

Department of Meteorology, University of Reading, UK, October 29, 2018

Ensemble Data Assimilation

Algorithms – Software – Applications

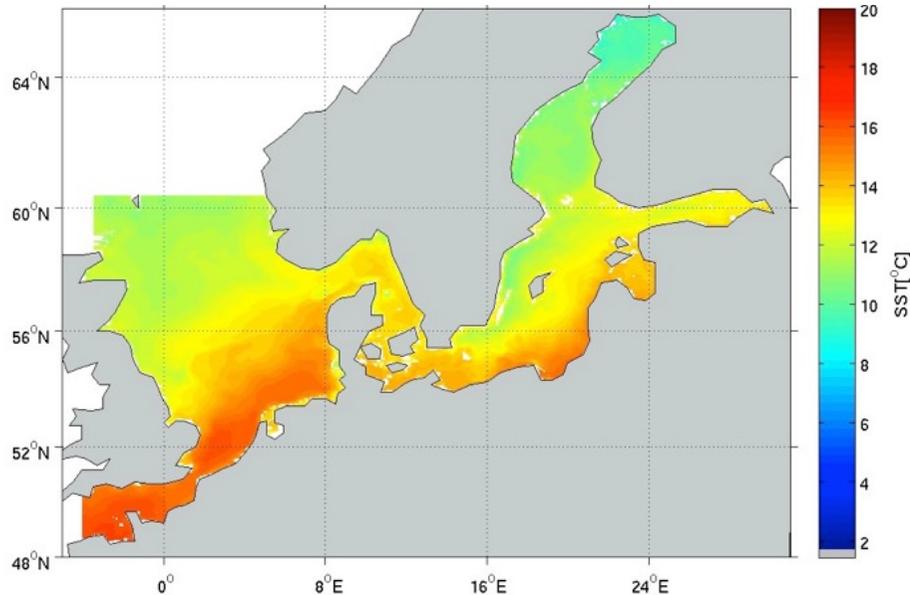
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Bremerhaven, Germany

Acknowledgements:

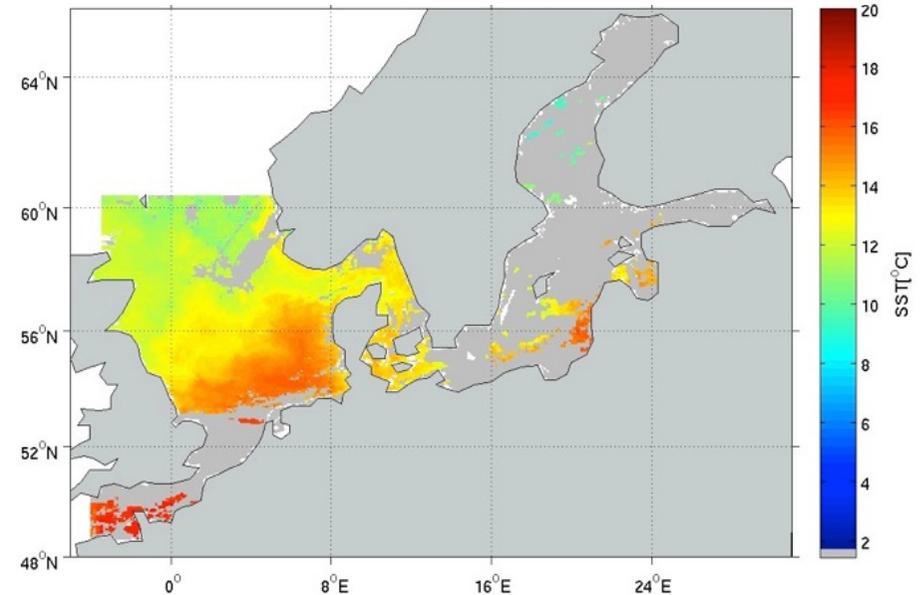
W. Hiller, J. Schröter, S. Losa, A. Androsov,
H. Pradhan, M. Goodliff, Q. Tang, Q. Yang, L. Mu

Model surface temperature



Information: Model

Satellite surface temperature



Information: Observations

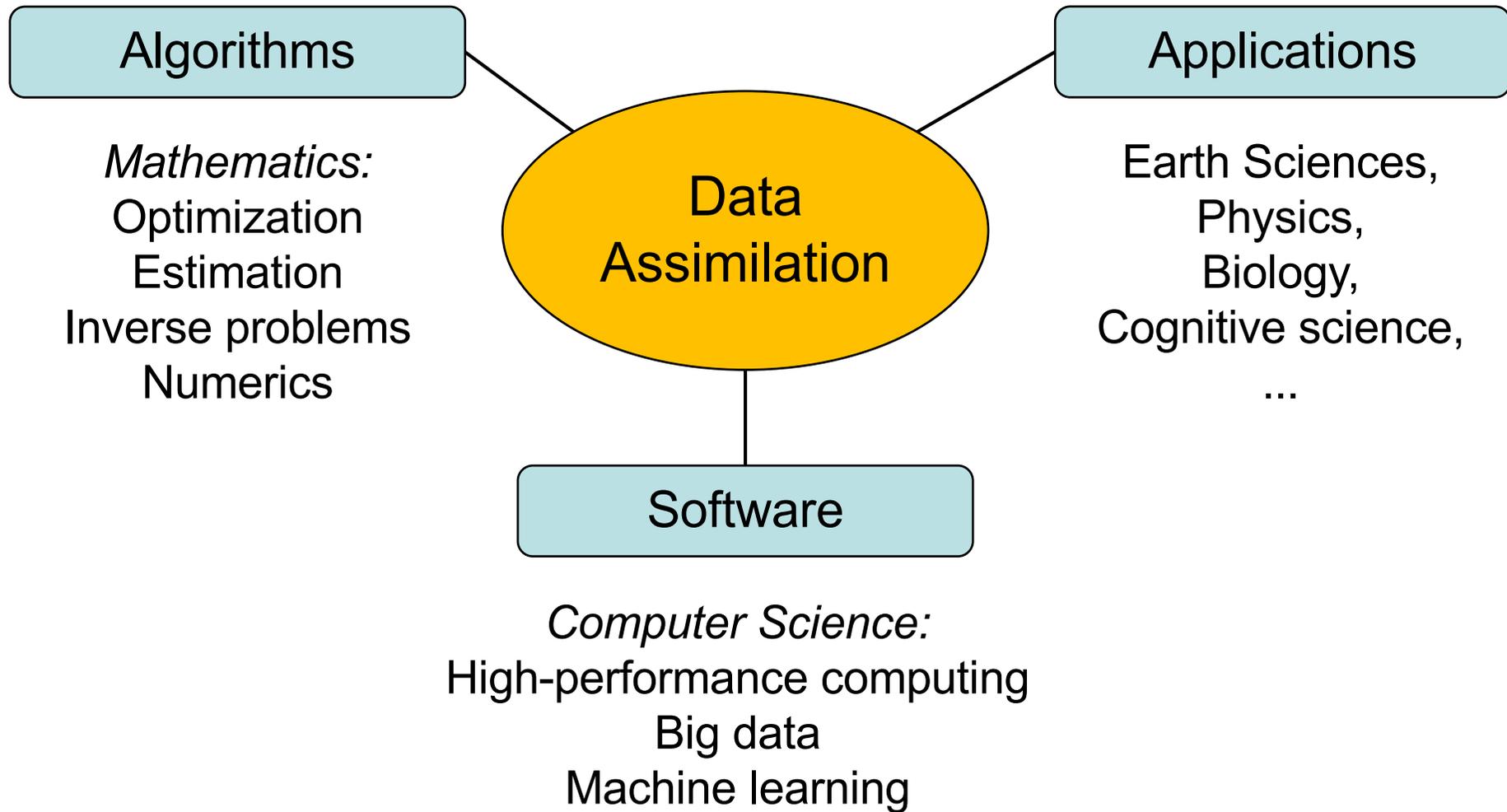
Combine both sources of information
quantitatively by computer algorithm
→ Data Assimilation

Data Assimilation

Methodology to combine model with real data

- Optimal estimation of system state:
 - initial conditions (for weather/ocean forecasts, ...)
 - state trajectory (temperature, concentrations, ...)
 - parameters (ice strength, plankton growth, ...)
 - fluxes (heat, primary production, ...)
 - boundary conditions and ‘forcing’ (wind stress, ...)
- More advanced: Improvement of model formulation
 - Detect systematic errors (bias)
 - Revise parameterizations based on parameter estimates

Interdisciplinarity of Data Assimilation



Outline

Ensemble Data Assimilation

Algorithms

- Understand behavior of different existing methods
- Develop efficient methods for high-dimensional nonlinear systems

Software

- Make data assimilation easily usable

Applications

- Assess assimilation into realistic model configurations
- Develop methodology for new modeling applications and data types

