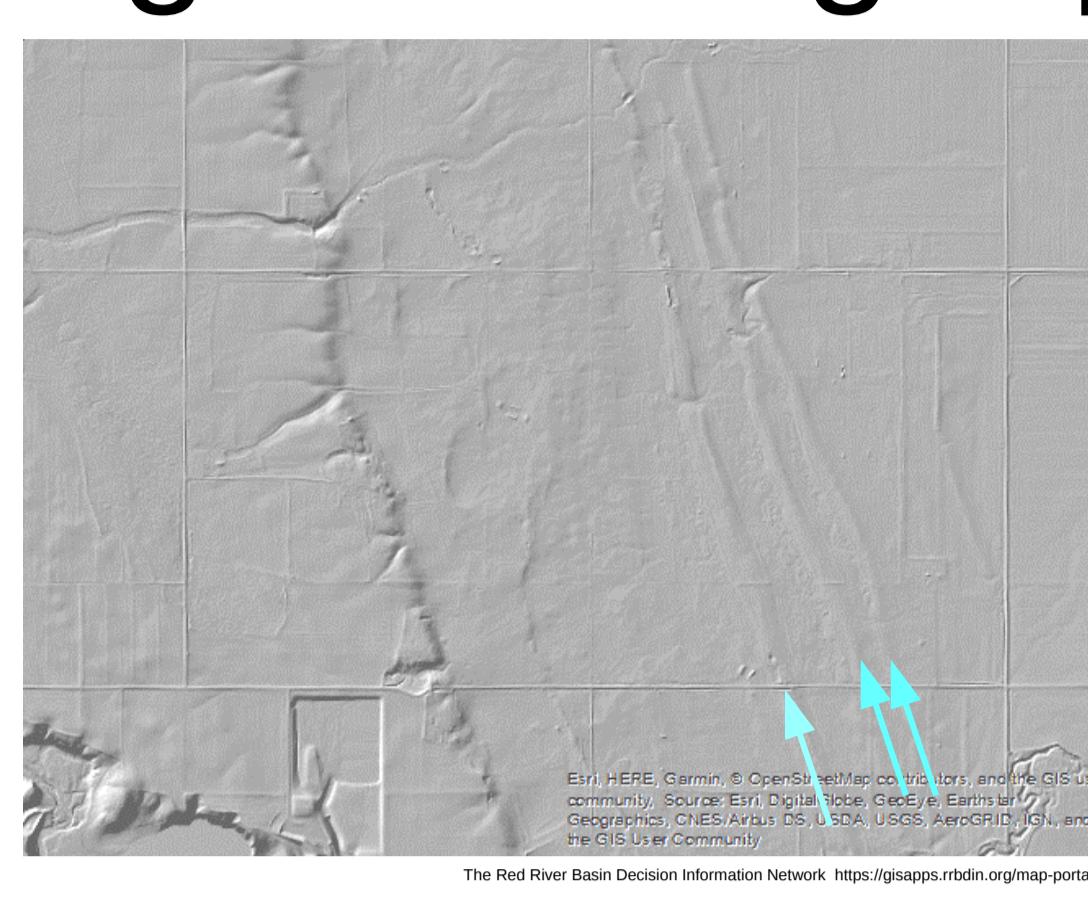
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HELMHOLTZ-ZENTRUM FÜR POLAR- UND MEERESFORSCHUNG

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Ice sheet reconstructions based on geological and geophysical information

- Geophysical modelling of glacial-isostatic adjustment (GIA) processes has long been used to reconstruct paleo-ice sheets (e.g. Tarasov et al 2012, Peltier et al 2015, Gowan et al 2016a, Lambeck et al 2017). In order to do this efficiently, it is necessary to have strict control on the geometry of the ice sheet.
- These data have limitations due to the spatial distribution (i.e. sea level indicators are only located in coastal regions, glacial lake strandlines exist only in paleo-lake basins, end moraines are only located where a glacial margin remained stationary for some time).
- Ultimately, the reconstruction should have at least a minimal amount of glaciological realism. This can be achieved using our model, ICESHEET (Gowan et al 2016b), which uses perfectly plastic rheology.



