

QOS 2021

October 3 (Sunday) - 9 (Saturday), 2021

 Online Meeting

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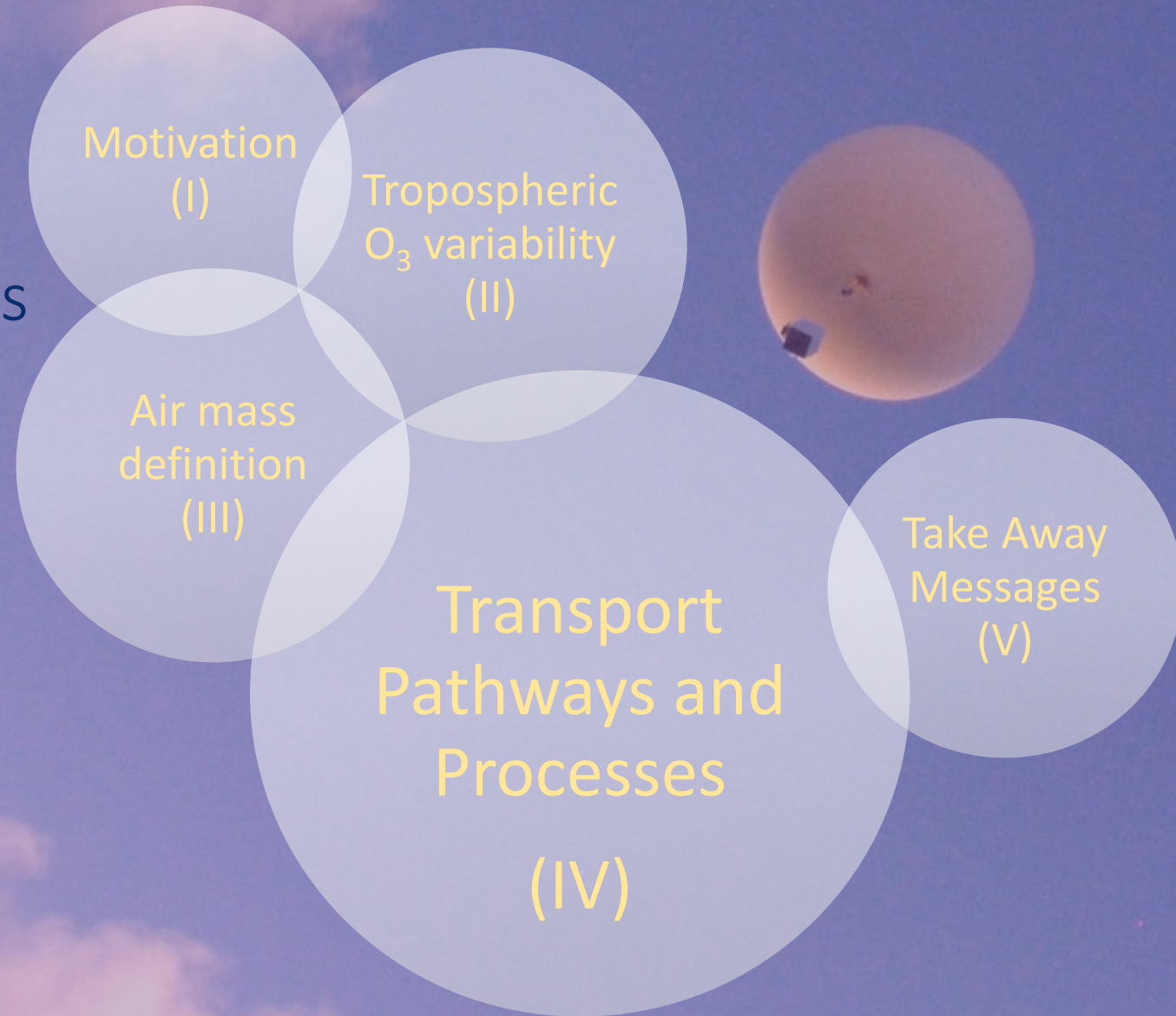
Origin of Tropospheric Air Masses in the Tropical West Pacific and related transport processes inferred from balloon-borne Ozone and Water Vapour observations from Palau



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Contact: Katrin.Mueller@awi.de

PhD thesis + 2 Manuscripts *



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UND MEERESFORSCHUNG

NAVIGATION



Main Menu

(X)

Chapters

*



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Why the Tropical West Pacific (TWP)?

Major source region for stratospheric air in boreal winter

Persistent tropospheric **Ozone minimum**

Origin and transit region of corresponding air masses in boundary layer and troposphere
(Rex et al. 2014)

PALAU
7°N 134°E



clean air

Corresponding **OH minimum**

and prolonged life times of various chemical species

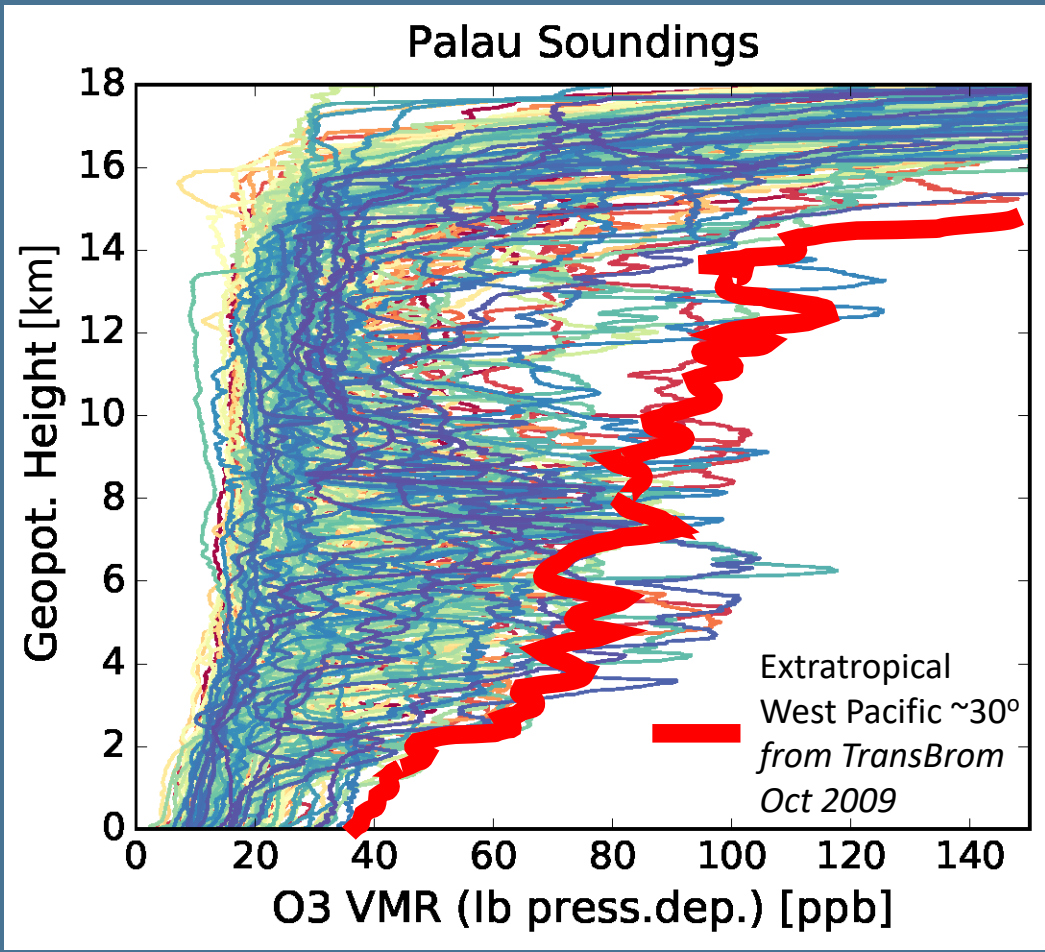
Important region for supply of chemical species to the stratosphere *

Need for monitoring of air composition and understanding of underlying processes and transport pathways to TWP

Key feature of the clean TWP troposphere: close coupling of the O_3 concentration and oxidizing capacity (OH), influencing overall transport of chemical species to the stratosphere.

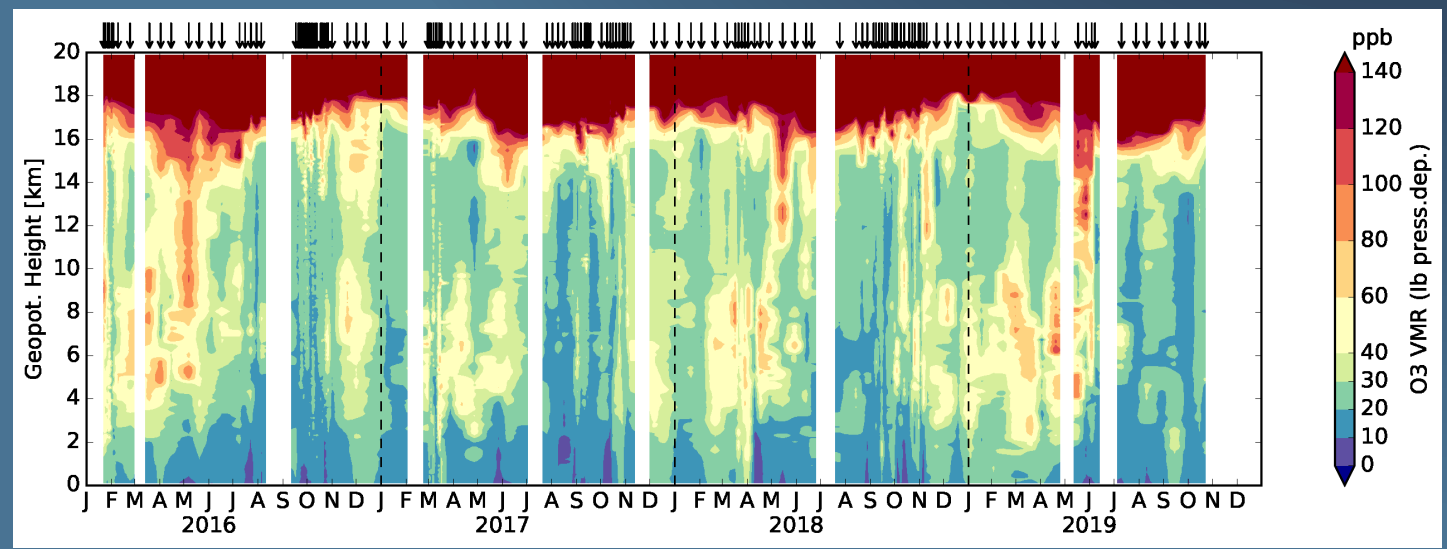
To improve the limited availability of tropospheric O_3 observations from this key region, the **Palau * Atmospheric Observatory** was established in 2016 as part of the EU-project StratoClim.