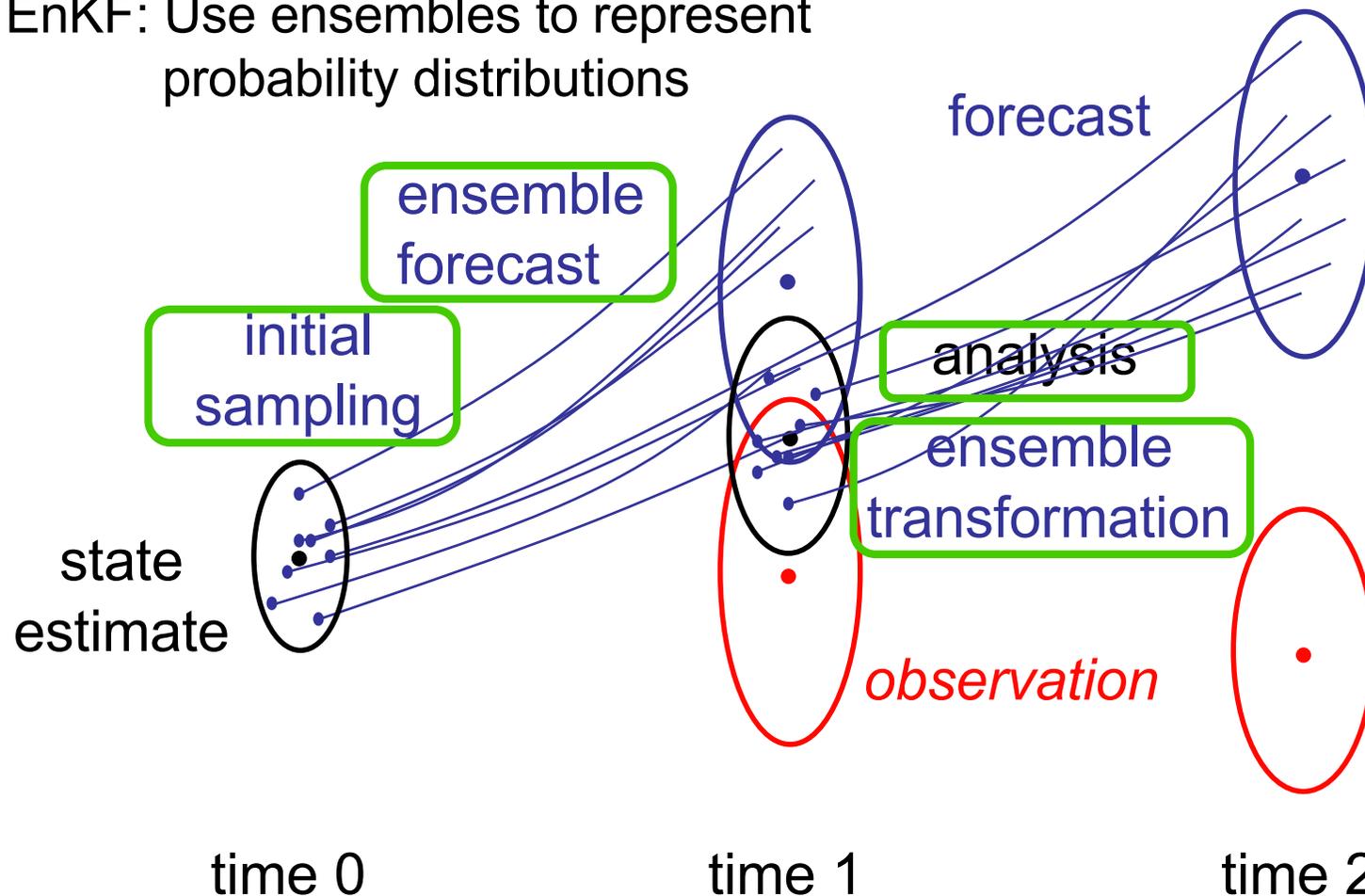
Algorithms

Ensemble-based Filtering

Original EnKF by G. Evensen (J. Geophys. Res. 1994)

Kalman filter: express probability distributions by mean and covariance matrix

EnKF: Use ensembles to represent probability distributions



There are many possible choices!

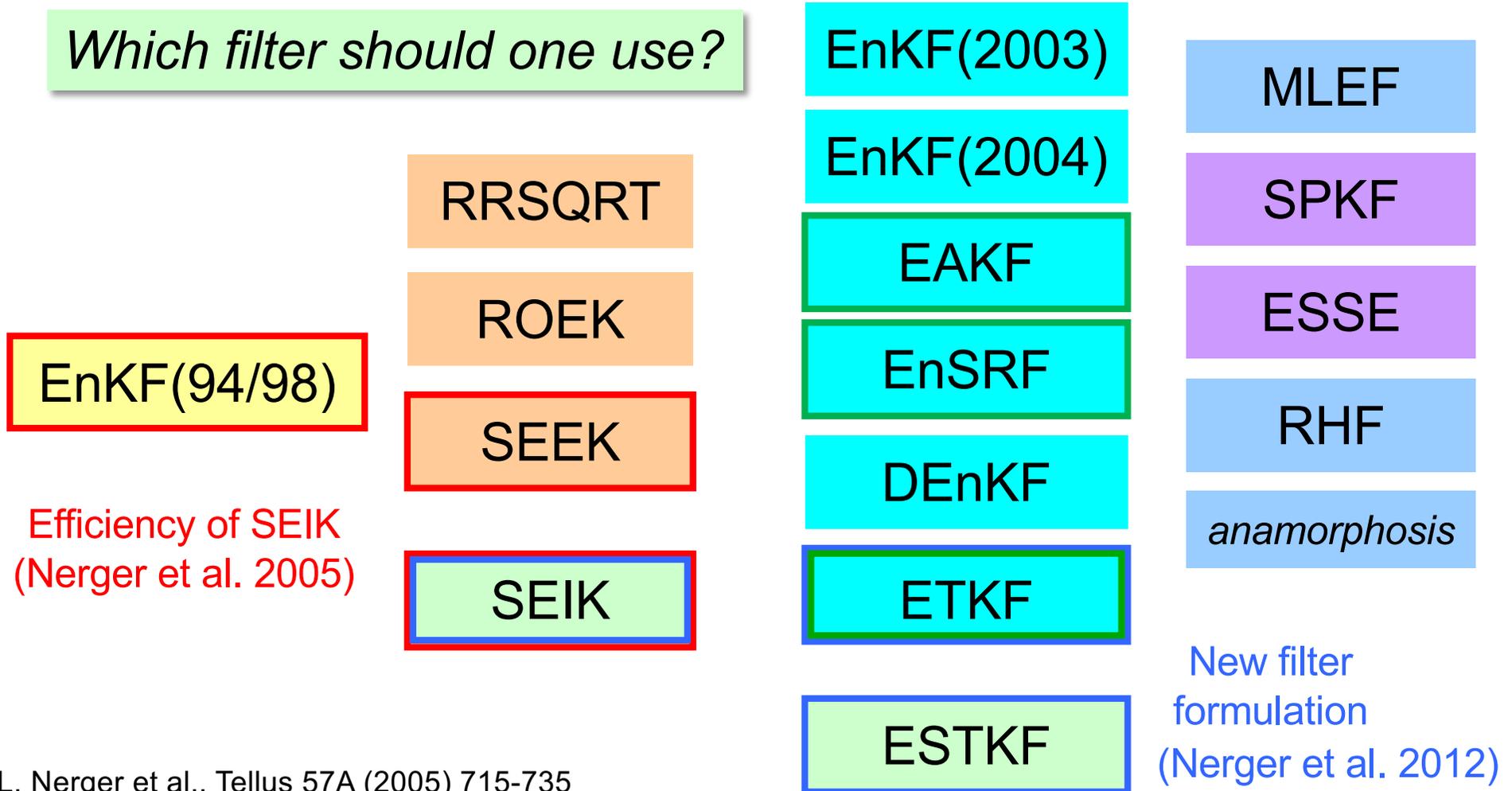
What is optimal is a research topic

Ensemble-based/error-subspace Kalman filters

A little “zoo” (not complete):

Which filter should one use?

Filter instability
(Nerger 2015)



L. Nerger et al., Tellus 57A (2005) 715-735

L. Nerger et al., Monthly Weather Review 140 (2012) 2335-2345

L. Nerger, Monthly Weather Review 143 (2015) 1554-1567

S. Vetra-Carvalho et al., Tellus A 70 (2018) 1445364



Assessing Ensemble Kalman Filters

Mathematical assessment of ensemble Kalman filters limited by

- optimality only proven for Gaussian error distributions
- convergence properties only clear for large ensemble limit

but

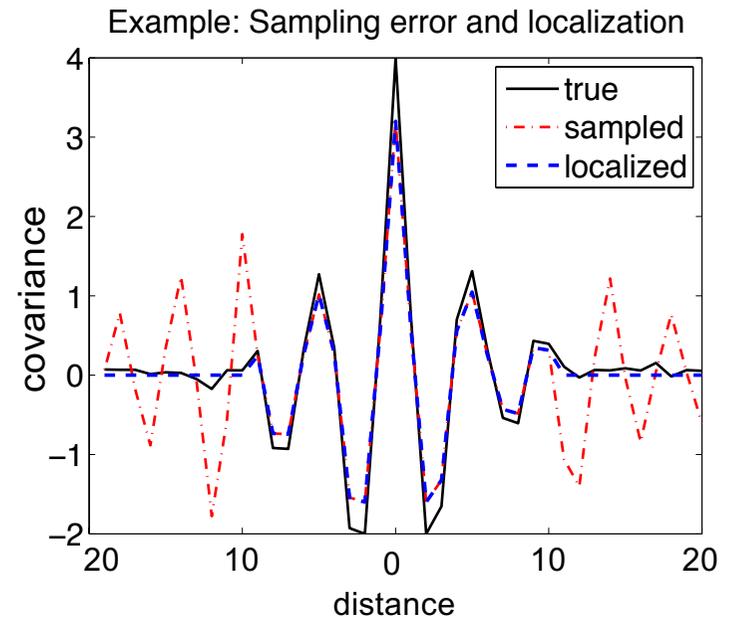
- models are nonlinear -> non-Gaussian distributions
- only small ensemble feasible to run for high-dimensional models

My approach

- compare and characterize behavior of different methods
- reach general conclusions from analyzing differences mathematically