Strandlines of Lake Agassiz



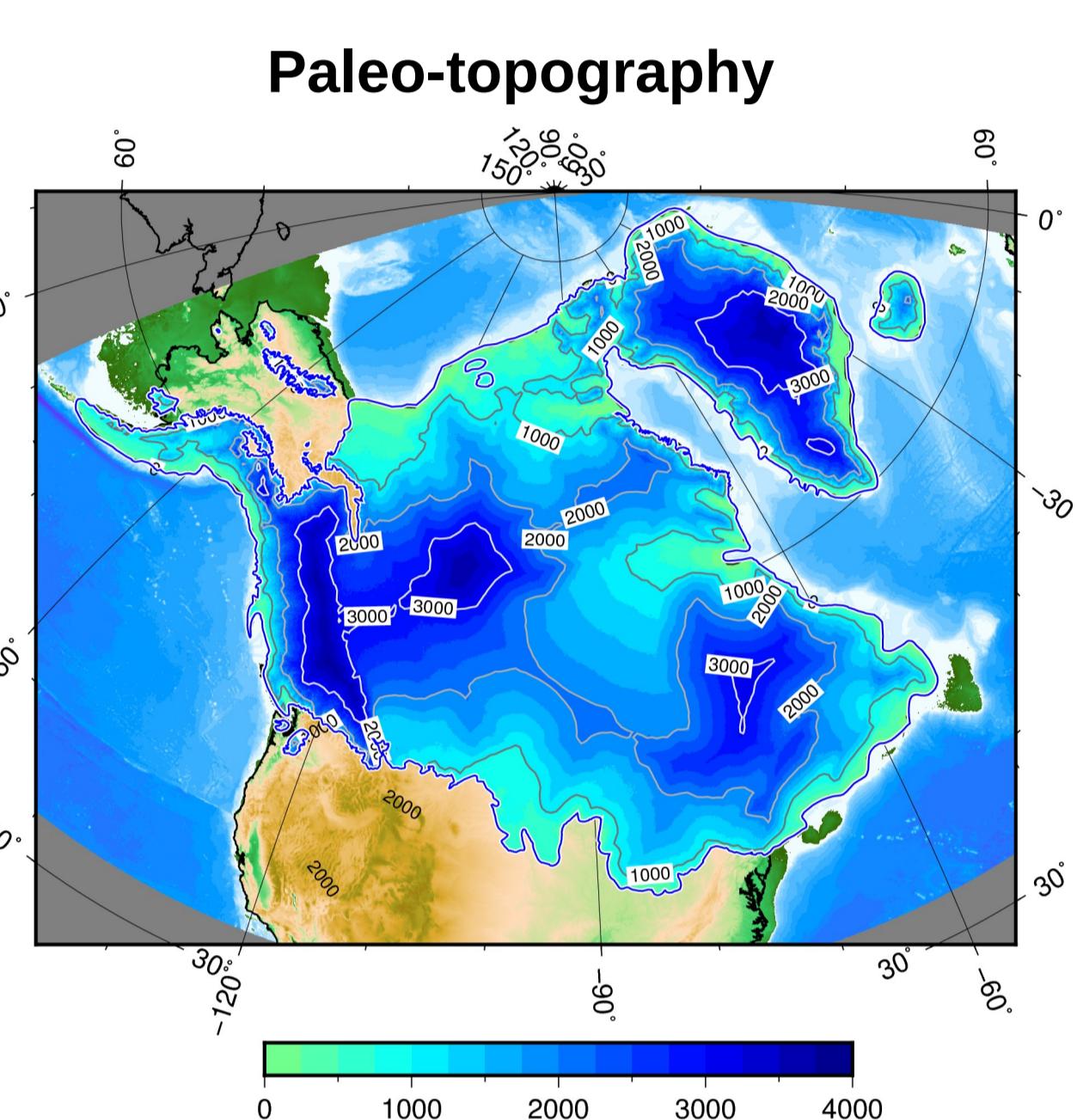
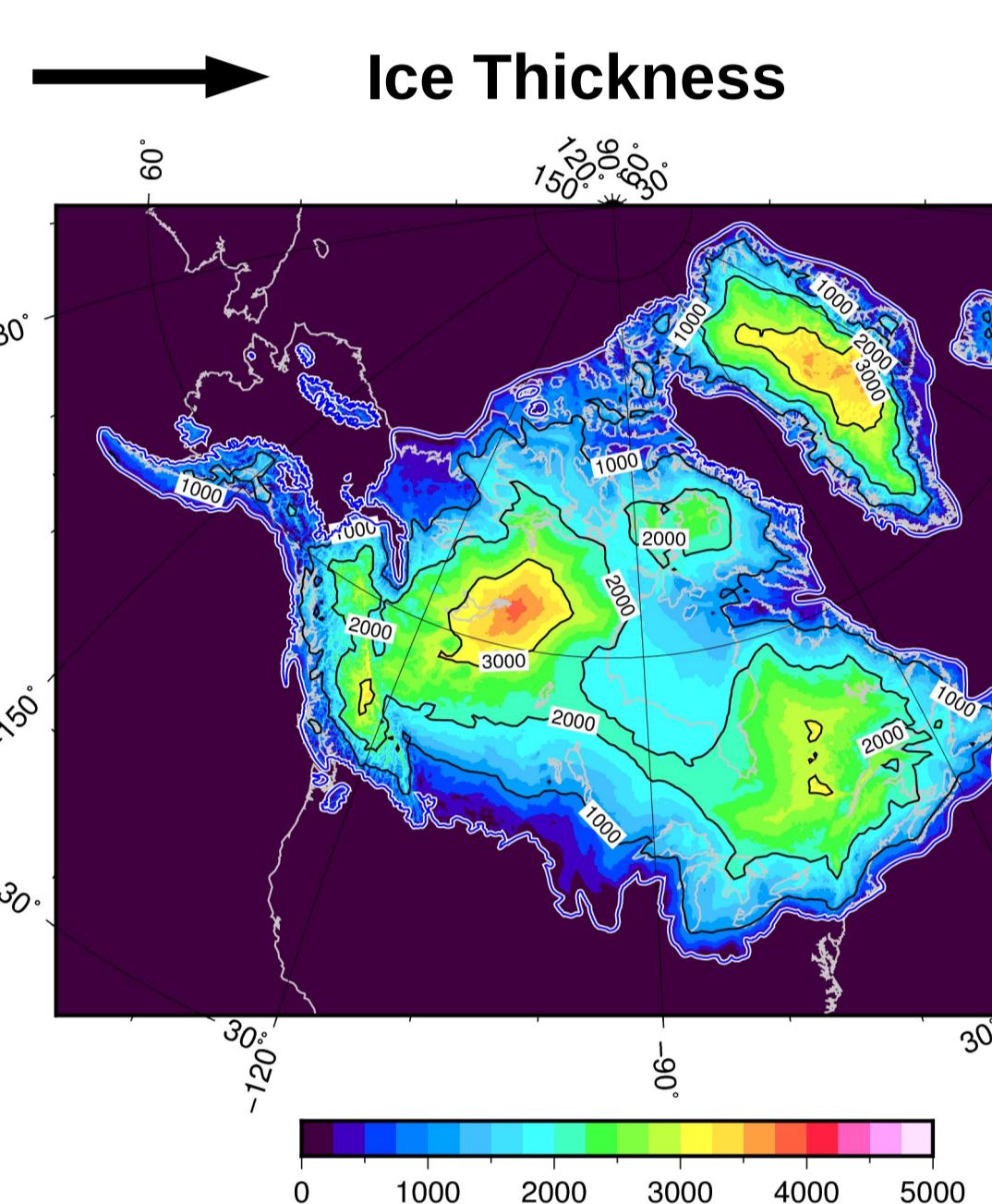