First characterization of tropospheric O₃ seasonality in the TWP with a multi-year continuous time series from ECC ozonesonde measurements every two weeks or in intensive campaigns (SPC 6A, Vaisala RS92/41). → Müller 2020

Special focus on quality issues of tropical soundings due to controversy around near-zero O₃ observations in the TWP (*e.g. Voemel and Diaz 2010, Rex et al. 2014, Thompson et al. 2019*)



Tropospheric profiles with altitude:
145 sondes, 01/2016-10/2019

Time-height-cross-section

(I) Why O₃? As a chemical tracer...



...for **local convective activity** in clean maritime air: **“low” O₃**
(e.g. *Folkins, 2002; Folkins et al., 1999; Kley et al., 1996; Paulik and Birner, 2012; Solomon et al., 2005*) and

... for **long range transport** processes to the region, either related to air pollution or stratospheric intrusions: **“high” O₃**
(e.g. *Andersen et al., 2016; Browell et al., 2001; Randel et al., 2016; Tao et al. 2018; Thouret et al., 2000; Pan et al. 2015*).

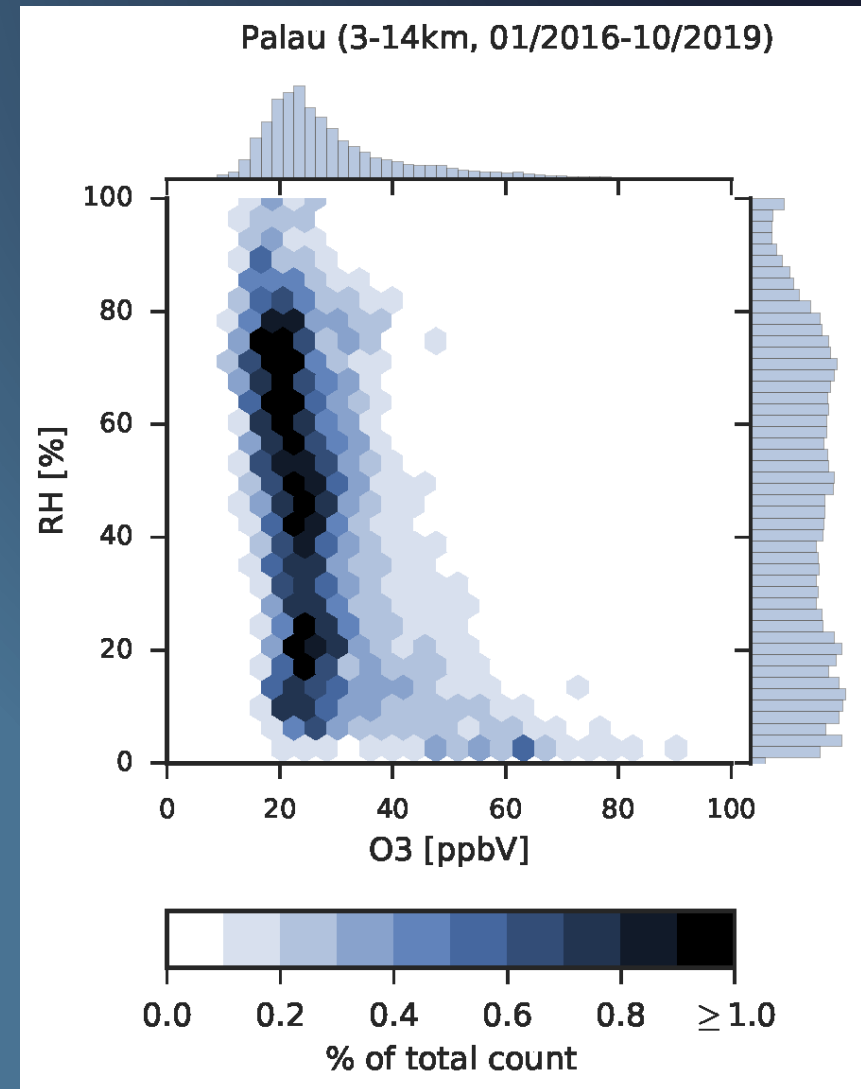
RH as a tracer for **vertical displacement**:

High humidity due to convection, dryness due to large scale descent
(e.g. *Hayashi et al., 2008; Andersen et al., 2016; Cau et al. , 2007; Dessler and Minschwaner, 2007*)

Central Question:

Can we identify air mass origin and its seasonality with the observed O₃/RH relation?

Free-tropospheric O₃/RH distribution

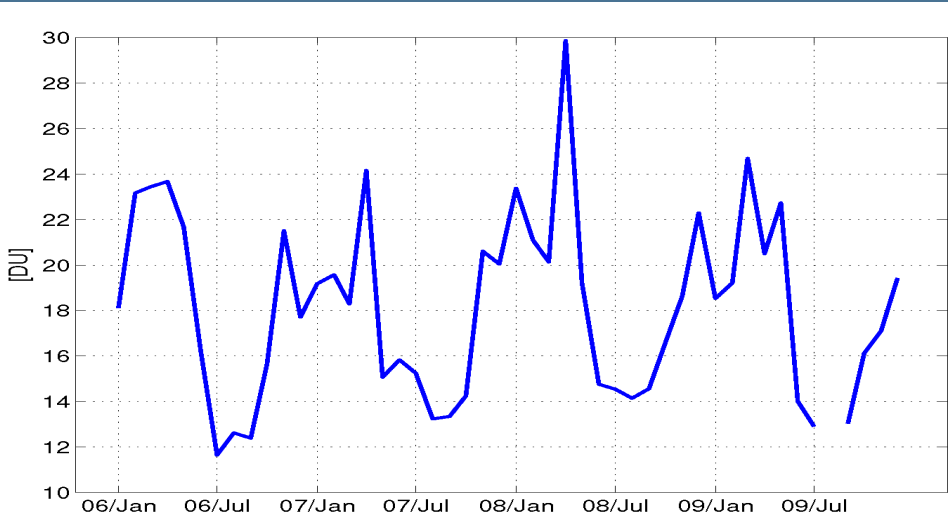


unique for Palau compared to stations of the tropical **SHADOZ ozonesonde network** *

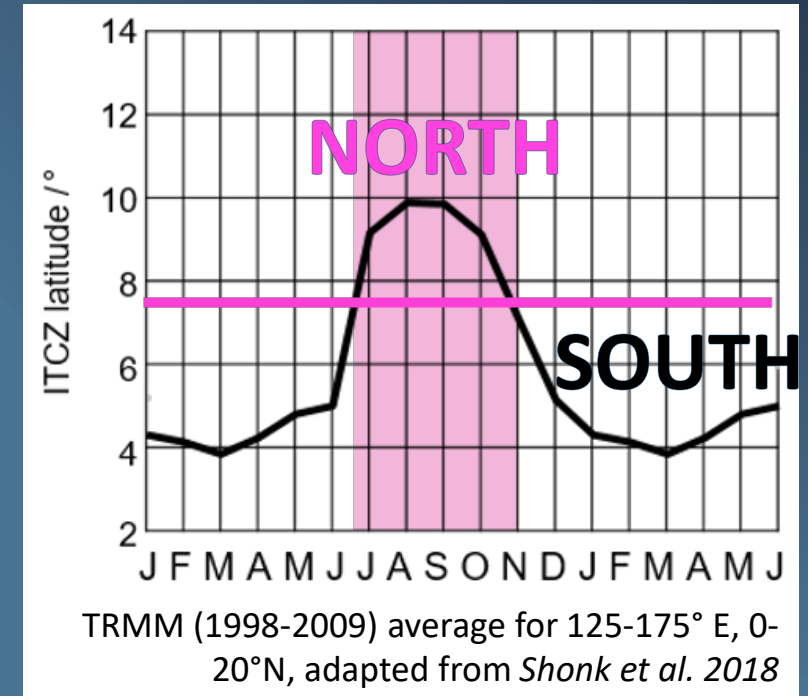
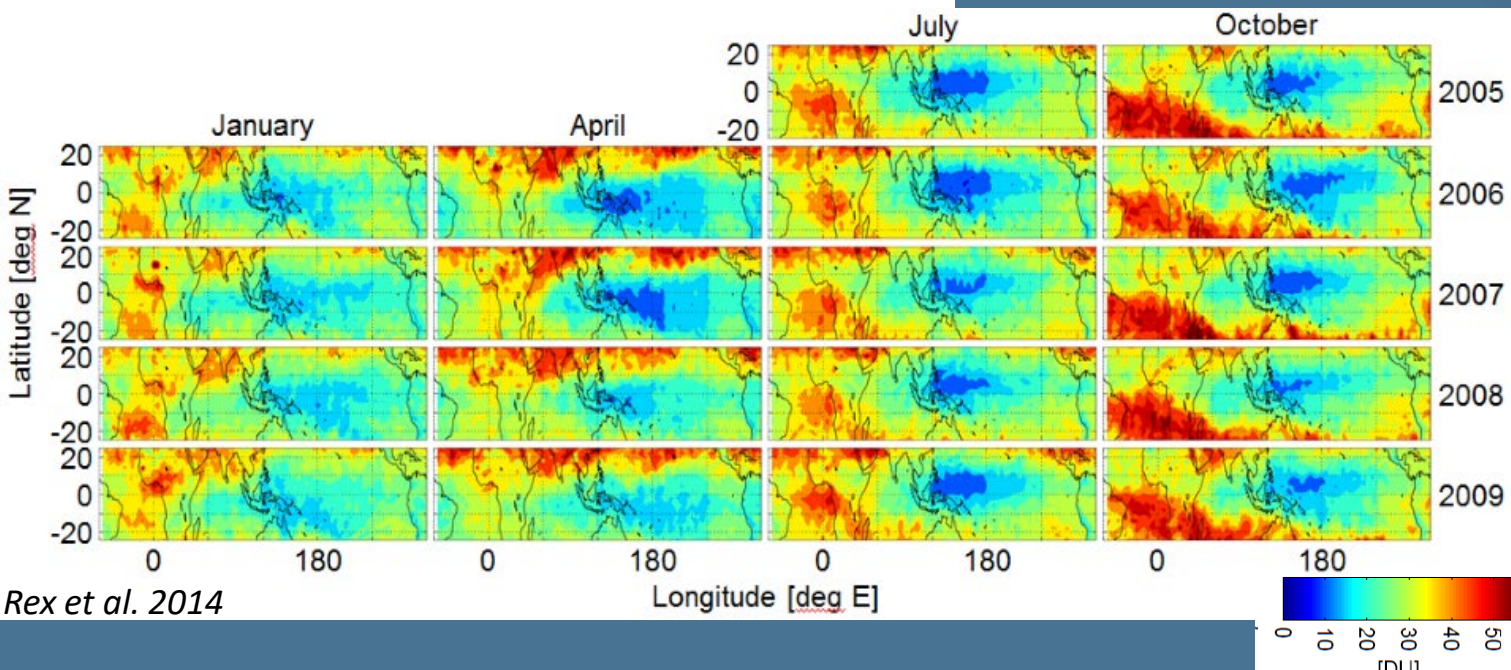
(II) Tropospheric O₃ variability