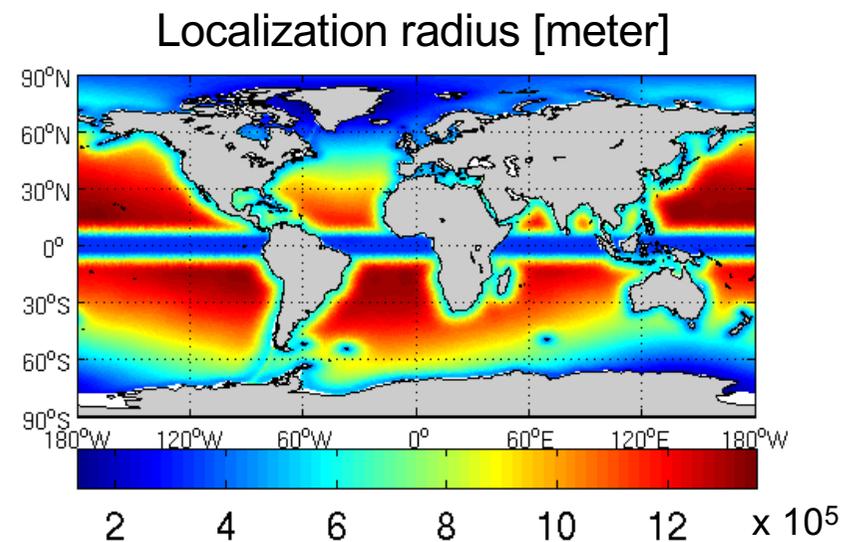
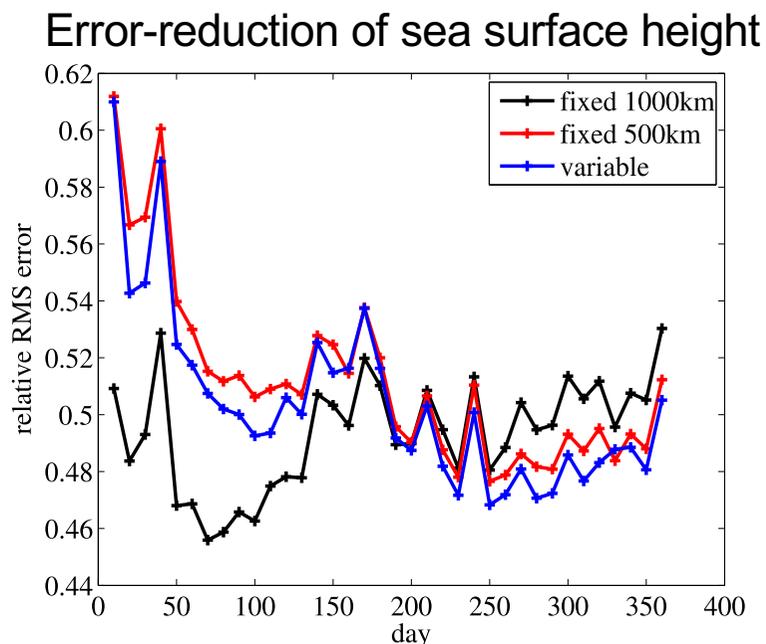
Localization: Why and how?

- Combination of observations and model state based on ensemble estimates of error covariance matrices
- Finite ensemble size leads to significant sampling errors
 - errors in variance estimates
 - usually too small
 - errors in correlation estimates
 - wrong size if correlation exists
 - spurious correlations when true correlation is zero
- Assume: long-distance correlations are small in reality
- Localization: damp or remove estimated long-range correlations (Houtekamer & Mitchell, 1998, 2001)



Adaptive localization radius in global ocean model

- Localization radius is usually hand-tuned
- Numerical analysis in small models shows:
errors minimal when localization radius chosen such that
local sum of observation weights = ensemble size
- Application with FESOM (Finite Element Sea-ice Ocean Model):
 - Fixed 1000km radius leads to increasing errors in 2nd half of year
 - Lower RMS error in sea surface height than fixed 500km radius



Instability of serial observation processing

Two widely used filter categories:

Serial observation processing

EnSRF, EAKF

- Perform a loop assimilating each single observation
- Efficient: Avoids matrix-matrix operations
- Requires diagonal observation error covariance matrix
- Localization of state error covariance matrix

Synchronous assimilation

ETKF, SEIK, ESTKF, (EnKF)

- Assimilation all observation at a given time at once
- Usually using ensemble-space transformations
- Possible for arbitrary observation error covariance matrices
- Localization of observation error covariance matrix

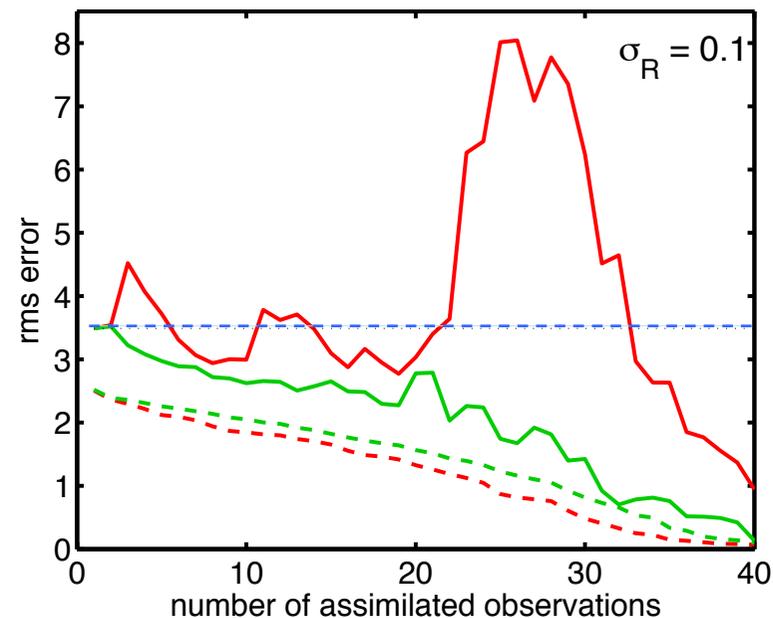
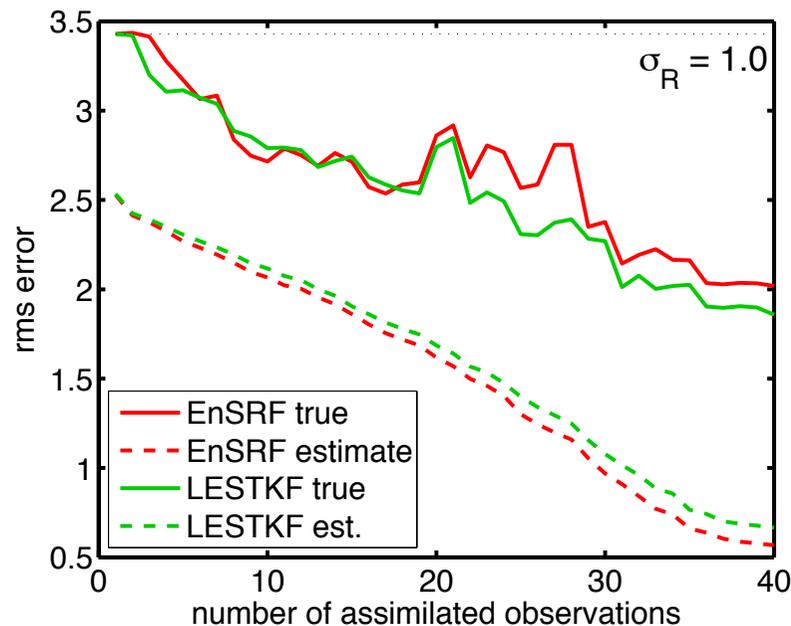
(EnSRF: Whitaker & Hamill, 2002; EAKF: Anderson, 2001)

RMS error over number of observations

How does the RMS error develop during the loop over all observations?

Test at first analysis step (Lorenz-96 toy model):

- EnSRF: Compute RMS errors at each iteration
- LESTKF: Do 40 experiments with increasing number of obs.



- Instability leads to larger error for EnSRF in full-length experiments
- Can be relevant in real applications: if observations have locally strong impact



Inconsistent Matrix Updates

P State error covariance
R Obs. error covariance
H Observation operator

Kalman filter updates covariance matrix according to

$$\mathbf{P}^a = (\mathbf{I} - \mathbf{K}\mathbf{H}) \mathbf{P}^f (\mathbf{I} - \mathbf{K}\mathbf{H})^T + \mathbf{K}\mathbf{R}\mathbf{K}^T \quad (1)$$

With Kalman gain

$$\mathbf{K} = \mathbf{P}^f \mathbf{H}^T (\mathbf{H}\mathbf{P}^f \mathbf{H}^T + \mathbf{R})^{-1} \quad (2)$$

this simplifies to

$$\mathbf{P}^a = (\mathbf{I} - \mathbf{K}\mathbf{H}) \mathbf{P}^f \quad (3)$$

Be careful when
introducing new
adaptions!

(1) and (3) yield same result **only** with gain (2)!