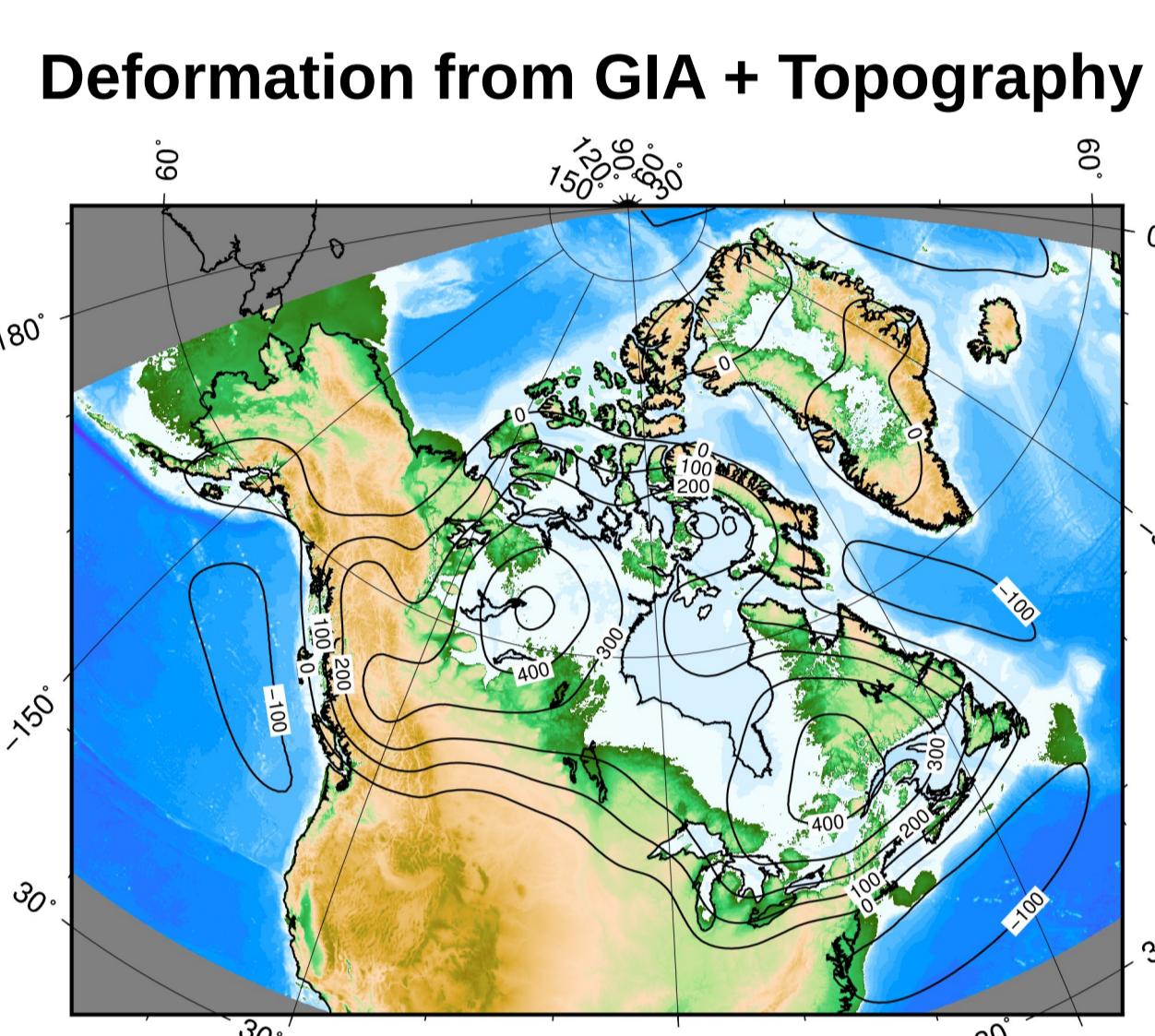
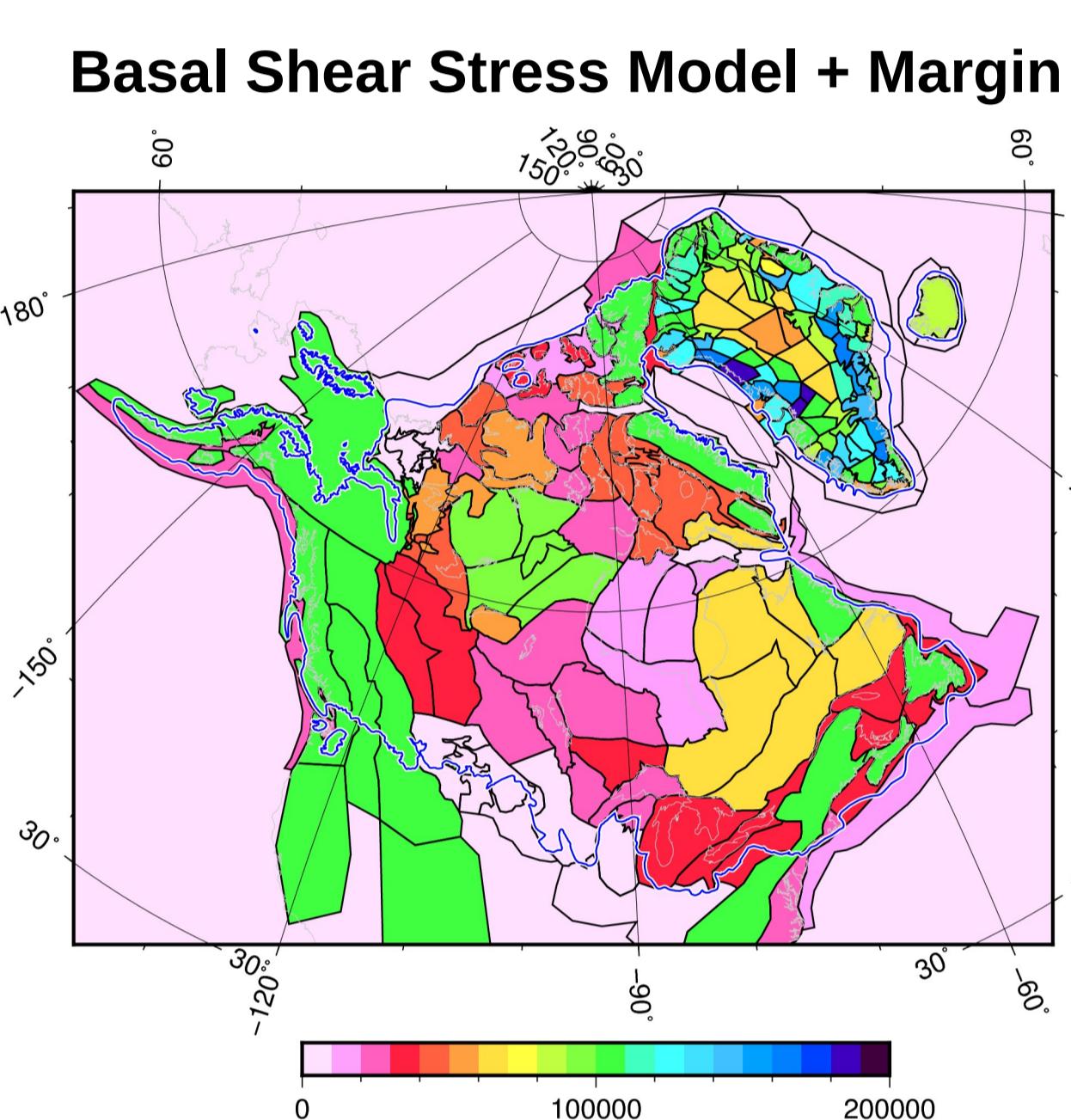
End moraines from the Erie Lobe of the Laurentide Ice Sheet