- **Seasonal drivers:** circulation (Walker, Hadley, West Pacific Monsoon, Brewer-Dobson) → modulated by Inter Tropical Convergence Zone (ITCZ)
- Hot, humid & wet climate all year: high convective activity
- **Important Variation:** ENSO



Tropospheric O₃ column from TES (on Aura Satellite) @ Palau and in the tropics



TRMM (1998-2009) average for 125-175° E, 0-20°N, adapted from *Shonk et al. 2018*

Movement of the ITCZ

Rex et al. 2014

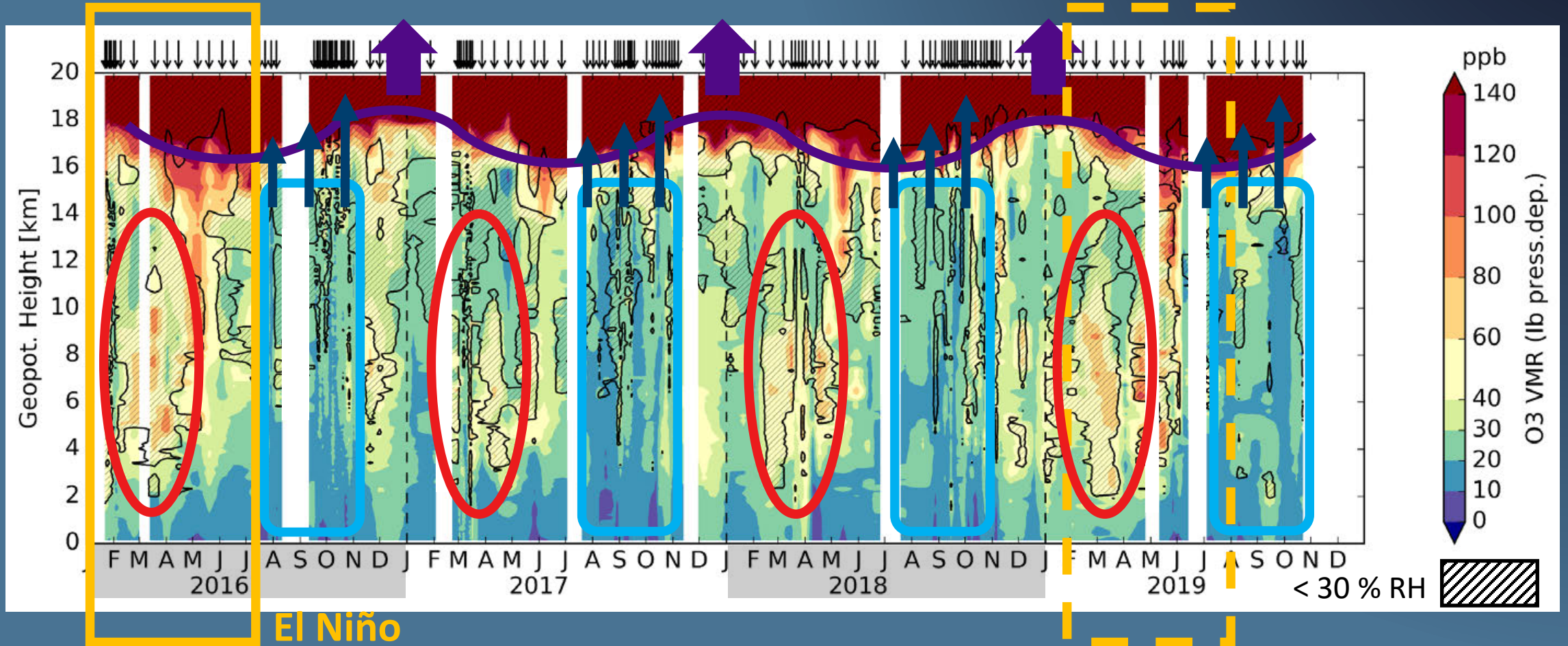
(II) Tropospheric O₃ variability



Annual TTL cycle: Brewer Dobson circulation *

+ enhanced high altitude convective outflow

Onset of transport to the stratosphere and the tropospheric O₃ minimum occur simultaneously



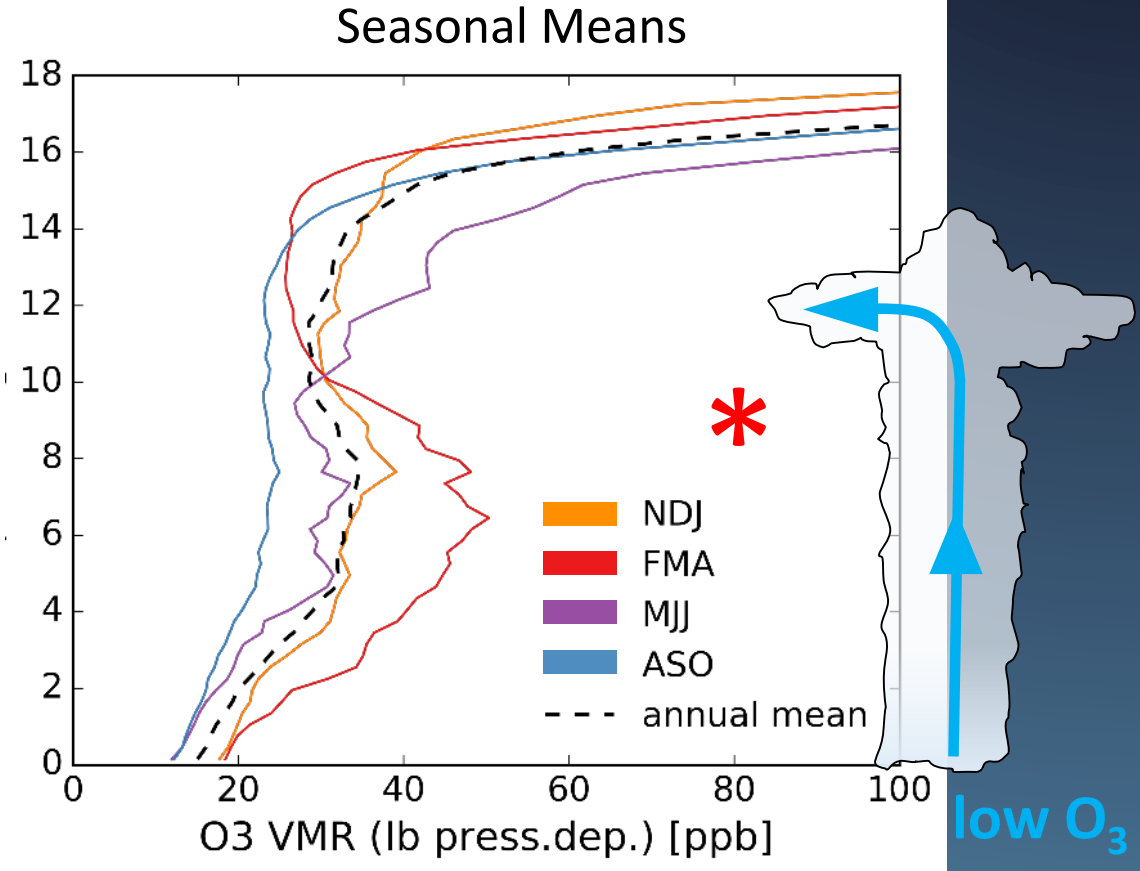
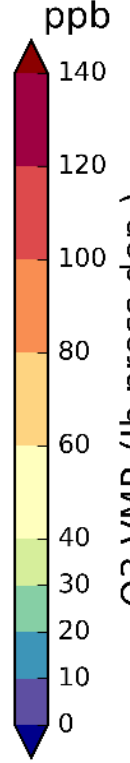
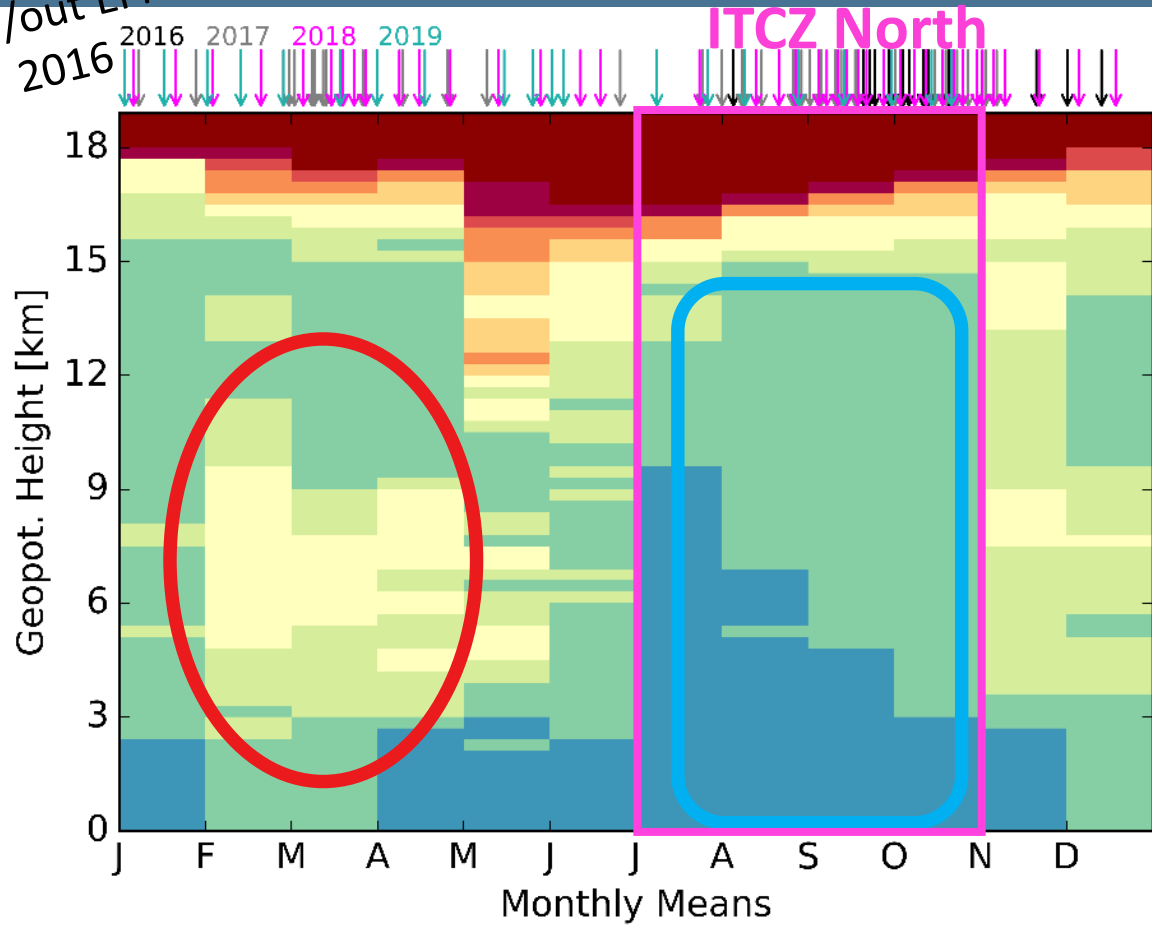
Mid-troposphere: O₃ minimum from July-October, layers of enhanced O₃ from February-April, often **anti-correlated with RH**