Not fulfilled with localization:

$$\mathbf{K}_{loc} = (\mathbf{C} \circ \mathbf{P}^f) \mathbf{H}^T (\mathbf{H} (\mathbf{C} \circ \mathbf{P}^f) \mathbf{H}^T + \mathbf{R})^{-1}$$

- Update of **P** is inconsistent in localized EnSRF (noted by Whitaker & Hamill (2002), but never further examined)
- Inconsistency also occurs in localized synchronous assimilation ... but update is only done once followed by ensemble forecast

Linear and Nonlinear Ensemble Filters

- Represent state and its error by ensemble \mathbf{X} of N states
- Forecast:
 - Integrate ensemble with numerical model

- Analysis:

- update ensemble mean $\bar{\mathbf{x}}^a = \bar{\mathbf{x}}^f + \mathbf{X}'^f \tilde{\mathbf{w}}$
- update ensemble perturbations $\mathbf{X}'^a = \mathbf{X}'^f \mathbf{W}$

(both can be combined in a single step)

- Ensemble Kalman & nonlinear filters: Different definitions of
 - weight vector $\tilde{\mathbf{w}}$
 - Transform matrix \mathbf{W}

ETKF (Bishop et al., 2001)

- Ensemble Transform Kalman filter
 - Assume Gaussian distributions
 - Transform matrix

$$\mathbf{A}^{-1} = (N - 1)\mathbf{I} + (\mathbf{H}\mathbf{X}'^f)^T \mathbf{R}^{-1} \mathbf{H}\mathbf{X}'^f$$

- Mean update weight vector

$$\tilde{\mathbf{w}} = \mathbf{A}(\mathbf{H}\mathbf{X}'^f)^T \mathbf{R}^{-1} \left(\mathbf{y} - \mathbf{H}\overline{\mathbf{x}}^f \right)$$

(depends linearly on \mathbf{y})

- Transformation of ensemble perturbations

$$\mathbf{W} = \sqrt{(N - 1)} \mathbf{A}^{-1/2} \mathbf{\Lambda}$$

(depends only on \mathbf{R} , not \mathbf{y})

NETF (Tödter & Ahrens, 2015)

- Nonlinear Ensemble Transform Filter

- Mean update from Particle Filter weights: for all particles i

$$\tilde{w}^i \sim \exp \left(-0.5 (\mathbf{y} - \mathbf{H}\mathbf{x}_i^f)^T \mathbf{R}^{-1} (\mathbf{y} - \mathbf{H}\mathbf{x}_i^f) \right)$$

(Nonlinear function of observations \mathbf{y})

- Ensemble update

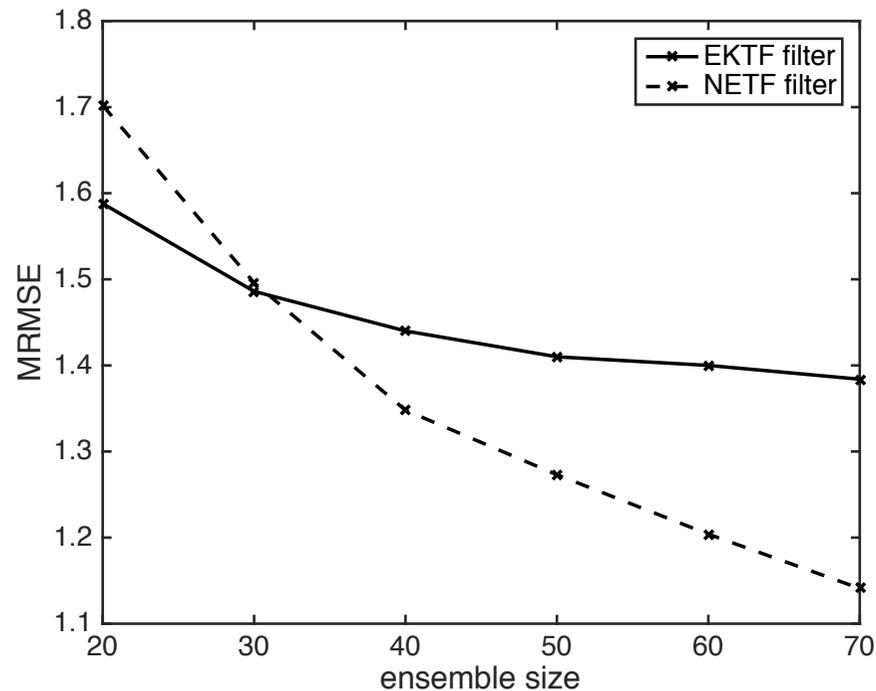
- Transform ensemble to fulfill analysis covariance (like ETKF, but not assuming Gaussianity)
- Derivation gives

$$\mathbf{W} = \sqrt{N} \left[\text{diag}(\tilde{\mathbf{w}}) - \tilde{\mathbf{w}}\tilde{\mathbf{w}}^T \right]^{1/2} \mathbf{\Lambda}$$

($\mathbf{\Lambda}$: mean-preserving random matrix; useful for stability)

Performance of NETF – Lorenz-96

- Double-exponential observation errors
- Run all experiments 10x with different initial ensemble



- NETF beats ETKF for ensemble size $N > 30$
- Larger ensemble needed for Gaussian errors

ETKF-NETF – Hybrid Filter Variants

1-step update (*HSync*)

$$\mathbf{X}_{HSync}^a = \bar{\mathbf{X}}^f + (1 - \gamma)\Delta\mathbf{X}_{NETF} + \gamma\Delta\mathbf{X}_{ETKF}$$

- $\Delta\mathbf{X}$: assimilation increment of a filter
- γ : hybrid weight (between 0 and 1; 1 for fully ETKF)

2-step updates

Variant 1 (*HNK*): NETF followed by ETKF

$$\tilde{\mathbf{X}}_{HNK}^a = \mathbf{X}_{NETF}^a[\mathbf{X}^f, (1 - \gamma)\mathbf{R}^{-1}]$$

$$\mathbf{X}_{HNK}^a = \mathbf{X}_{ETKF}^a[\tilde{\mathbf{X}}_{HNK}^a, \gamma\mathbf{R}^{-1}]$$

- Both steps computed with increased \mathbf{R} according to γ

Variant 2 (*HKN*): ETKF followed by NETF

Choosing hybrid weight γ

- Hybrid weight shifts filter behavior
- How to choose it?

Possibilities:

- Fixed value
- Adaptive
 - According to which condition?

- Base on effective sample size $N_{eff} = \sum_i 1/(w^i)^2$

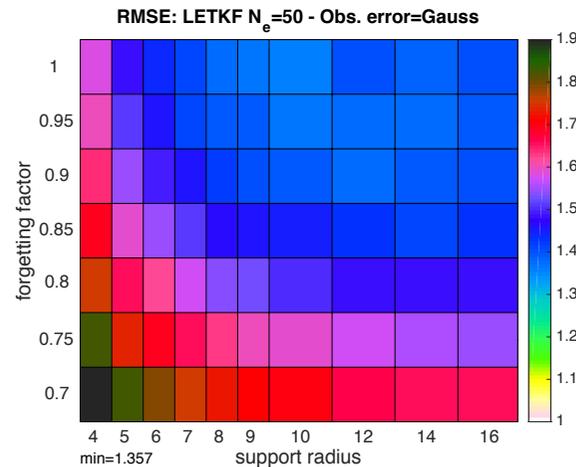
set

$$\gamma_{adap} = 1 - N_{eff}/N$$

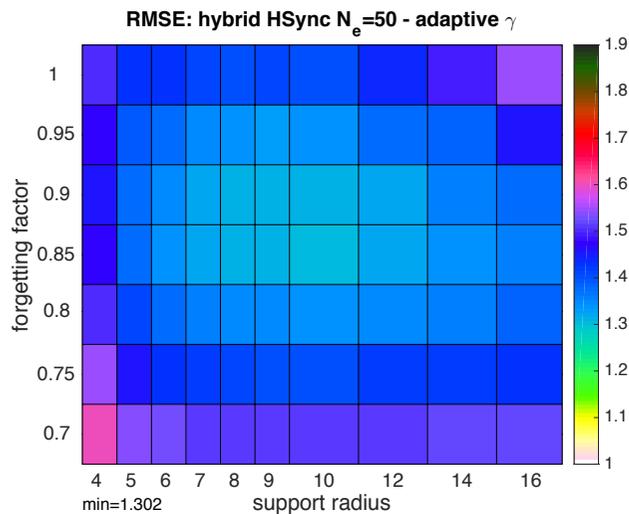
(close to 1 if N_{eff} small, i.e. small contribution of NETF)

Test with Lorenz-96 Model (ensemble size N=50)

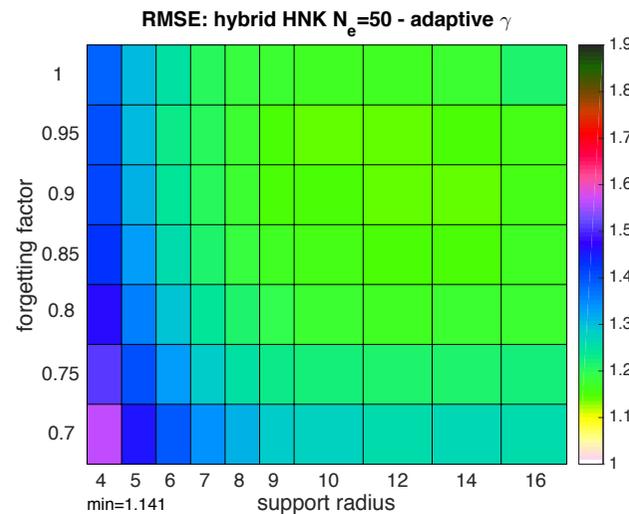
Ensemble size N=50



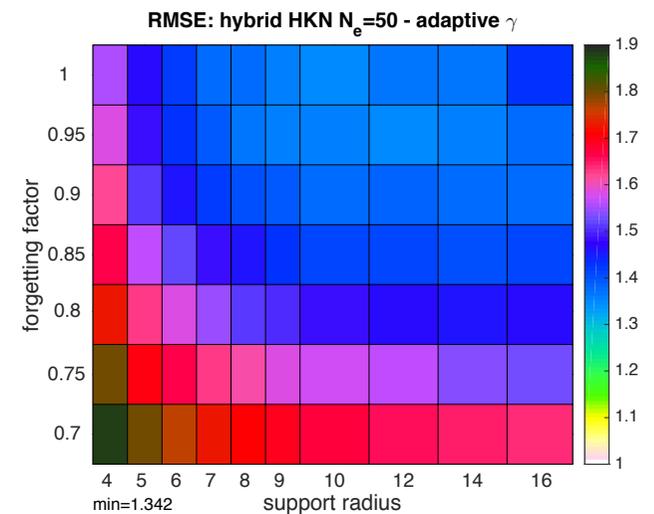
- All hybrid variants improve estimates compared to LETKF & NETF
- Dependence on forgetting factor & localization radius like LETKF
- Similar optimal localization radius
- Largest improvement for variant HNK (NETF before LETKF)
- Currently testing in a larger model ...