Methodology to make ice sheet reconstructions using ICESHEET

- Inputs for ICESHEET include the margin at discrete time periods, and a temporal variable basal shear stress model which controls the ice surface profile.
- Can include iterations of GIA to account for changes in basal topography from loading and sea level change. We use SELEN (Spada et al., 2012) to compute this.
- At present, we have setups for North American and Eurasian ice sheets.

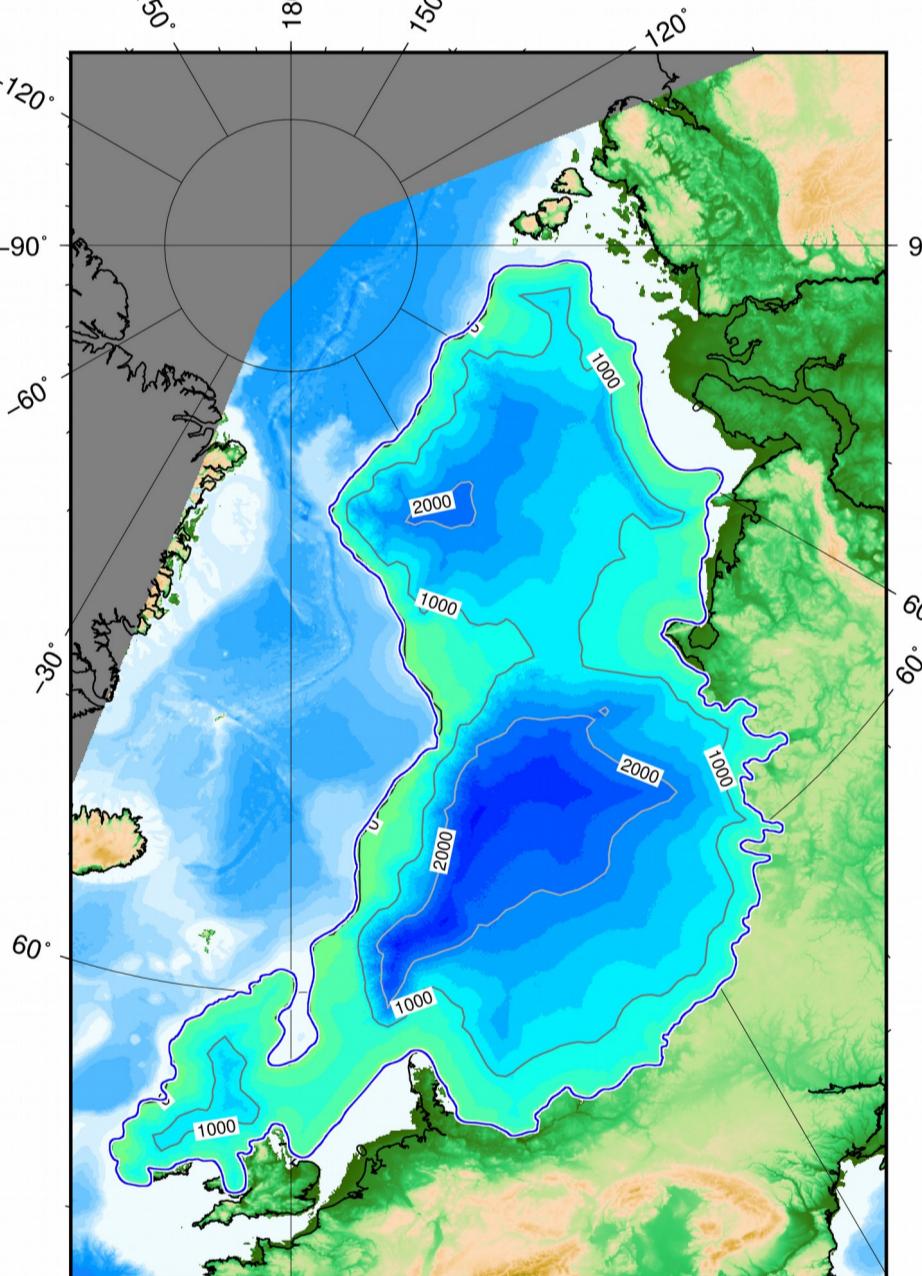
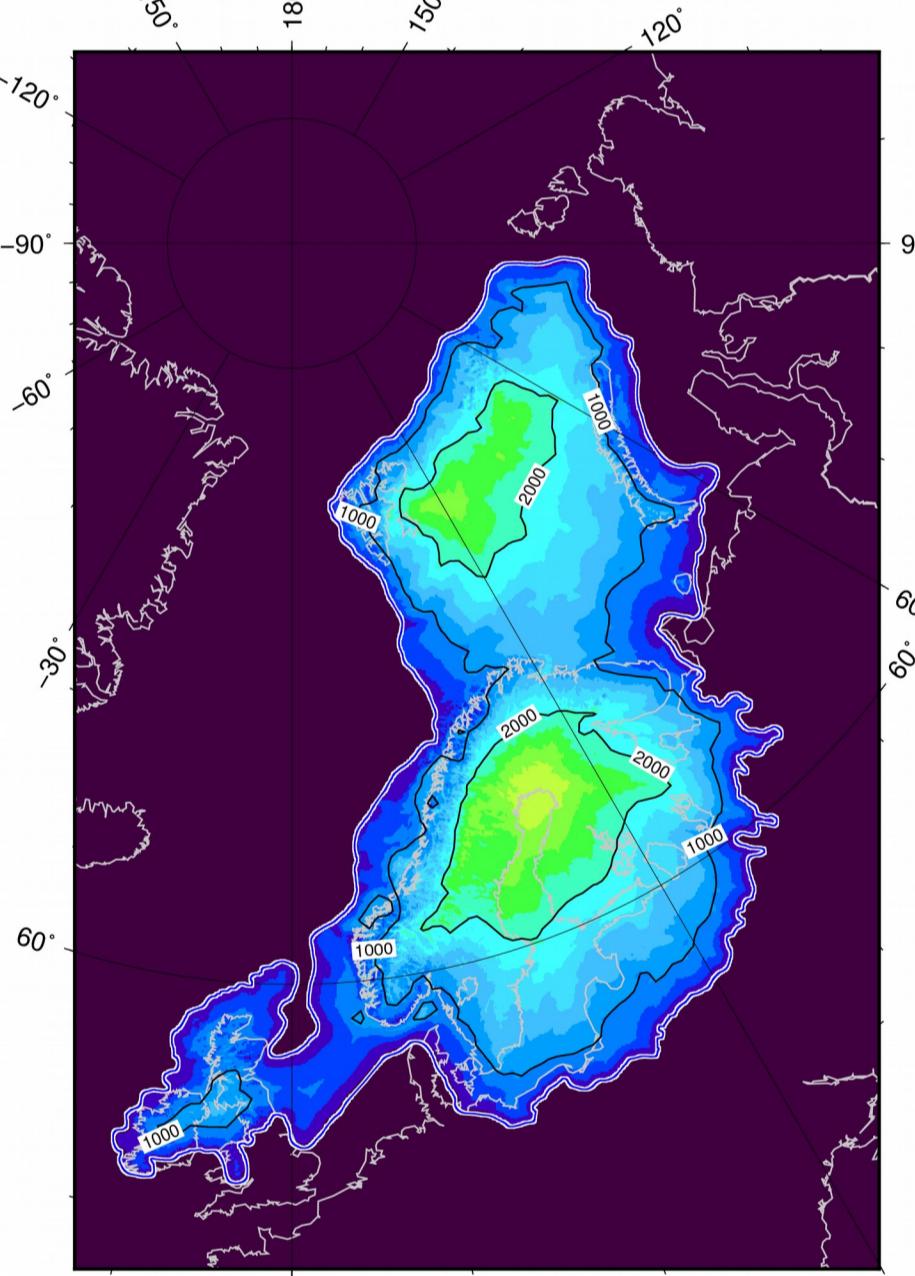
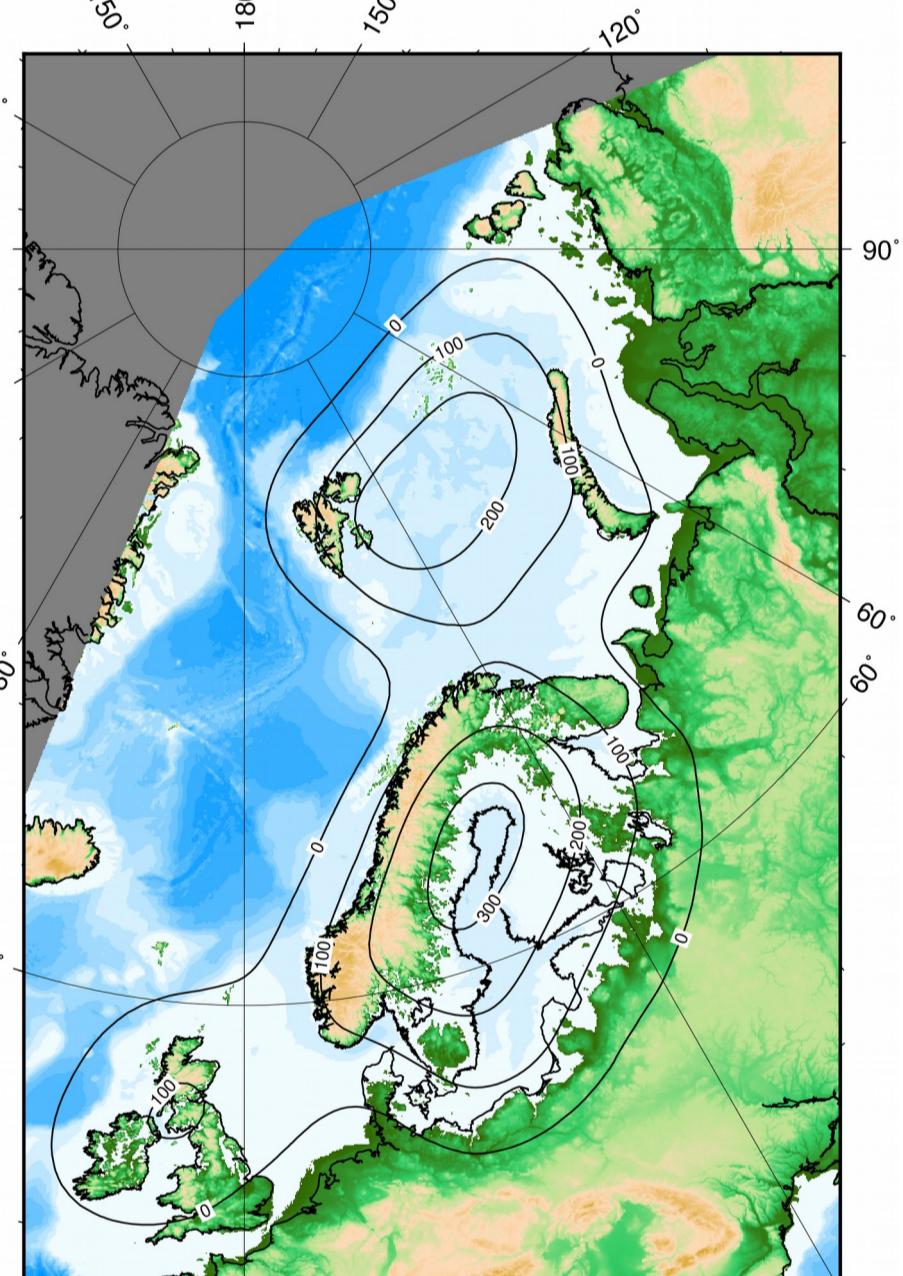
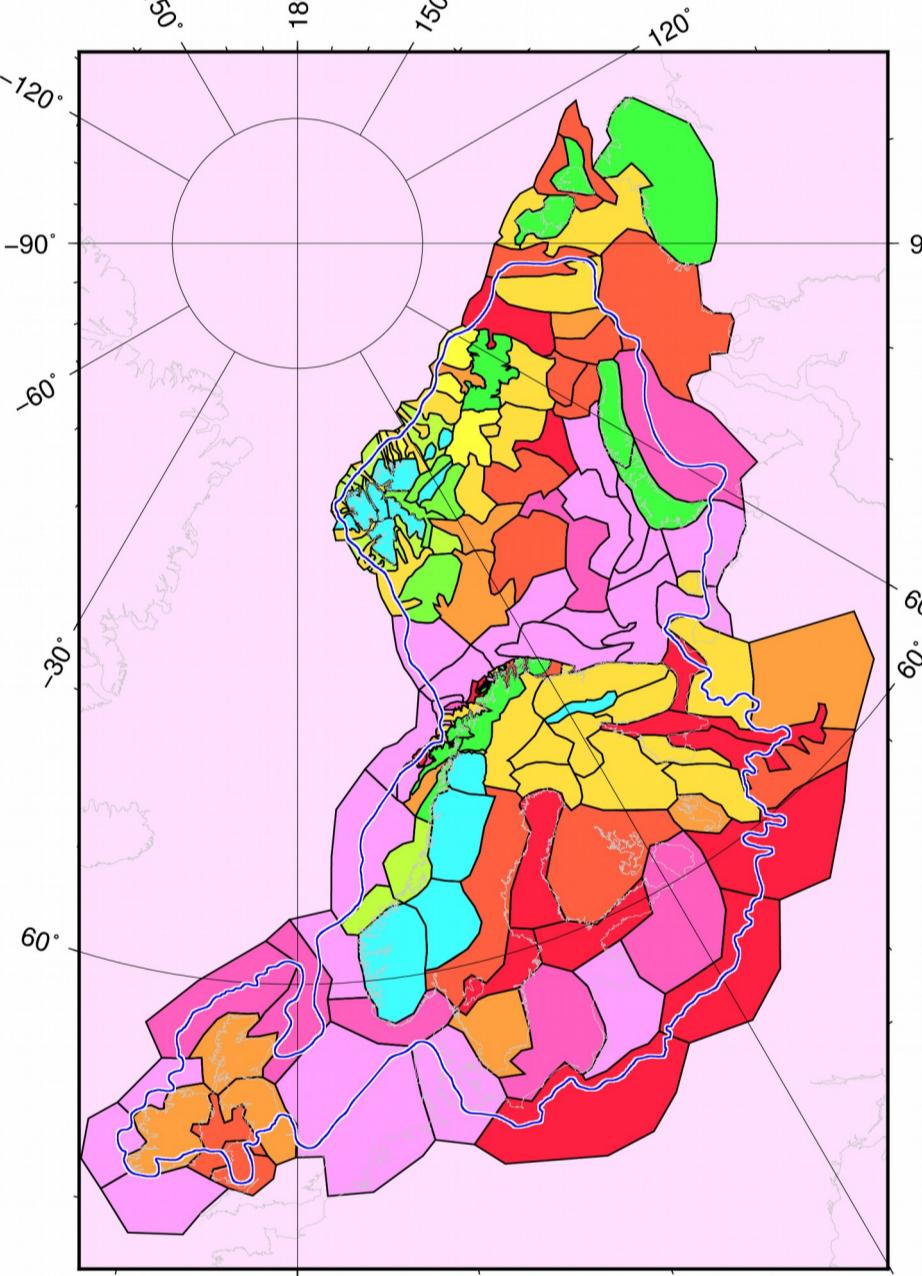
North American Ice sheets at 20000 yr BP