(II) Tropospheric O₃ variability



Annual mean: typical (tropical) „S-Shape“, monthly means grouped according to similar shape: **4 distinct types of profiles (seasons)**.

w/out El Nino

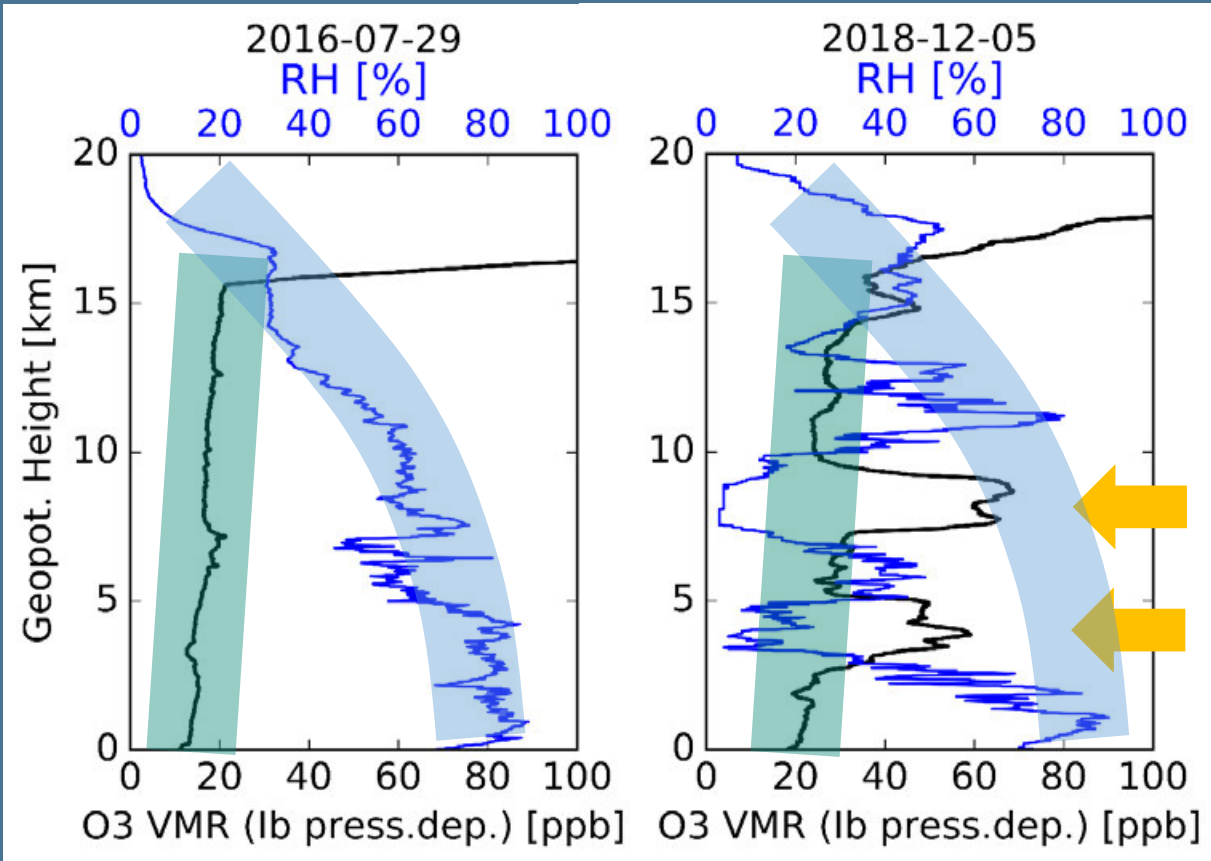


Monthly means highlight annual cycles, **O₃ minimum** corresponds with the **ITCZ located North** of Palau.

Deep convective detrainment can explain upper dent in the „S“ (10-14 km); between 5-10 km or the belly of the „S“: weak cloud-mass divergence, greatest anomalies from annual mean in ASO and FMA.



Example tropospheric O₃ and RH profiles



convectively controlled,
well-mixed **background:**

LOCAL mode

interruptive layers,
controlled by transport:

NON LOCAL mode

Underlying processes:

Local boundary layer air masses lacking pollution (\rightarrow low in O₃) are lifted locally by convection (**humid**), creating a uniform profile.

No known mechanism for in situ production of **high O₃** or **dehydration** in the mid-troposphere - origin either transport from the (extratropical) stratosphere or non-local ground pollution, lifted convectively in the area of origin then undergoing dehydration during transport, e.g. via large-scale descent and radiative cooling.

(compare Dessler and Minschwaner, 2007; Andersen et al., 2016)

Layered structures and respective background are hidden in **the belly of the „S“** of mean profiles!

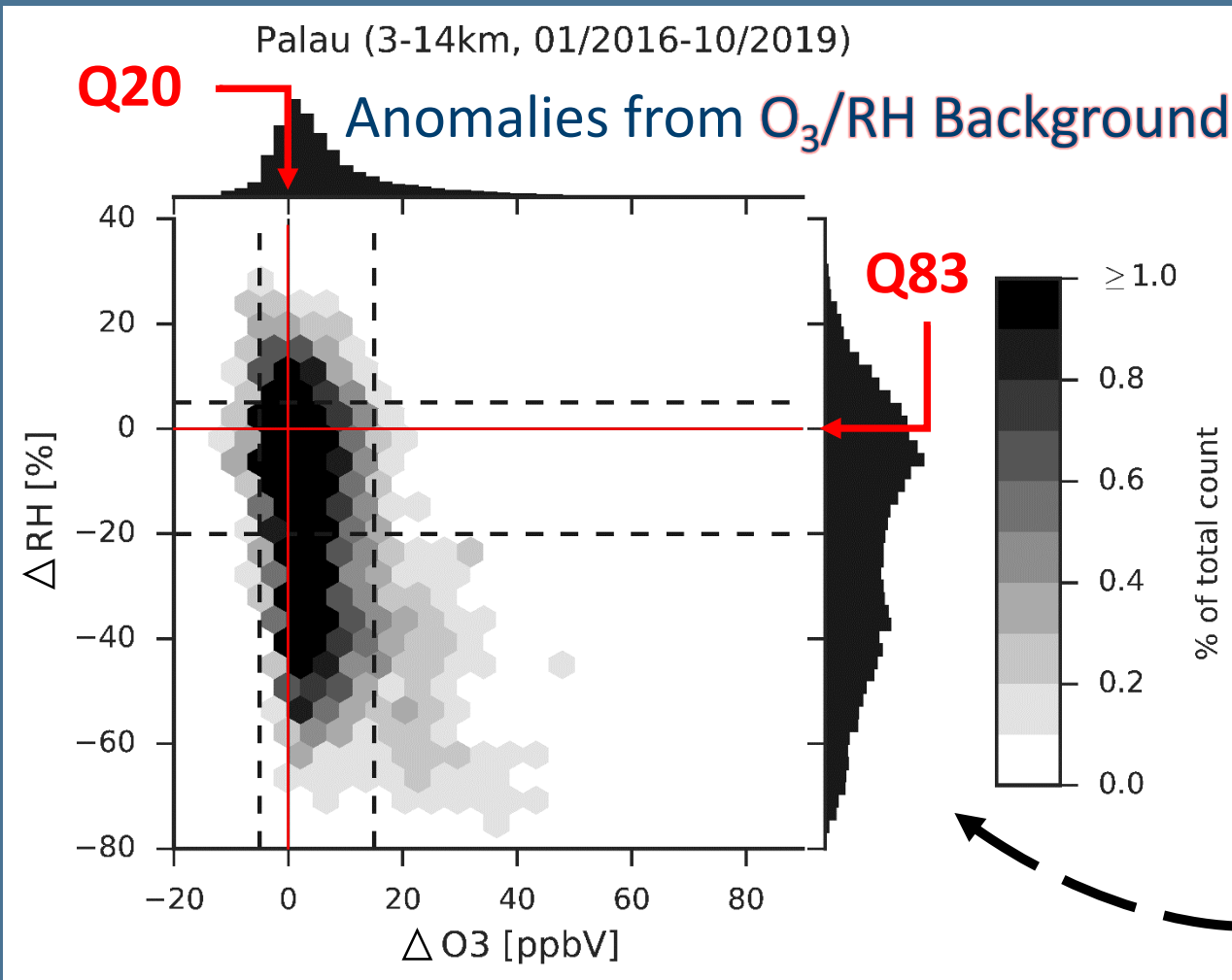
mode
LOCAL

mode
NON LOCAL

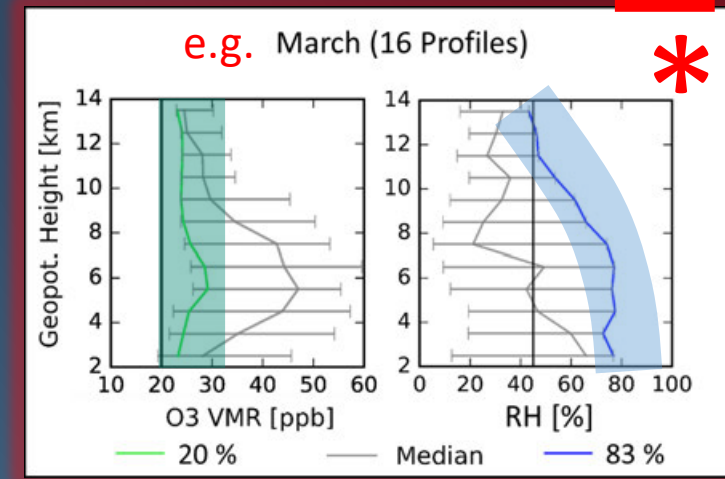
(III) Air mass definition



- First step: define **background profiles** for both tracers
- Second step: determine **anomalies** against this background

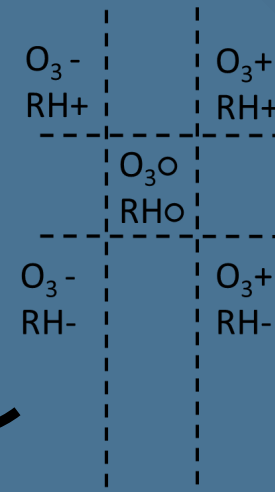


Background profiles: the monthly 20th (O₃) and 83rd (RH) quantile, **Q20 and Q83** (altitude dependent)



- Third step: **bimodality in RH anomalies** motivates classification in O₃RH groups