4% improvement



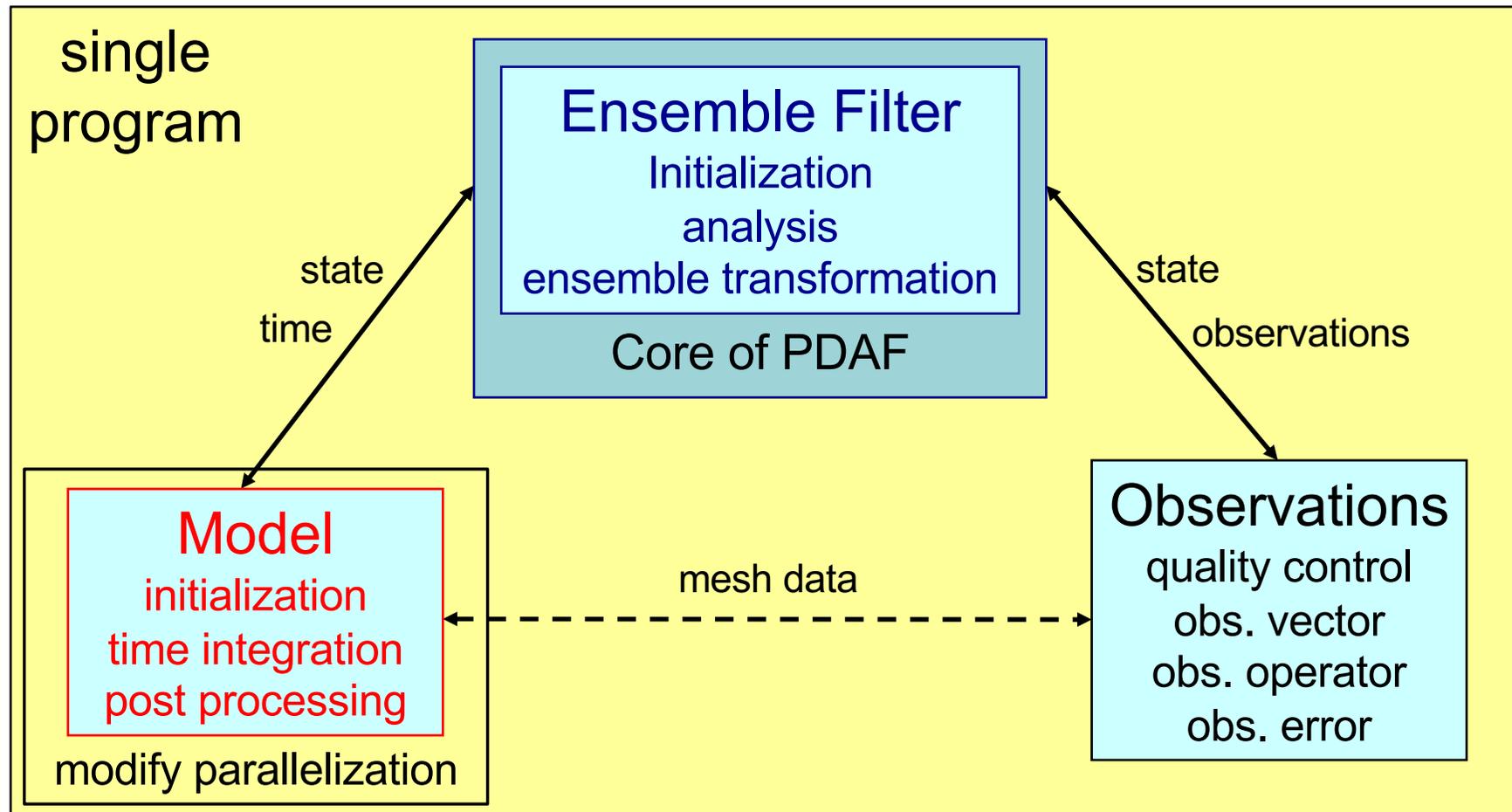
16% improvement



1% improvement

Software

Components of an Assimilation System



↔ Explicit interface

⋯ Indirect exchange (module/common)

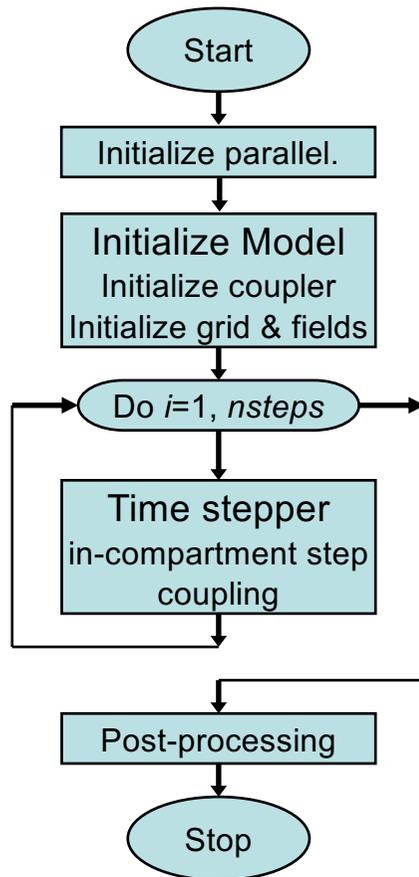
PDAF - Parallel Data Assimilation Framework

- a program library for ensemble data assimilation
- provide support for parallel ensemble forecasts
- provide fully-implemented & parallelized filters and smoothers (EnKF, LETKF, NETF, EWPF ... easy to add more)
- easily useable with (probably) any numerical model (applied with NEMO, MITgcm, FESOM, HBM, TerrSysMP, ...)
- run from laptops to supercomputers (Fortran, MPI & OpenMP)
- first public release in 2004; continuous further development
- ~310 registered users; community contributions

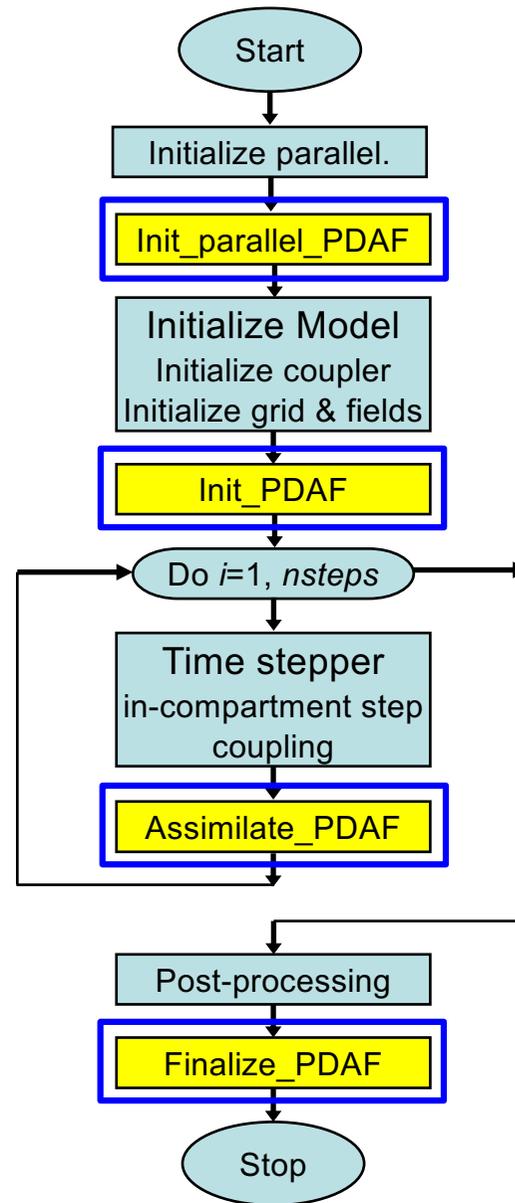
Open source:
Code, documentation & tutorials at
<http://pdaf.awi.de>

Extending a Model for Data Assimilation

Model
single or multiple executables
coupler might be separate program



Extension for data assimilation



plus:
 Possible model-specific adaption
 e.g. in NEMO: treat leap frog time stepping

revised parallelization enables ensemble forecast

(similar to EMPIRE, but more efficient)