(blue line is the margin reconstruction from Dyke, 2004 and Gowan et al. 2016a)

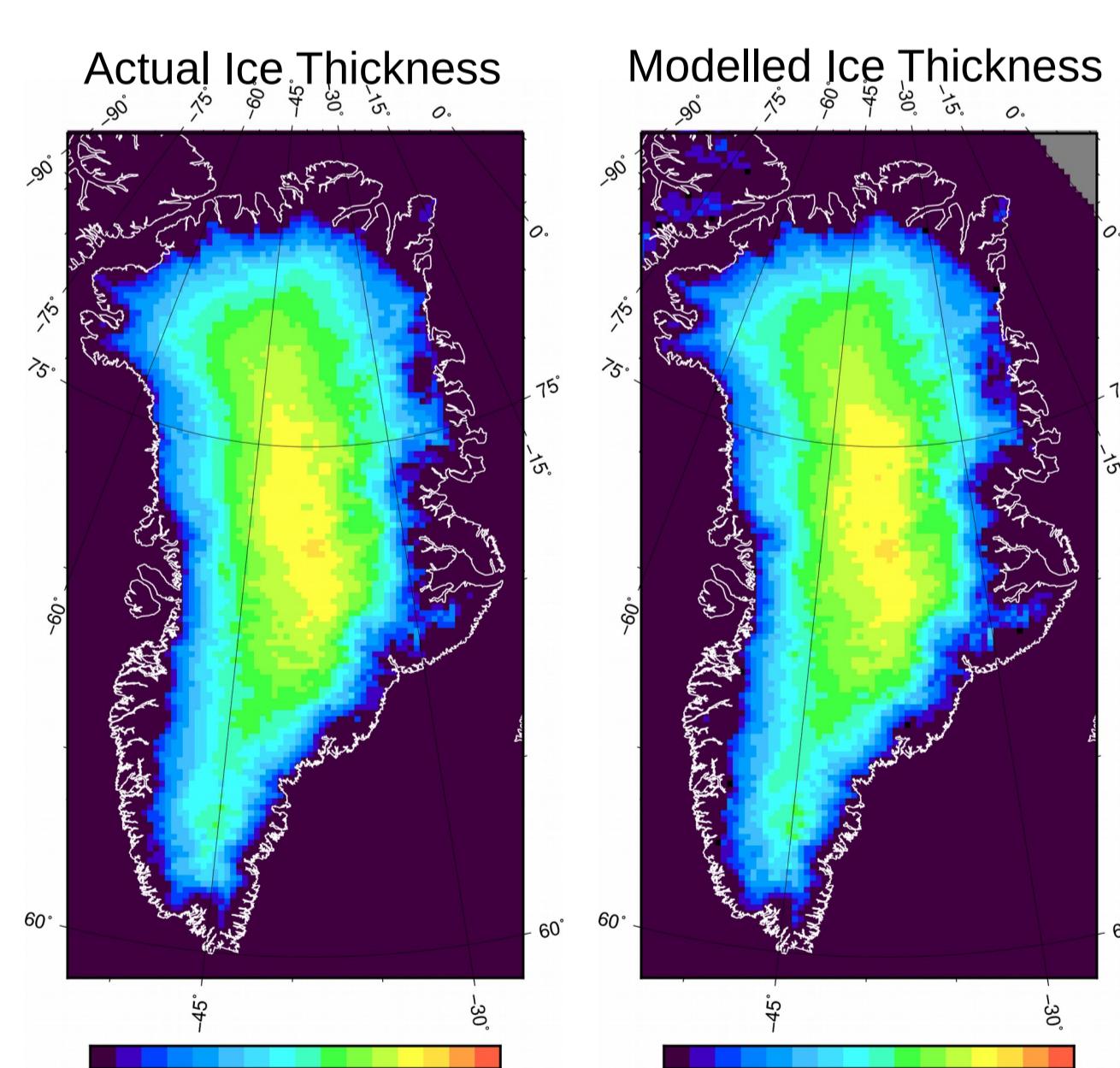


Eurasian Ice sheets at 20000 yr BP

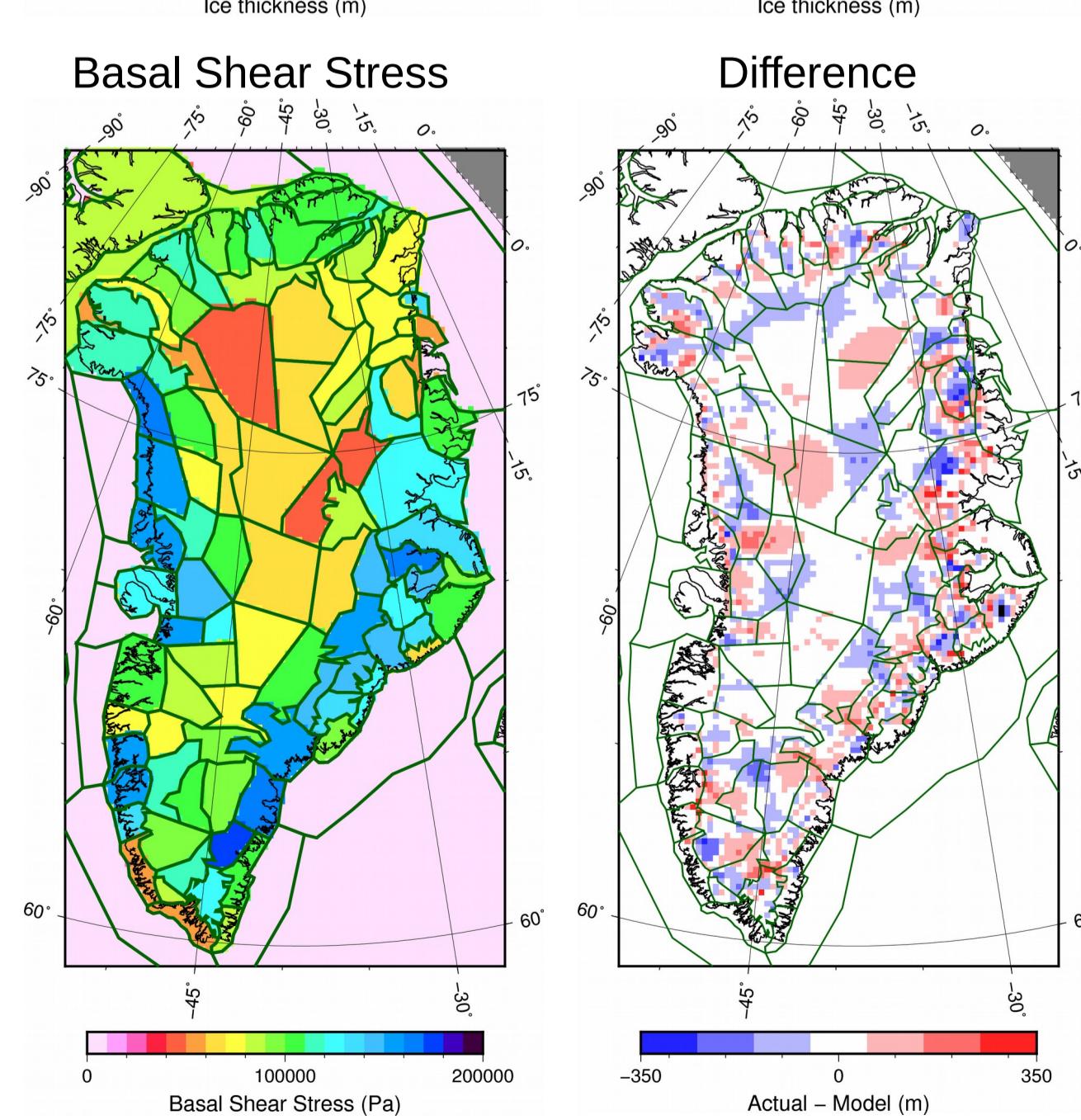
(blue line is the margin reconstruction from Hughes et al. 2016)



Greenland Ice Sheet



To test the utility of ICESHEET, it is instructive to show the results versus the contemporary Greenland Ice Sheet. The shear stress domains were adjusted to minimize the misfit of the modelled ice thickness and actual ice thickness. Even with the coarse resolution of the shear stress domains, the modelled ice thickness is generally within 250 m of the true thickness. The largest differences happen at the borders between the shear stress domains. Using coarser shear stress domains is advantages for paleo-reconstruction to reduce the amount of adjustable parameters.