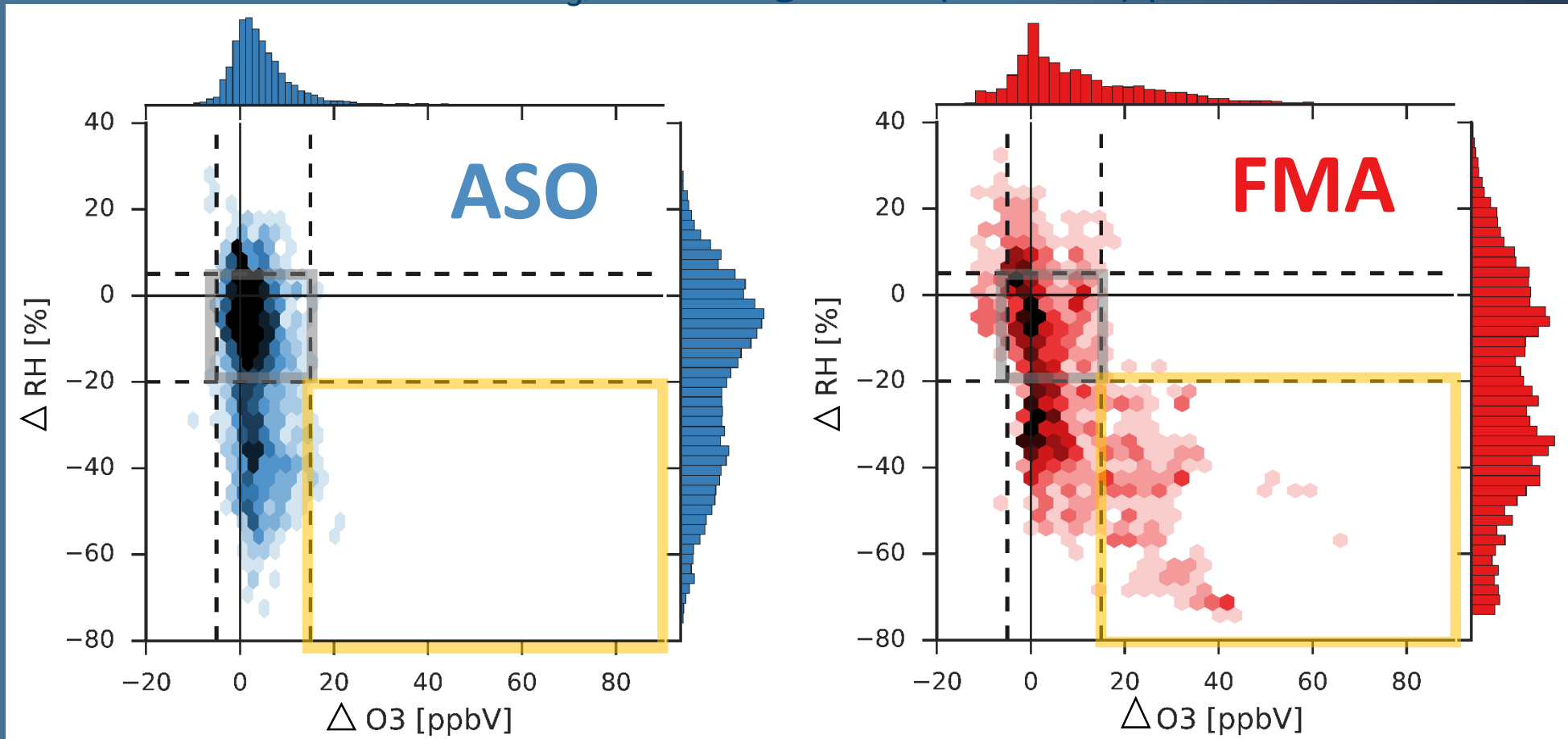
→ **air mass definition**



Classification in a 3x3 grid (dashed lines), with respect to the distributions, focus in the following on **central background group** (O₃o RHo) and **dry O₃-rich anomaly** (O₃+RH-)



Anomalies from O₃/RH Background (3-14 km) per season



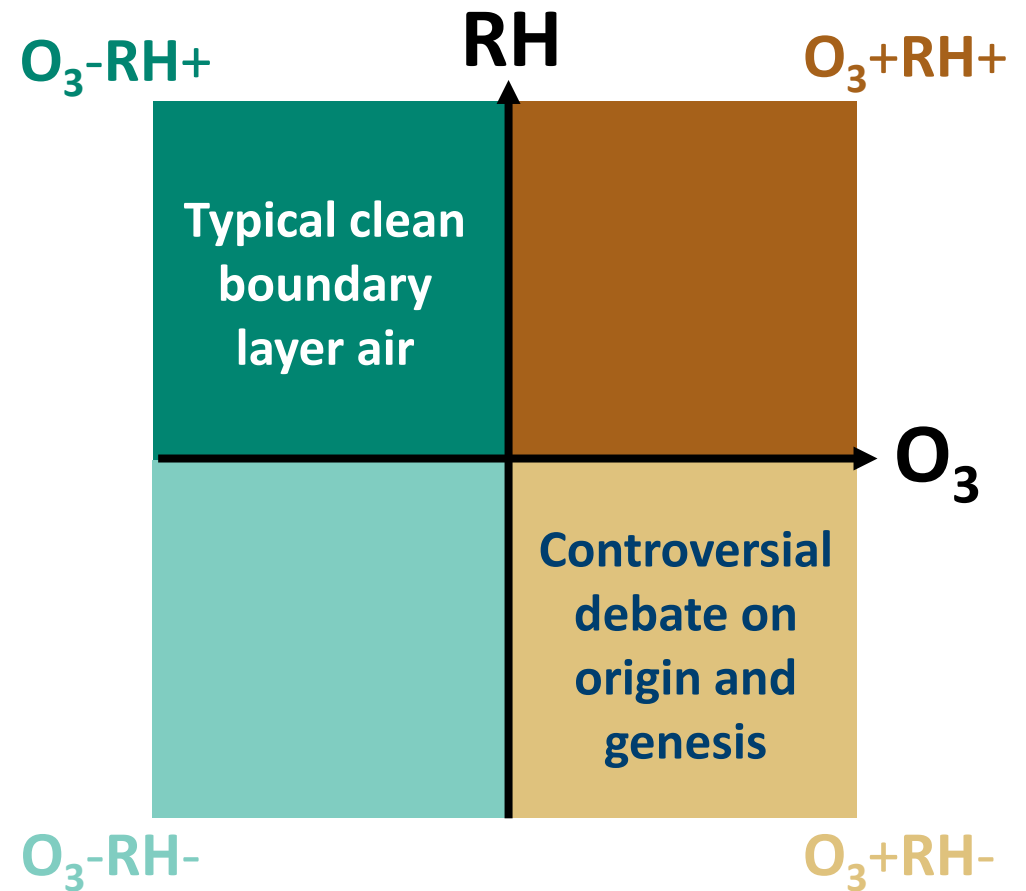
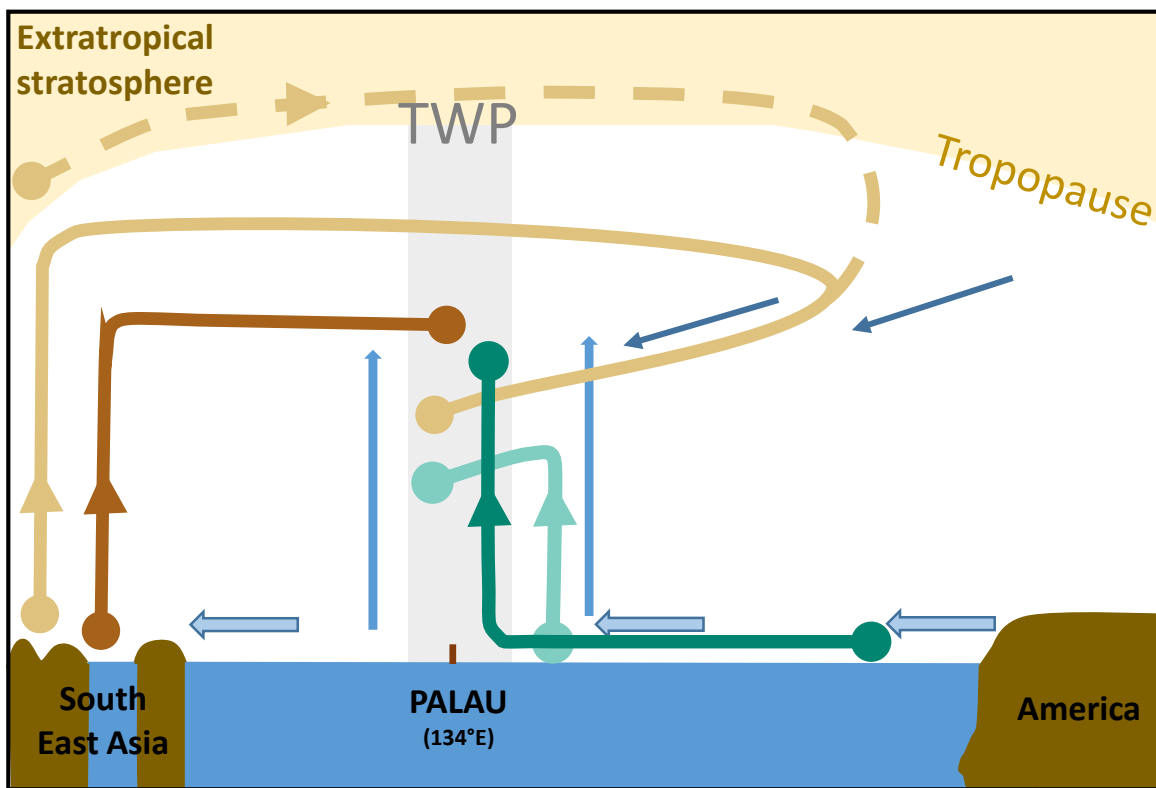
O₃oRHo:
humid, O₃-poor
background, present
year-round (!), but
dominates ASO

O₃+RH- :
dry, O₃-rich, most
frequent in FMA





With our process understanding, we identified major transport pathways related to the O_3 RH relation observed in Palau:



← Trade winds ↑ Convective uplift ↘ Clear sky cooling

(IV) Backward Trajectories



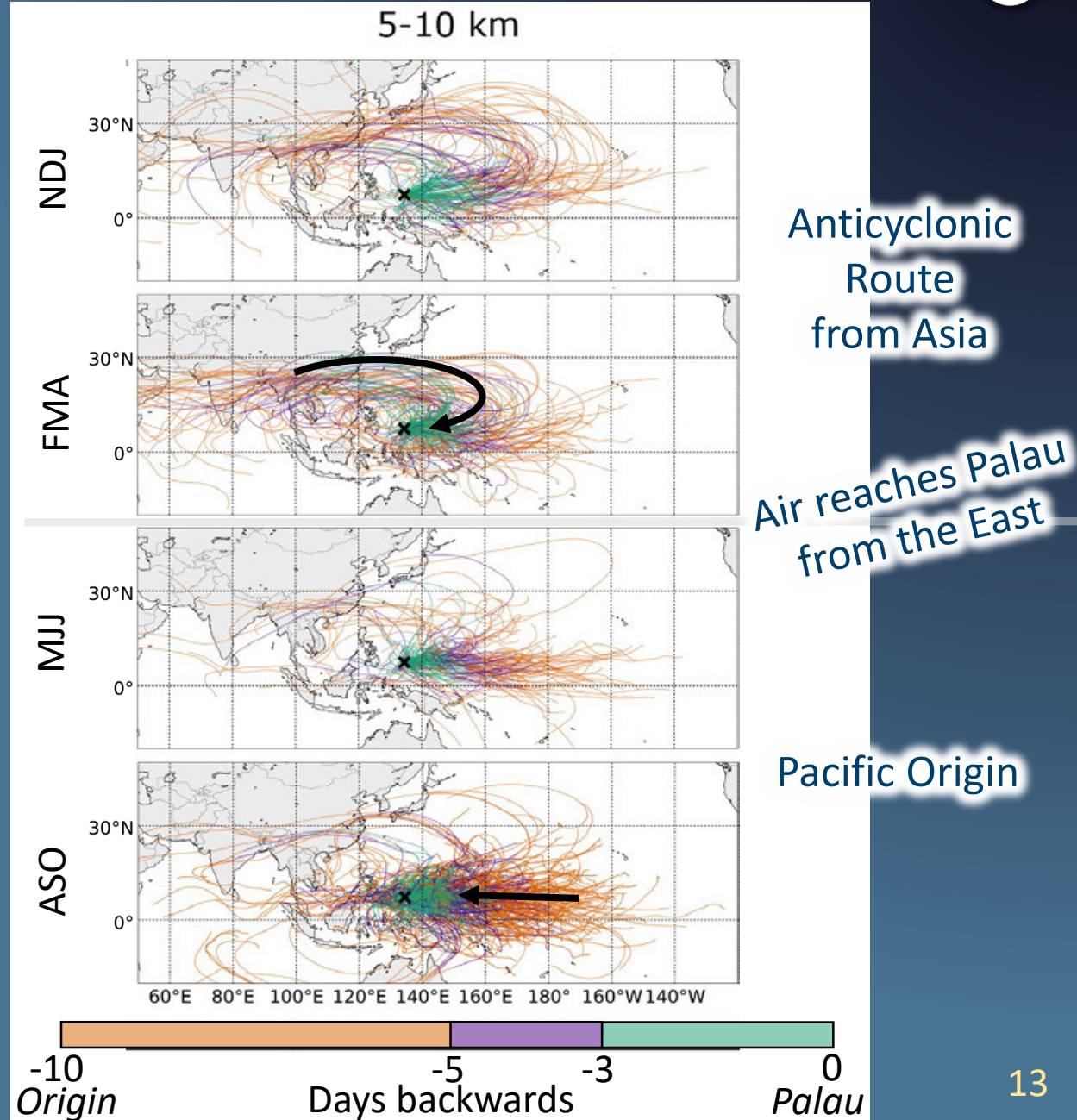
Transport module of Langrangian Chemistry and Transport Model ATLAS (Wohltmann et al. 2010)

Setup:

- driven by ERA5 reanalysis data, no diffusion, no convective model parameterization, 10-min time steps
- initialized from ozone sounding data, 01/2016-10/2019, 2-14 km, every 10th measurement
→ focus on 5-10 km altitude range

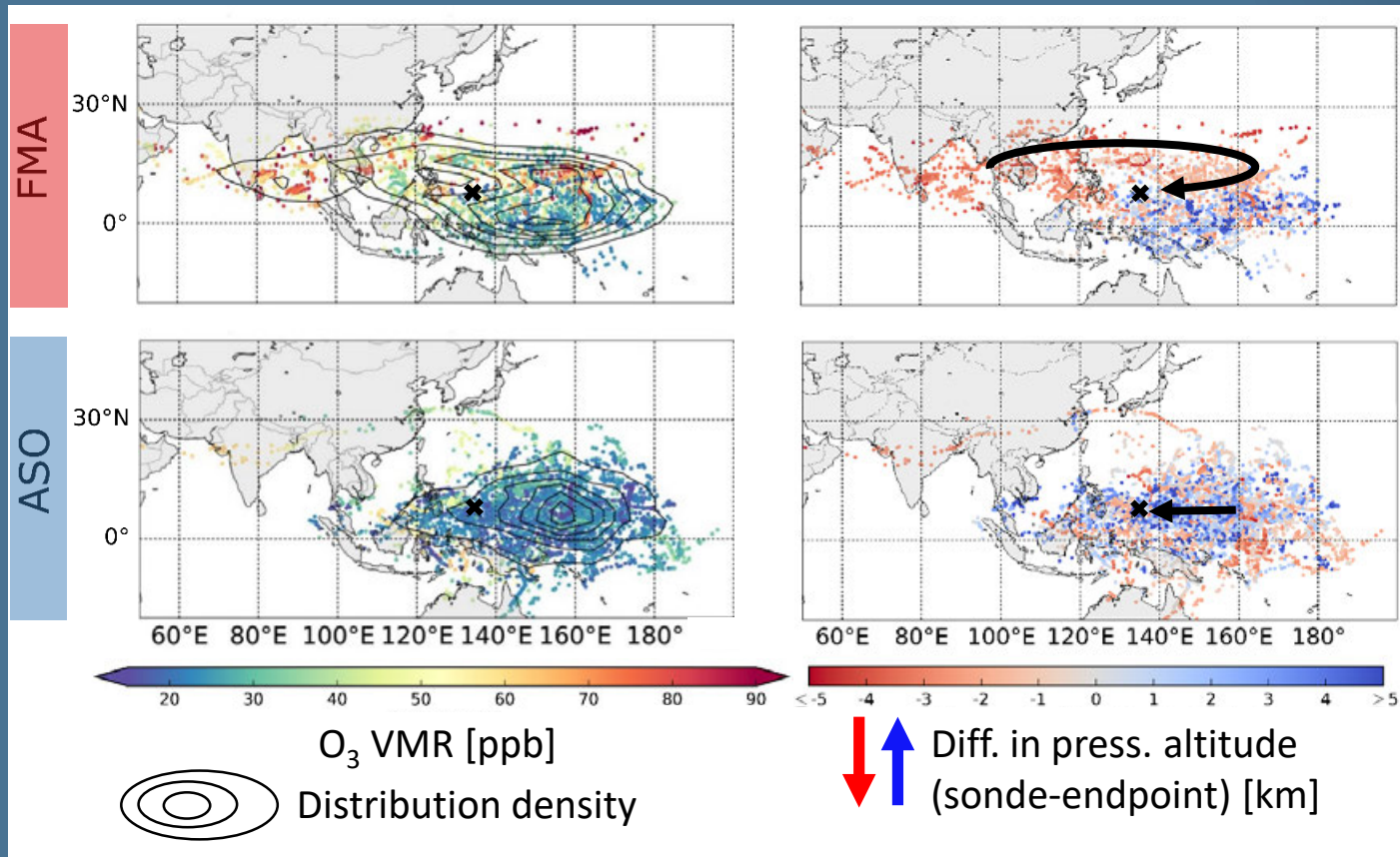
Assumptions:

- 10-day-backtrajectories for dynamical footprint
- Due to typical lifetime of marine boundary layer O_3 :
5-day-backtrajectory ending points
= origin of air mass composition





5-days-back trajectory ending points \equiv origin,
trajectory start @ 5-10 km in Palau **x**



All Observations per season:

O₃ VMR distributions:

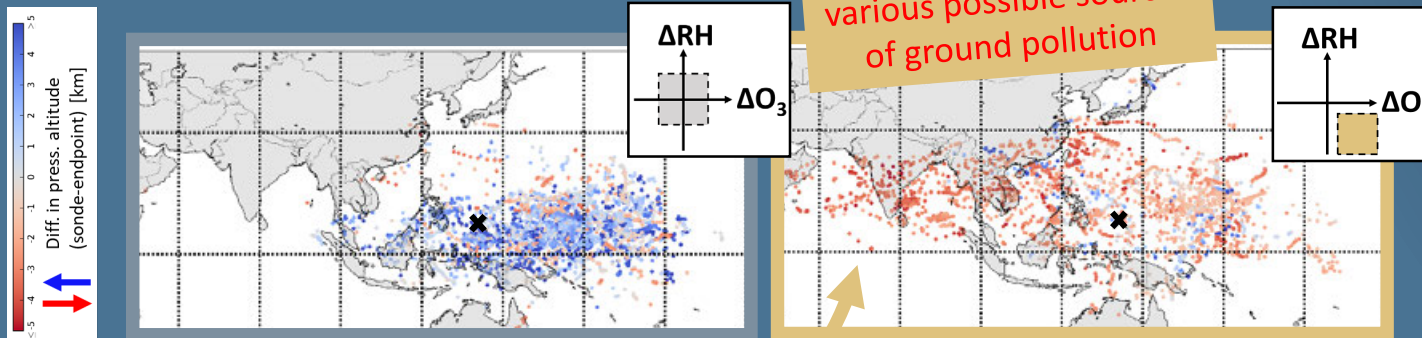
- Center of **low O₃** in both seasons, FMA and ASO, East of Palau
- Secondary center of **enhanced O₃** in FMA, North of Palau from India to East China

Vertical displacement:

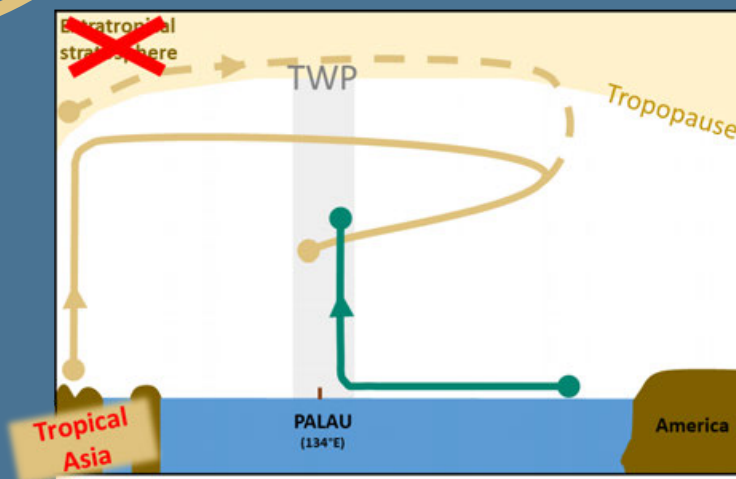
- Mainly in FMA, North of Palau air masses **descend** towards Palau (**anti-cyclonic route**), consistent with large-scale descent within the Hadley circulation and subsequent dehydration
- **Ascent** dominates ASO air masses (**Pacific origin**), corresponding well with the dominance of convective uplift



5-days-back trajectory ending points \equiv origin,
trajectory start @ 5-10 km in Palau **x**



Origin of **dry O₃-rich** air masses in areas of increased air pollution on the ground from industry or bio mass burning, speaking in favor for a **pollution based origin**



All Observations per O₃/RH group *

Selection of trajectories for air masses identified as humid, O₃-poor background (O₃o RHo) or **dry O₃-rich (O₃+RH-)** anomaly from the background separates air masses according to the processes controlling RH (Convective **uplift**, ASO; dehydration during **descent**, FMA) and locates spatially separate source regions

No indication for significant contribution of **stratospheric air**:

Potential Vorticity analysis for all trajectories (from 4 years, 138 profiles, 5-10 km) revealed essentially **no air mass crossing the 1.5 PVU threshold** for more than a day during 10 days backwards.