Assumption: Users know their model

→ let users implement assimilation system in model context

For users, model is not just a forward operator

→ let users extend their model for data assimilation

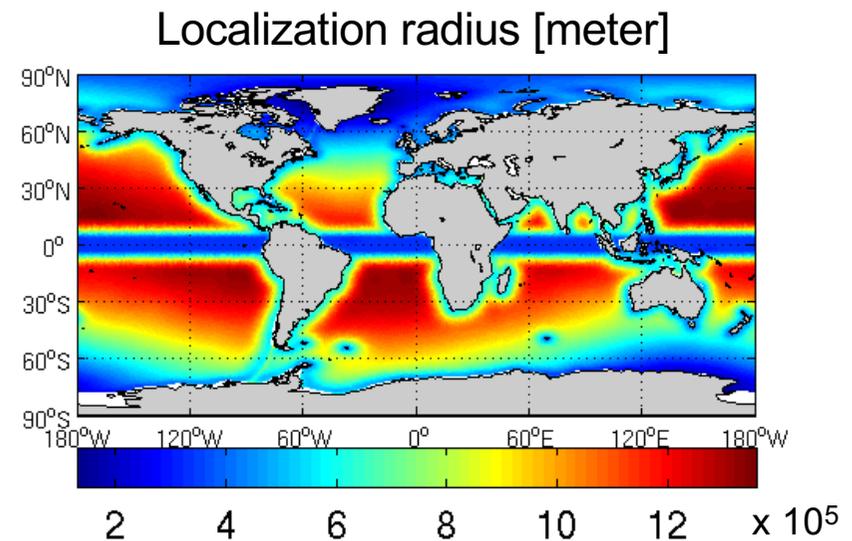
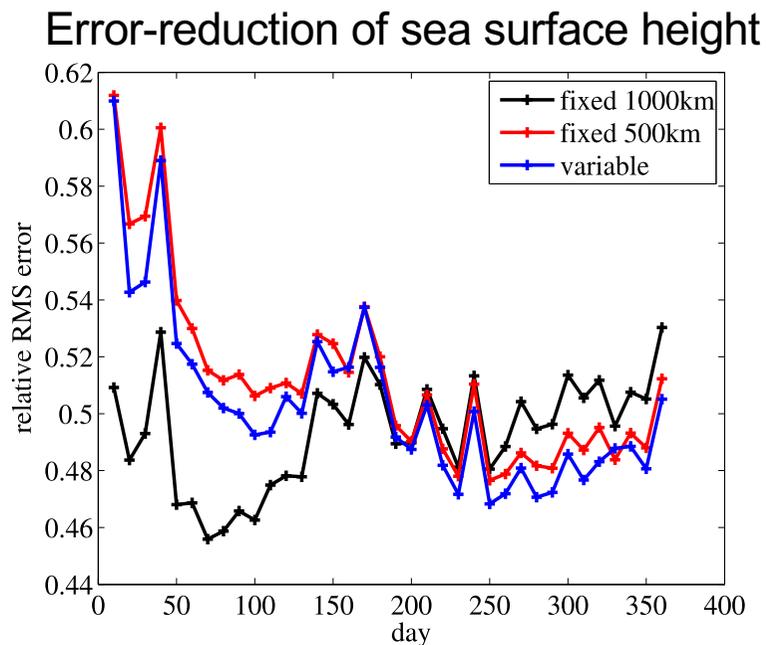
Keep simple things simple:

- Define subroutine interfaces to separate model and assimilation based on arrays
- No object-oriented programming (most models don't use it; most model developers don't know it; not many objects would be involved)
- Users directly implement observation-specific routines (no indirect description of e.g. observation layout)

Example: Value of Efficient Software

Adaptive Localization (Kirchgessner et al, 2012)

- Original study done with small models (Lorenz-96, shallow water)
 - Paper reviewer asked to apply it with full-scale forecast model
 - FESOM with PDAF was fully coded without adaptivity
 - Update PDAF library (just when recompiling)
 - Adding adaptivity routine and running experiment
- } 1 day!



Applications

Application Example

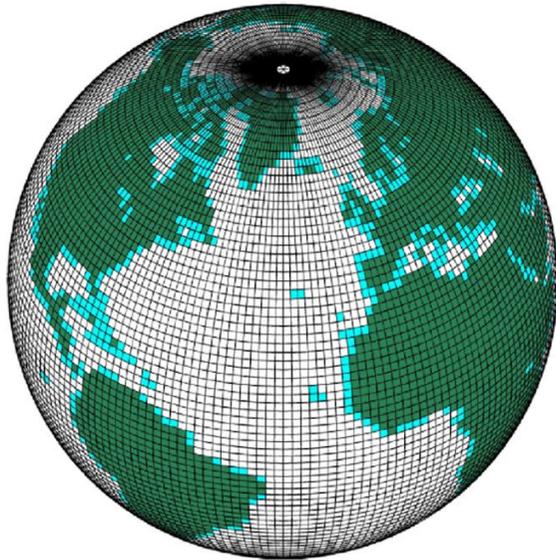
Coupled Atmosphere-Ocean Data Assimilation

Qi Tang

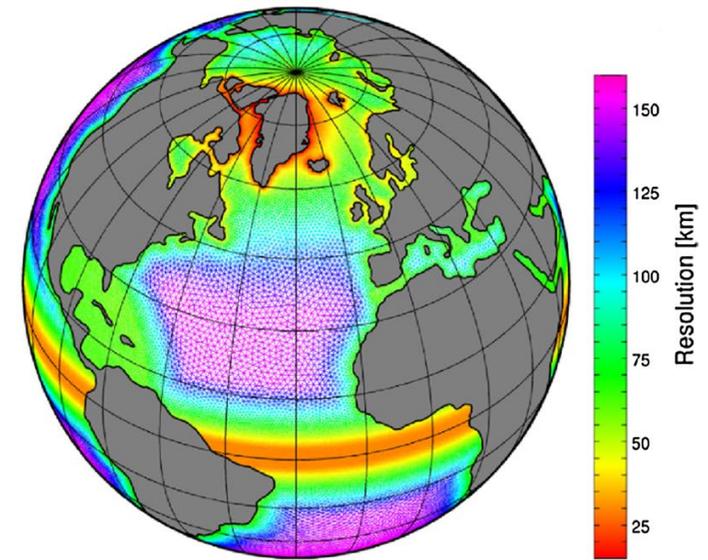


Example: ECHAM6-FESOM (AWI-CM)

Atmosphere



Ocean



OASIS3-MCT

fluxes



ocean/ice state

Atmosphere

- ECHAM6
- JSBACH land

Coupler library

- OASIS3-MCT

Ocean

- FESOM
- includes sea ice

Two separate executables for atmosphere and ocean

Goal: Develop data assimilation methodology for cross-domain assimilation (“strongly-coupled”)

Execution Times (weakly-coupled, DA only into ocean)

MPI-tasks

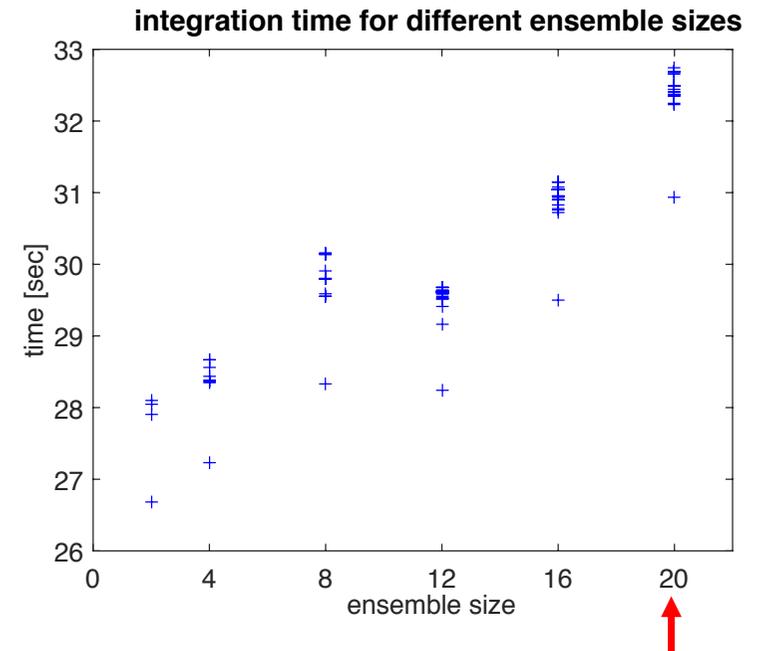
- ECHAM: 144
- FESOM: 384

Timings (1 day):

- Ens. forecast: 27 – 33 sec
- Analysis step: 0.5 – 0.9 sec

Scalability:

- Slowly increasing integration time with growing ensemble size (only 16% due to more parallel communication)
- some variability in integration time over ensemble tasks
- Need optimal distribution of programs over compute nodes/racks (here set up as ocean/atmosphere pairs)