Refining the ice sheet reconstruction

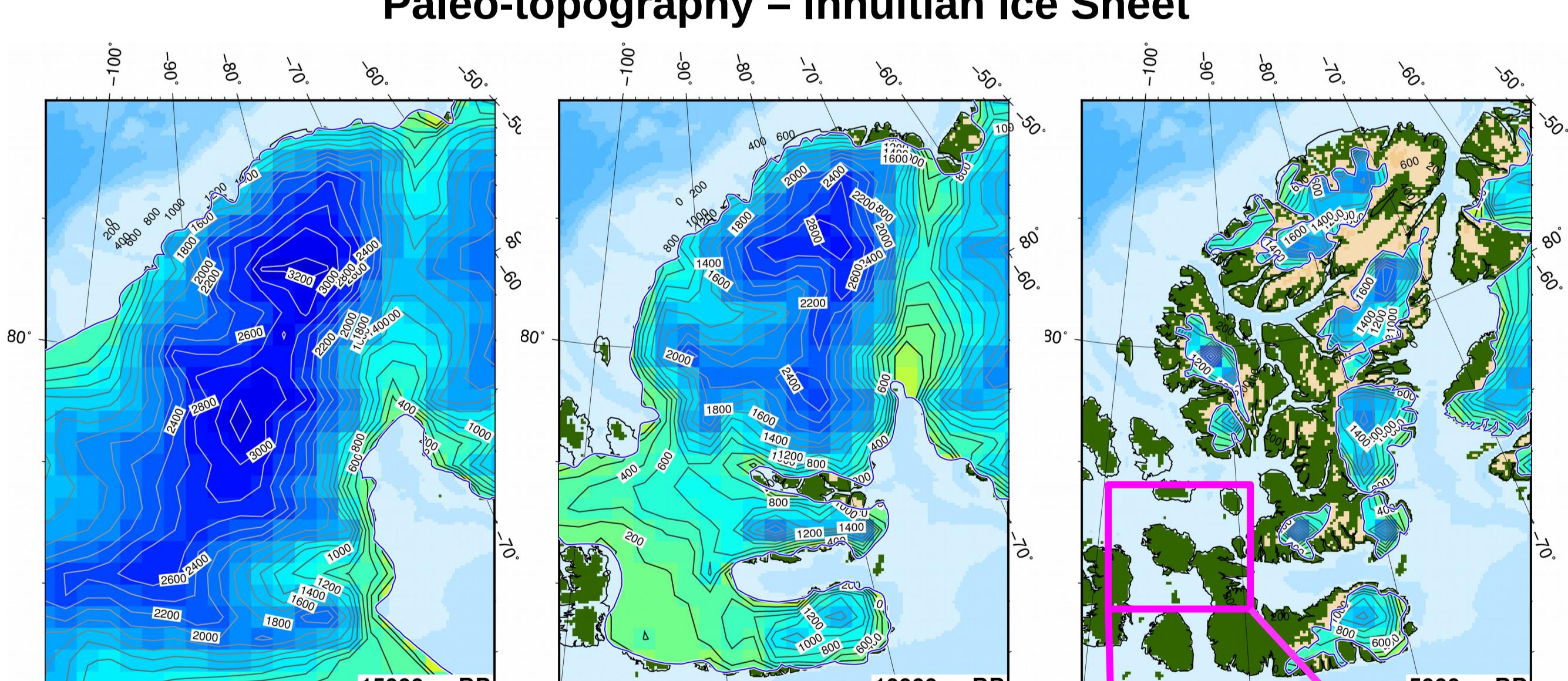
Paleo-topography – Innuitian Ice Sheet

Currently, we are refining the ice sheet reconstruction for the Innuitian Ice Sheet in Northern Canada. We are revising margins and sea level indicators using updated reservoir corrected radiocarbon dates (Butzin et al 2017).

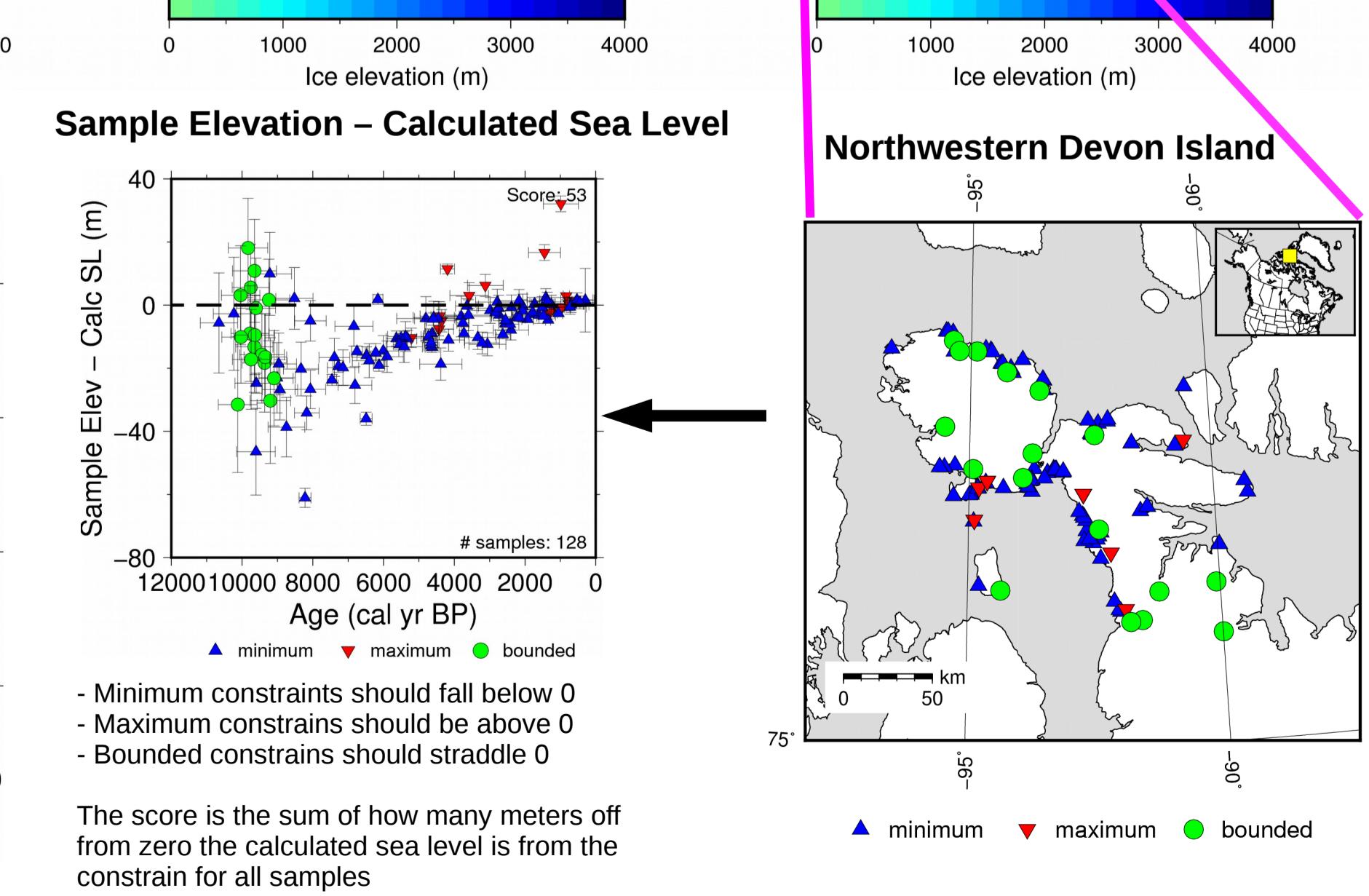
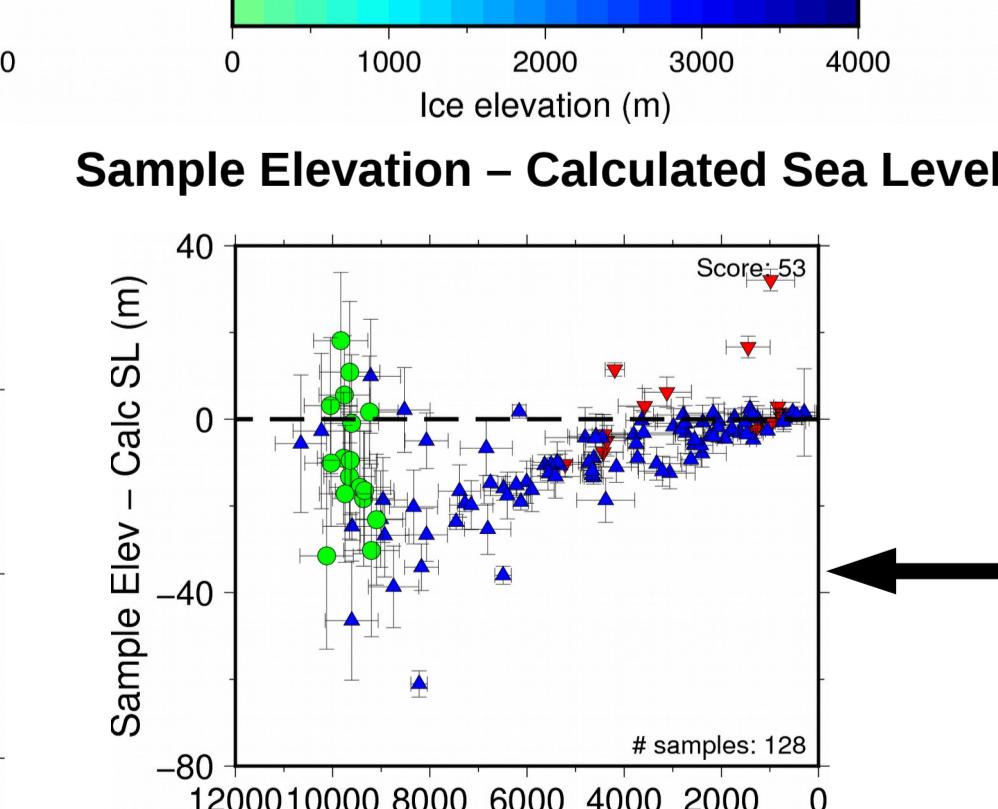
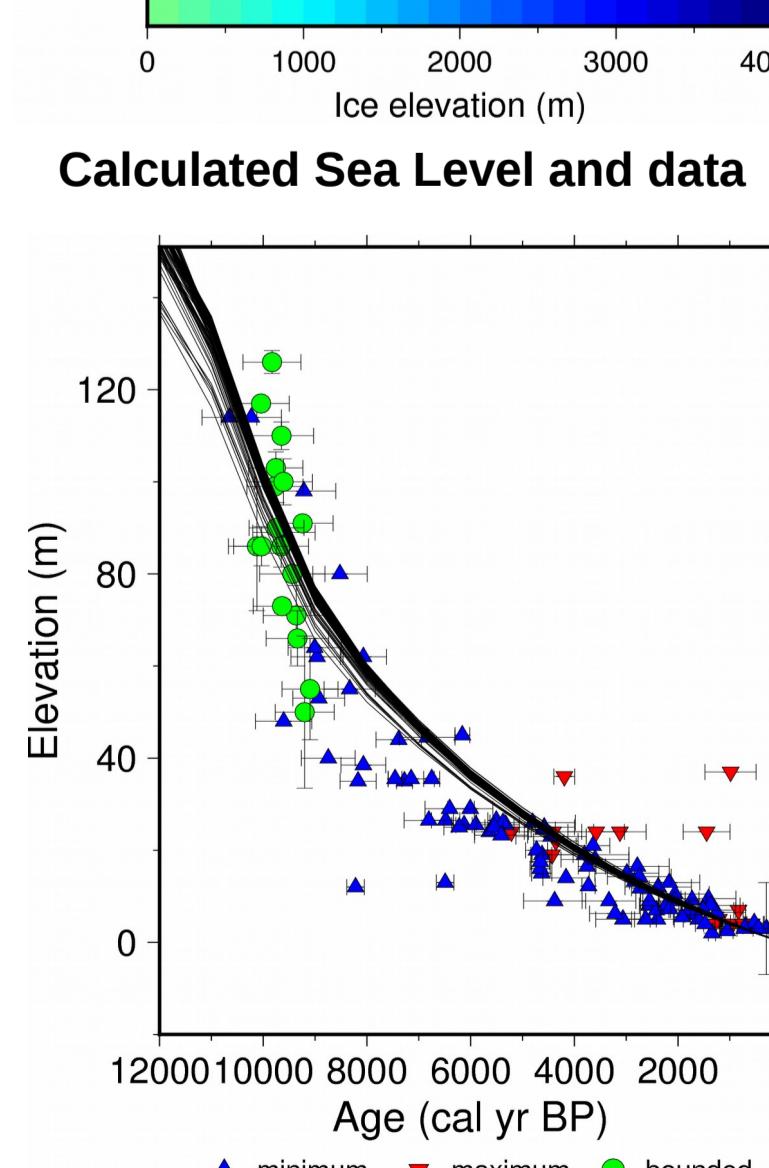
The sea level data are classified based on whether they indicate that sea level was above (minimum) or below (maximum) the sample elevation, or intermediate of the sample and the local highstand position (bounded).