- ✓ **Palau's four-year tropospheric O₃ time series** fills the observational gap in this key region of stratospheric entry.
- ✓ Using the ECC O₃ sounding data set (01/2016-10/1019), seasonal analysis, trajectory modelling and a statistical approach to distinguish air masses by O₃/RH relation, we **identified transport processes and pathways to the TWP:**

	Humid, O ₃ -poor	Dry, O ₃ -rich
Processes	Convective background	Large scale descent, pollution
Origin	Pacific or local	Tropical Asia (anticyclonic route)
Frequency	Year-round, dominates Aug-Oct	Most frequent in Feb-Apr

- ✓ **Watch out** for the upcoming publications!

References



Katrin Müller, 2020: Characterization of Ozone and the Oxidizing Capacity of the Tropical West Pacific Troposphere, **PhD thesis**, <https://doi.org/10.26092/elib/463>.

WATCH OUT: two papers to be submitted this fall:

K. Müller, Ingo Wohltmann, Peter von der Gathen and Markus Rex (2021): Air Mass Transport to the Tropical West Pacific Troposphere inferred from Ozone and Relative Humidity Balloon Observations above Palau, in prep.

K. Müller, Jordis Tradowsky, Christoph Ritter, Justus Notholt, Ingo Beninga, Wilfried Ruhe, Winfried Markert, Juergen Graeser, Sharon Patris and Markus Rex (2021): The Palau Atmospheric Observatory – continuous monitoring of tropospheric composition in the Tropical West Pacific, in prep.

This presentation is based on an two earlier conference presentations:

K. Müller. et al. (2020): AGU presentation, Origin of Tropospheric Air Masses in the Tropical West Pacific identified by Balloon-borne Ozone and Water Vapor Measurements from Palau, <https://doi.org/10.1002/essoar.10505805.1>

K. Müller. et al. (2021): EGU presentation, Origin of Tropospheric Air Masses in the Tropical West Pacific identified by Balloon-borne Ozone and Water Vapor Measurements from Palau

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*** EXTRA SLIDES**



Palau Atmospheric Observatory

MaxDOAS:

Pandora2S –
Pandonia Network
O₃, NO₂, AOD
(, H₂O, SO₂, ...)

FTIR Spectrometer:

Total abundances of
~ 20 chemical species

Lidar:

Vertical profiles of
aerosol properties
Since 2018: multi-λ cloud
and aerosol lidar ComCAL
in new lab

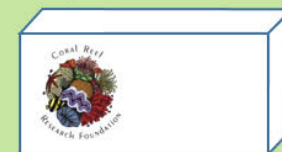
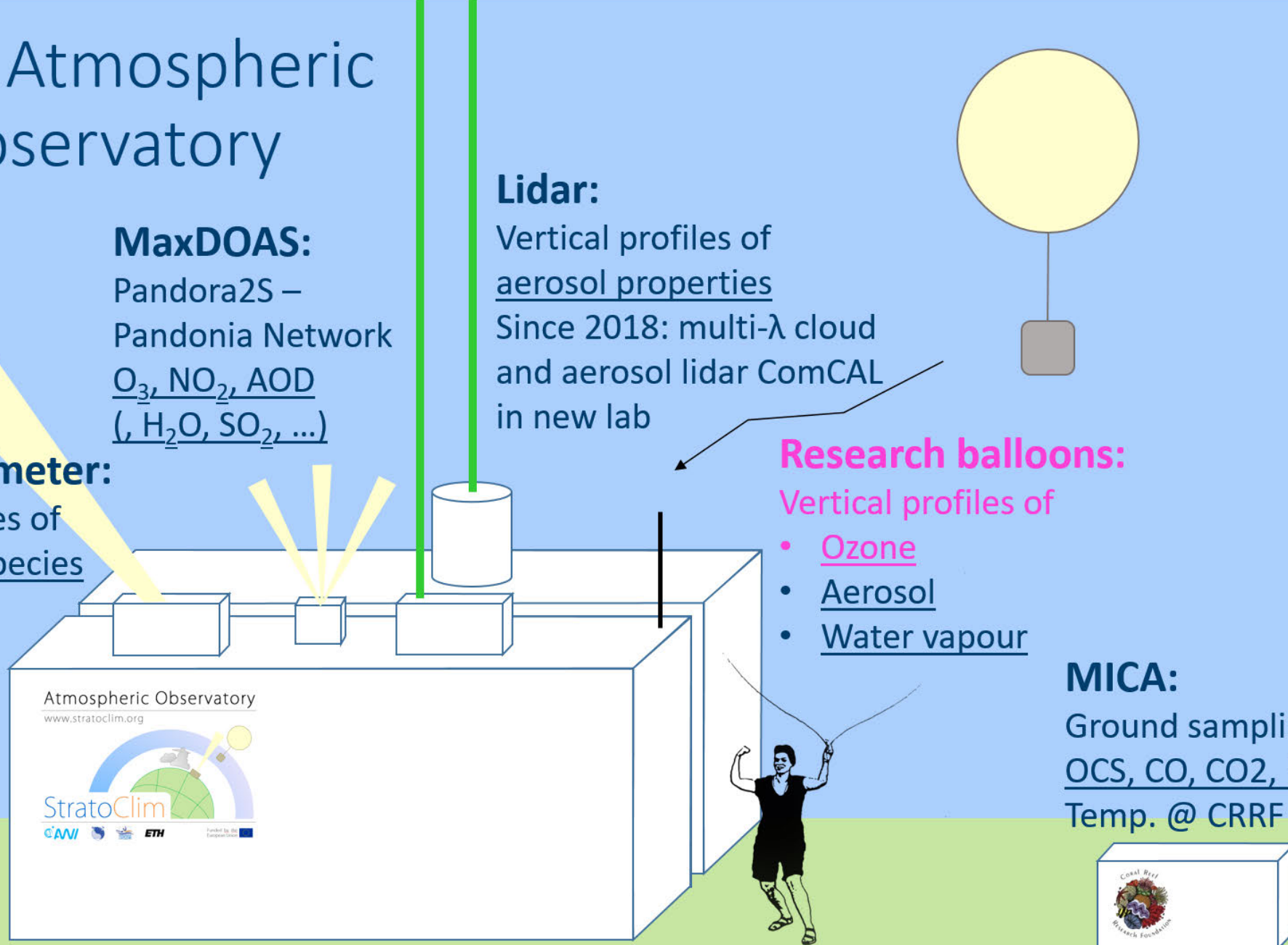
Research balloons:

Vertical profiles of

- Ozone
- Aerosol
- Water vapour

MICA:

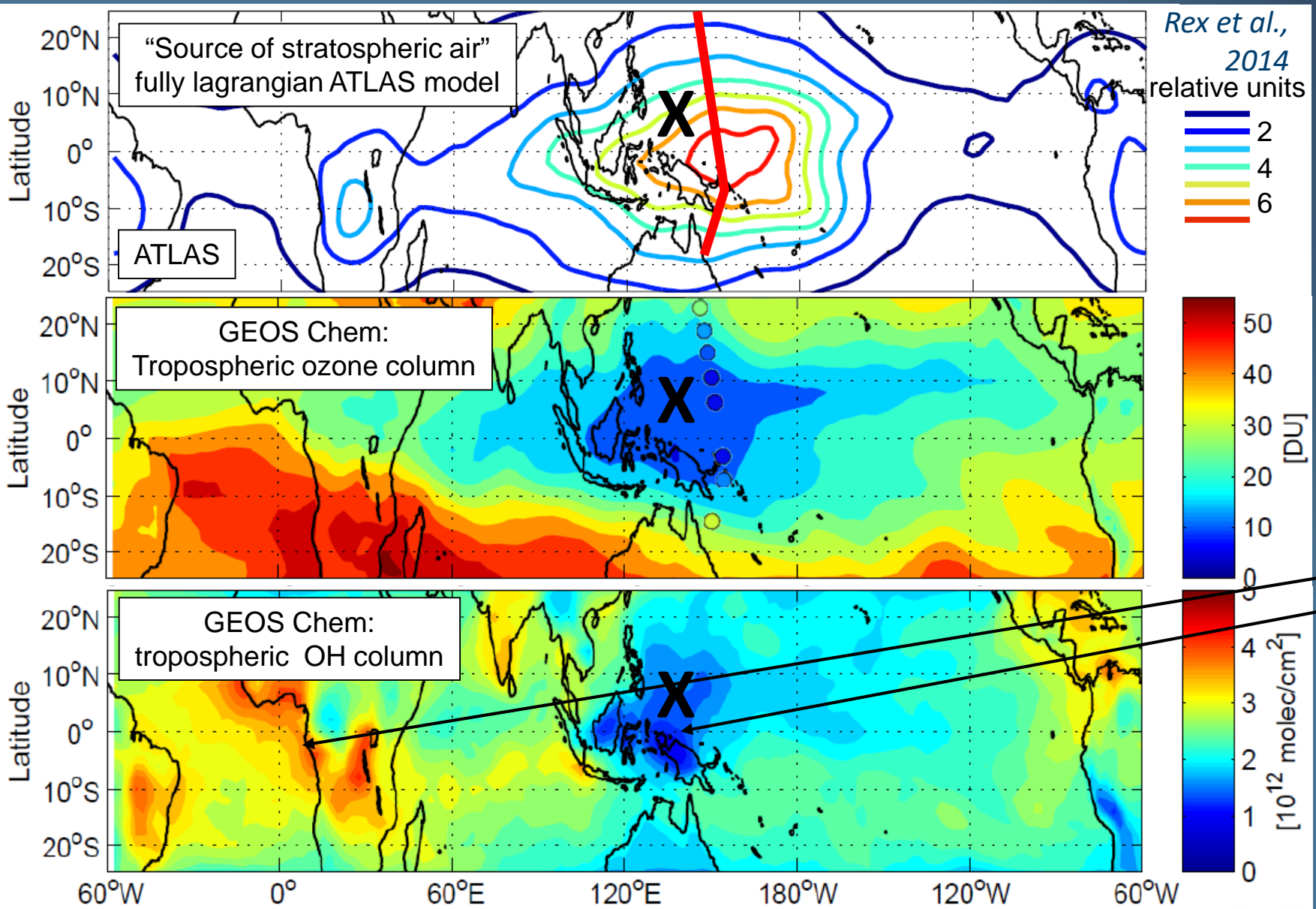
Ground sampling of
OCS, CO, CO₂, H₂O
Temp. @ CRRF site