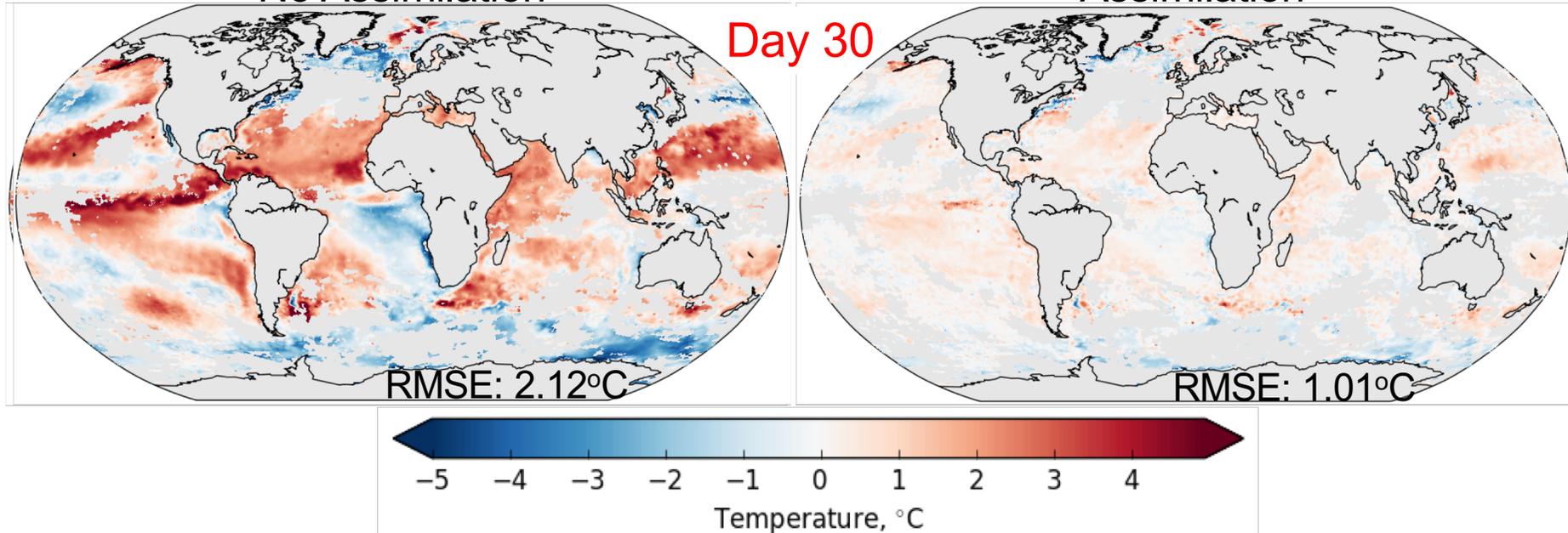
10,560
processor
cores

Assimilation Effect on Surface Temperature

Difference between model simulations and observations

No Assimilation

Assimilation



Surface temperature assimilation successful over 1 year

- Vertical localization required to avoid unrealistic subsurface temperatures

Current work

- Add subsurface profile data (temperature & salinity)
- Assess effect on atmosphere
- Final aim: strongly-coupled assimilation
(e.g. improve atmospheric state using ocean observations)

Application Example

Assimilation of Satellite Ocean Color Data into Ocean-biogeochemical Model

Himansu Pradhan

IPSO

Coupled Model: MITgcm - REcoM

MITgcm

General ocean circulation model of MIT (*Marshall et al., 1997*).

Global configuration

80°N - 80°S, 30 layers

Resolution:

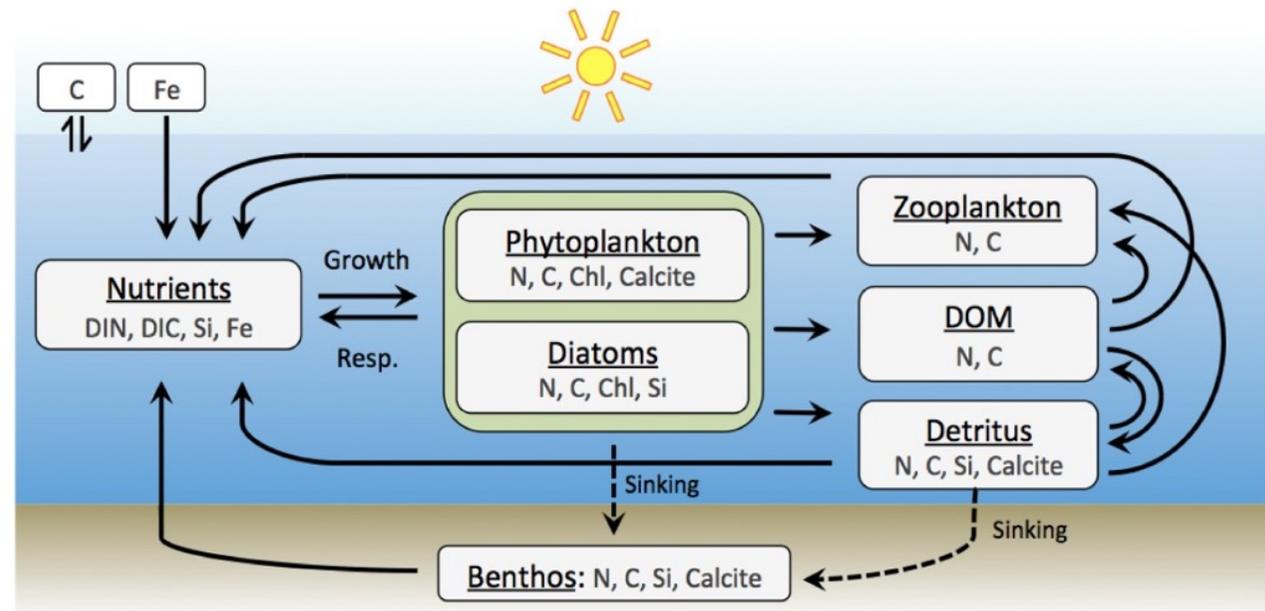
lon : 2 deg

lat : 2 deg in North
up to 0.38 deg in South

layers : 10 m – 500 m

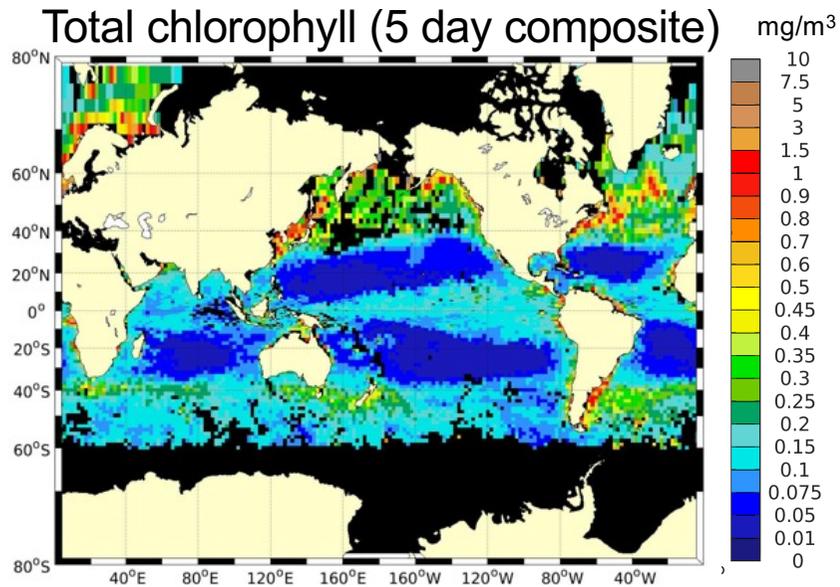
REcoM-2

Regulated Ecosystem Model – Version 2
(*Hauck et al., 2013*)

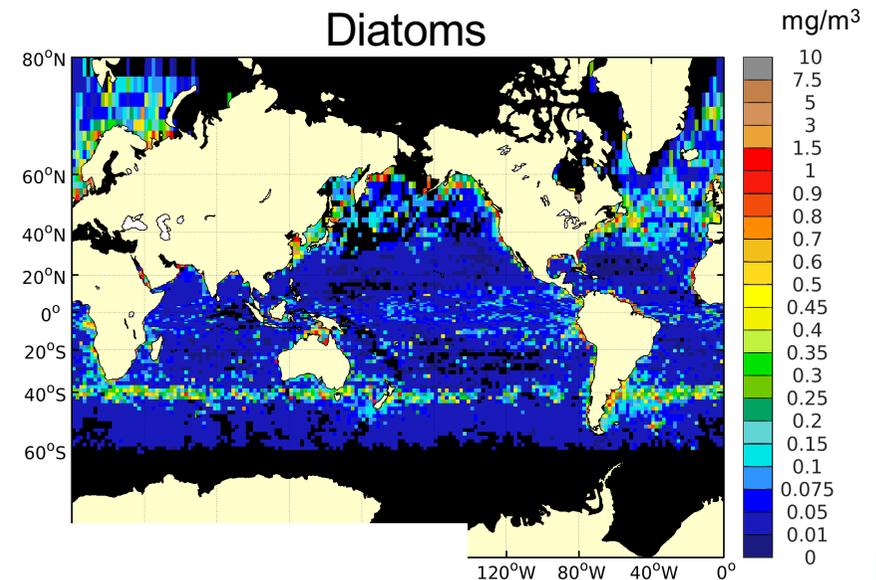
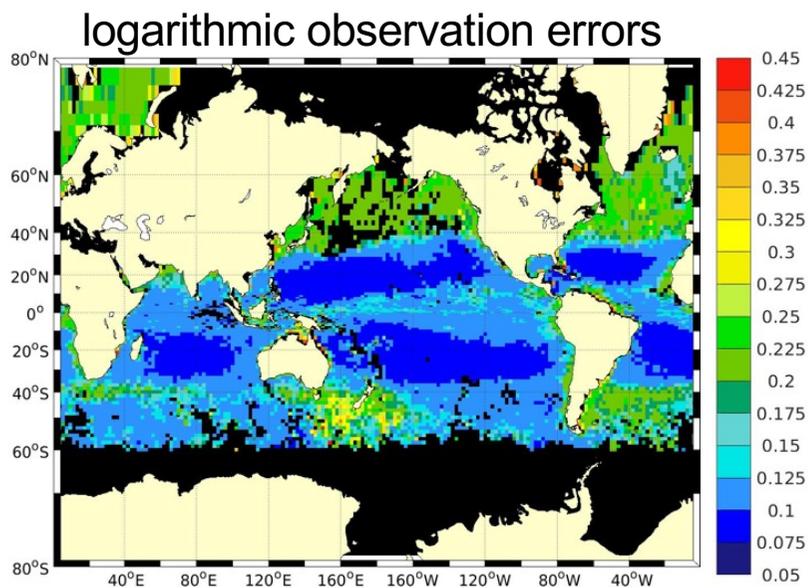
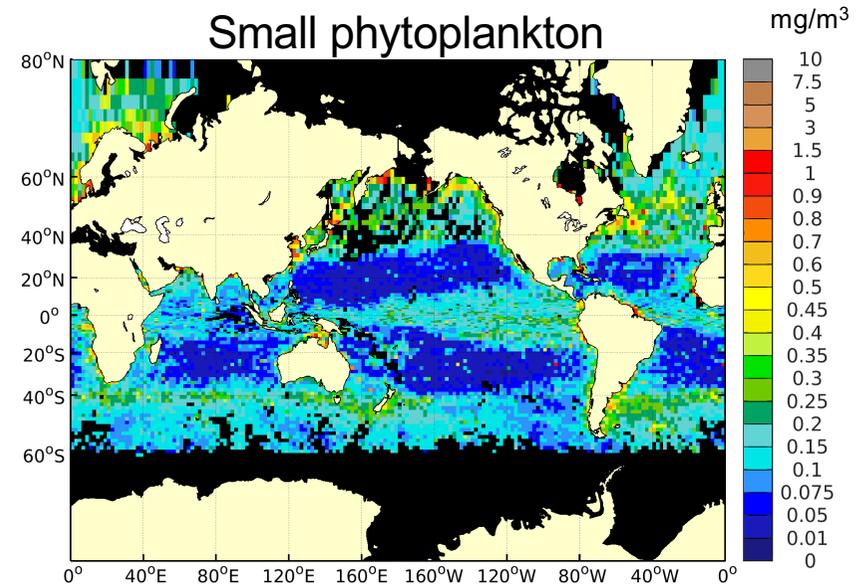