Sea level is calculated at the location of each sample, and a score is assigned based on the discrepancy between the observation and model (zero if there is no discrepancy). This score is used to assess the ice sheet reconstruction.

The basal shear stress or margin models are adjusted if there is a discrepancy in calculated sea level.



Calculated Sea Level and data



The score is the sum of how many meters off from zero the calculated sea level is from the constrain for all samples

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Acknowledgements

This work was funded by the Helmholtz Climate Initiative REKLIM (Regional Climate Change), a joint research project of the Helmholtz Association of German research centres (HGF). This study was also supported by the PACES-II programme at AWI and the BMBF-funded project PalMod. We thank Art Dyke for making his sea level indicator database available.

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Ice sheet reconstructions based on geological and geophysical information

- Geophysical modeling of glacial isostatic adjustment (GIA) processes has long been used to reconstruct paleo-ice sheets (e.g. Tarasov et al. 2002, Peltier et al. 2015, Gowar et al. 2020a, Lohmann et al. 2020). The main advantage of this approach is that it does not rely on the geometry of the ice sheet.
- However, the GIA model is only able to predict the glacial distribution if a sea-level indicator is only located in coastal regions; glacial lake standstills only provide paleo-lake basins, and therefore do not provide a sea-level indicator.
- Ultimately, the reconstruction should have at least a minimal amount of glaciological realism. This can be achieved using our model, ICESHEET (Gowar et al. 2020b), which uses geological, glaciological and geophysical data.

Northeast of Lake Agassiz
West margin of the Laurentide Ice Sheet

Methodology to make ice sheet reconstructions using ICESHEET

Inputs for ICESHEET include the margin or discrete step margin, the paleogeography, and a basal shear stress model which controls the ice surface profile.

- Can include iterations of GIA to account for changes in the paleogeography and ice thickness distribution. We use SELEN (Spada et al., 2020) to compute this.
- At present, we have set-ups for North American and Eurasian ice sheets.

North American Ice sheets at 20000 yr BP
(blue line is the margin reconstruction from Dyke, 2018; Dyke et al., 2018a)

Eurasian ice sheets at 20000 yr BP
(dark blue line is the margin reconstruction from Hughes et al., 2016)

Greenland Ice Sheet

Actual vs Thickness
Modelled vs Thickness
Basal Shear Stress
Difference

To show the utility of ICESHEET, it is instructive to show the result for the contemporary Greenland ice sheet. The ice sheet domains were adjusted to match the observed ice thickness and basal shear stress. The modelled ice thickness is generally within 25% of the observed ice thickness. The main differences happen at the ice sheet margins. Changing these domain boundaries and adjusting the reconstruction to reduce the difference is a possible parameter.

Refining the ice sheet reconstruction

Paleo-topography - Innuitian Ice Sheet
Paleo-topography - Innuitian Ice Sheet
Calculated Sea Level and data
Sea level is calculated at the location of each sample, and a linear regression is used to remove the discrepancy between the calculated sea level and the sea level above (minimum) the sample. The sample is then projected onto the sea level of the sample and the location of the sample is used to calculate the basal shear stress.

The basal shear stress or margin models are adjusted if there is a discrepancy. This process is repeated until the calculated sea level matches the observed sea level.

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