Assimilation of Total Chlorophyll

Assimilated:
Total chlorophyll from ESA OC-CCI

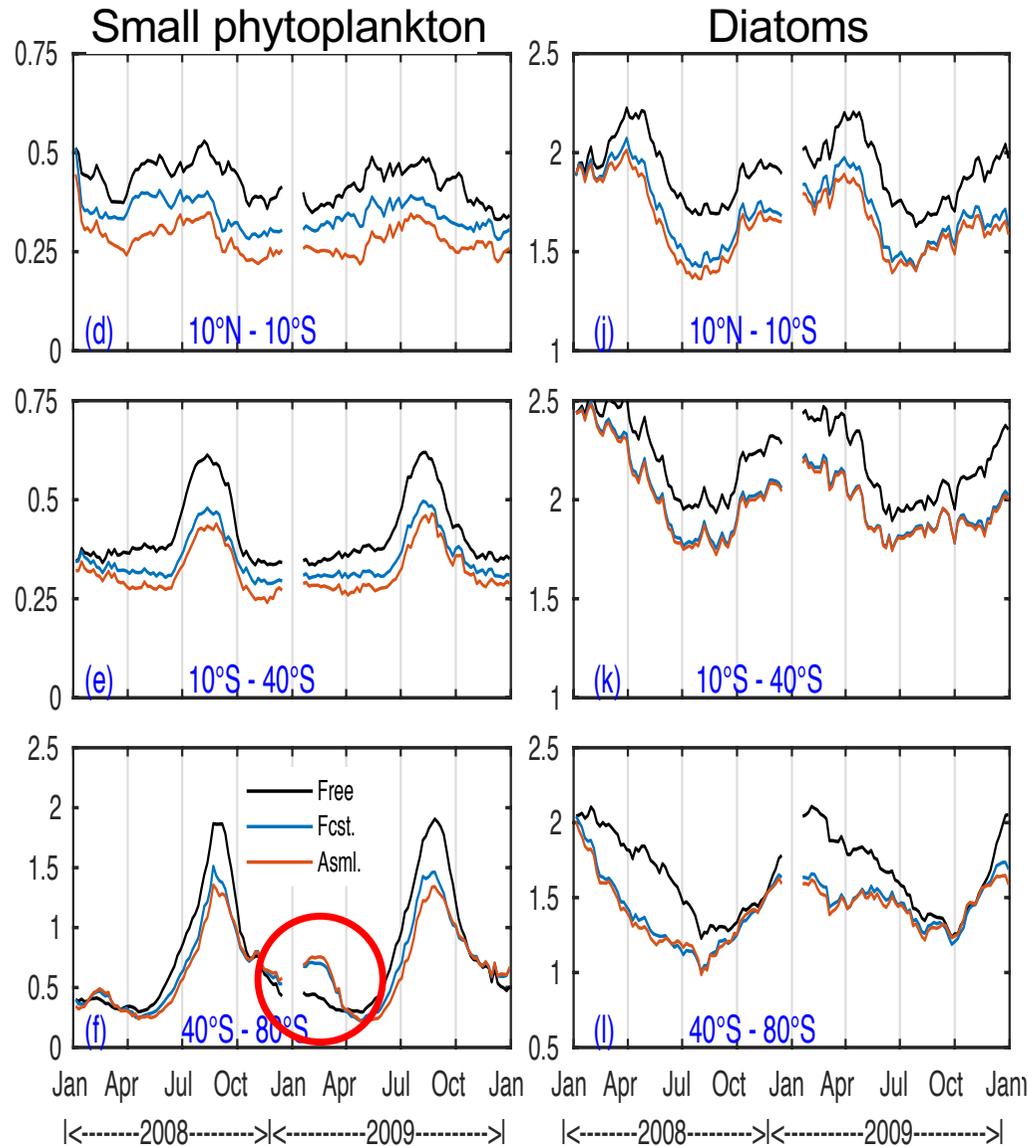


Verification: Phytoplankton group data
SynSenPFT (Losa et al. 2018)



Effect on Chlorophyll in Phytoplankton Groups

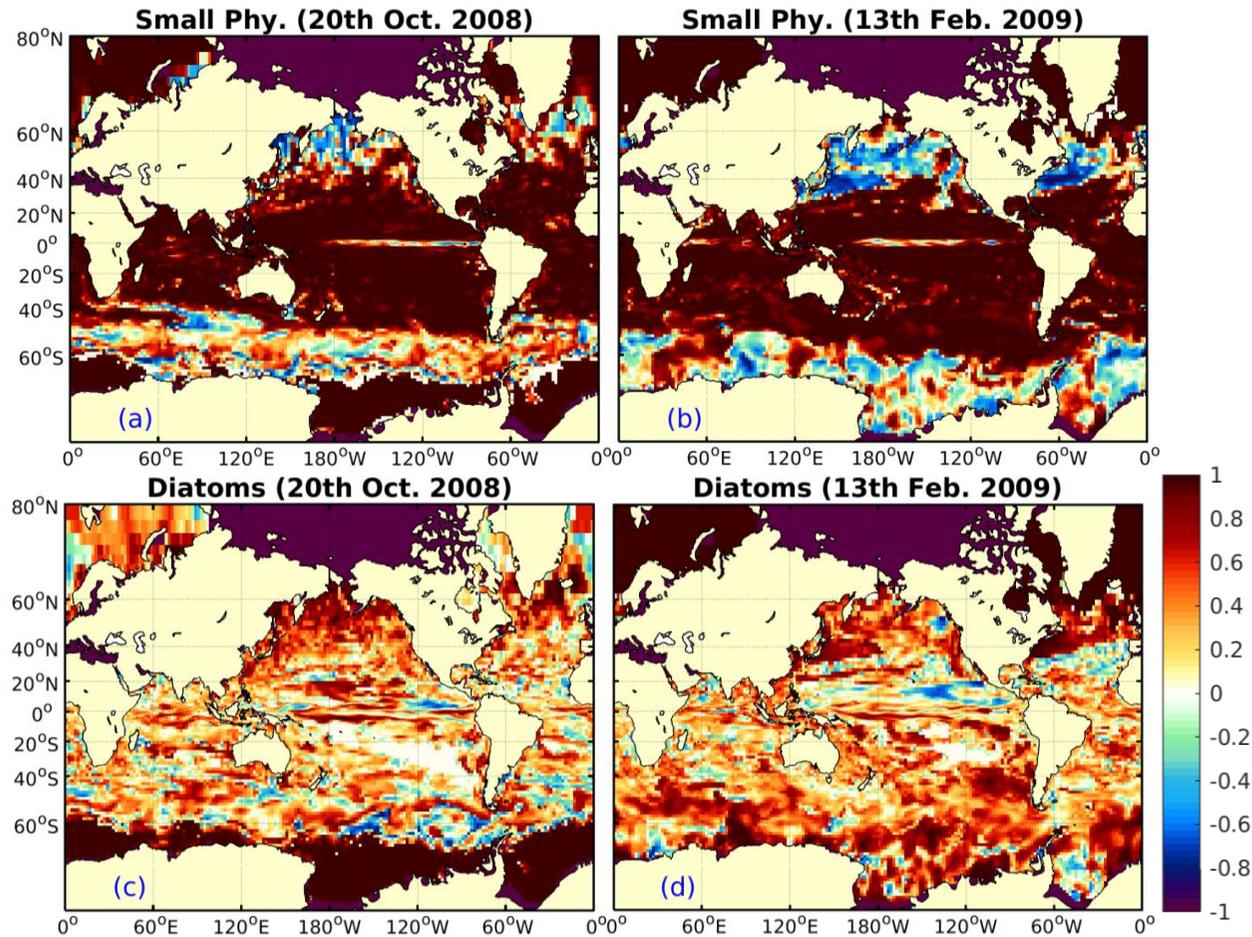
logarithmic RMS errors (southern regions)



- Assimilation improves groups individually through cross-covariances
- Stronger error-reductions for Diatoms
- Southern Ocean: Particular effect for small phytoplankton at very low concentration
- Current work
 - Asses impact of assimilating chlorophyll group data

Ensemble-estimated Cross-correlations

Cross correlations between total and group chlorophyll



- Significantly different correlations for small phytoplankton and diatoms
- Negative correlations exist