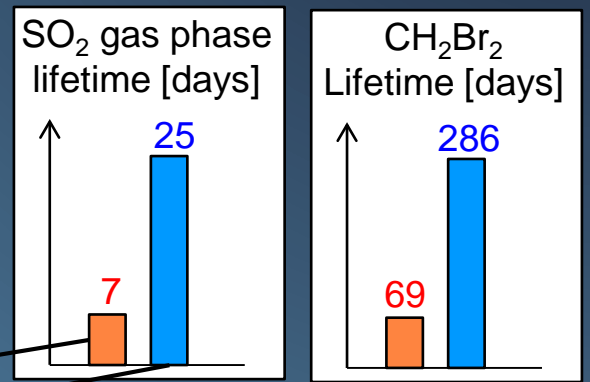
(I) Motivation - TransBrom



X Palau



Density distribution function of the horizontal positions of the trajectories between boundary layer and Lagrangian Cold Points; **red thick line**: TransBrom Cruise 2009; filled circles: from ozonesonde measurements

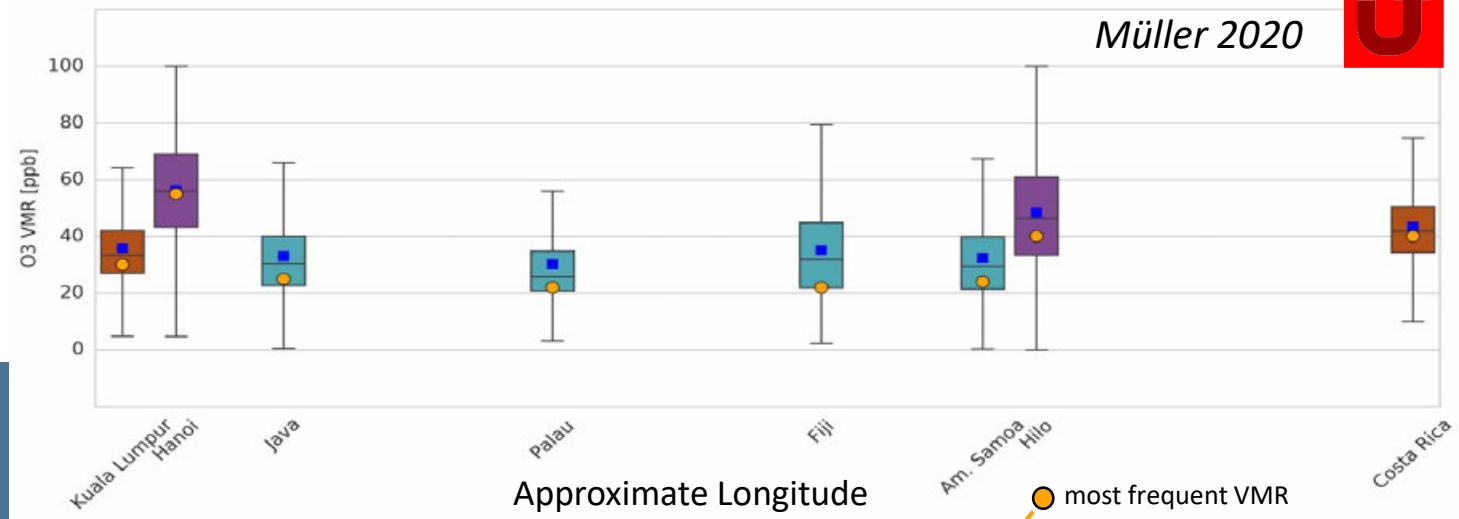


Lifetime comparison for tropical Atlantic and West Pacific (values for mid-troposphere- 500 hPa at the equator for typical conditions)

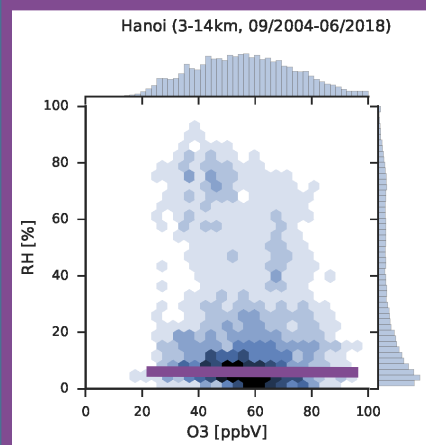
(I) O₃/RH Comparison with SHADOZ



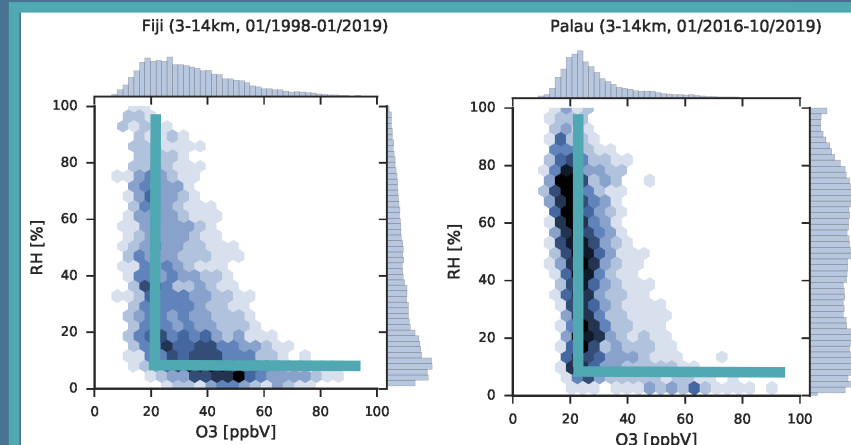
Müller 2020



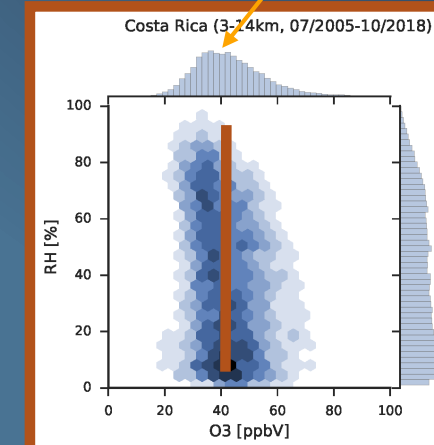
Analyses of 7 selected SHADOZ stations reveals **3 types of free-tropospheric O₃/RH distributions** (see Müller 2020). Seasonal distributions (not shown here) highlight uniqueness of Palau



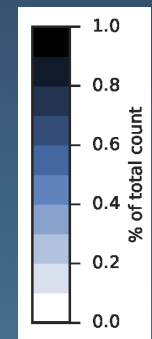
Predominantly dry air over a wider range of O₃ VMR

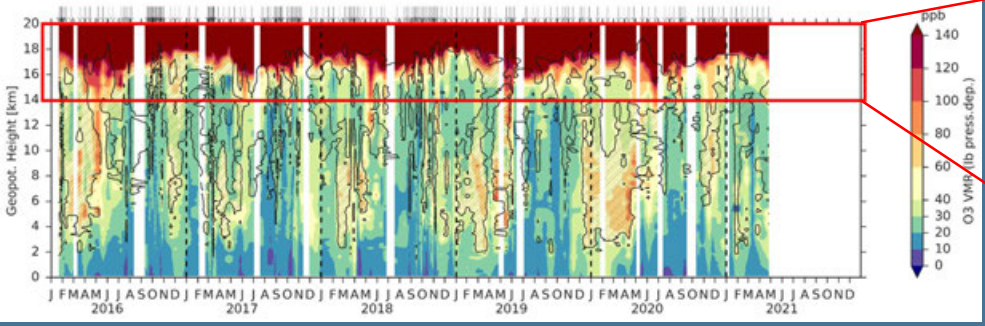


„L“-shaped: low O₃ over whole RH range + tail towards higher O₃ corresponding to low RH

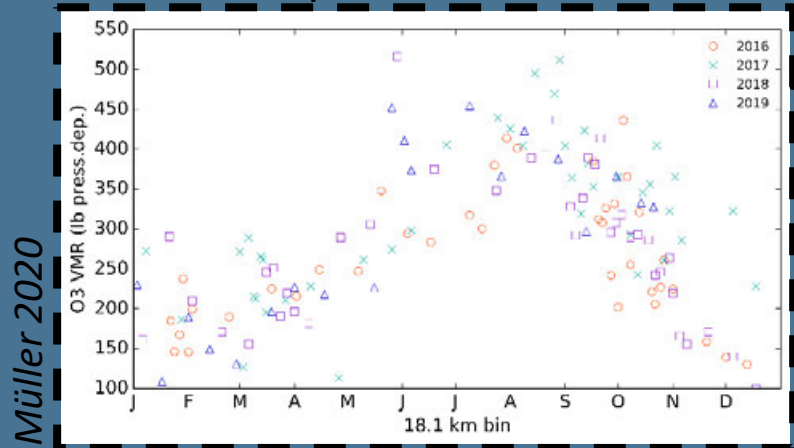


Gaussian distribution of O₃ + evenly distributed RH

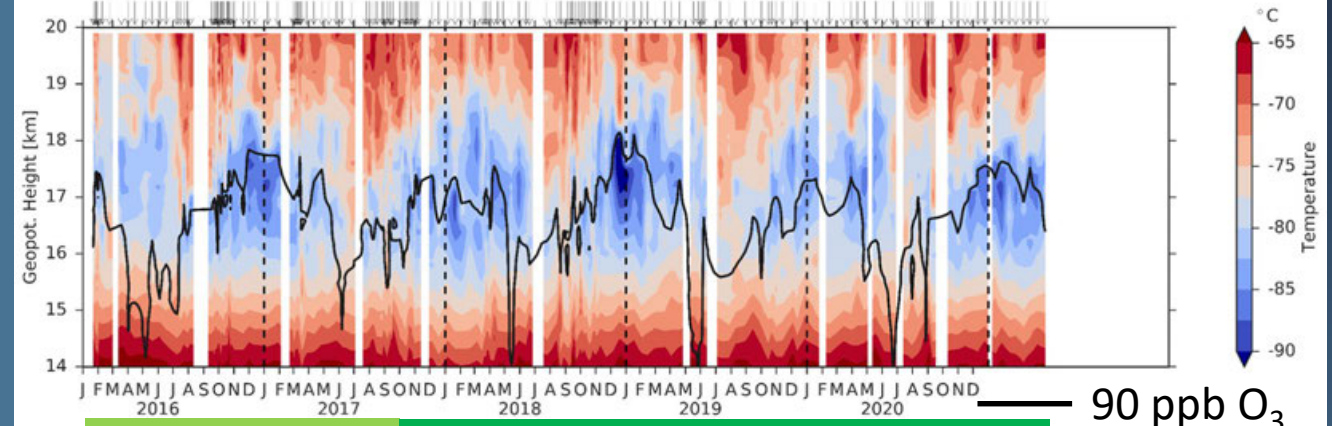
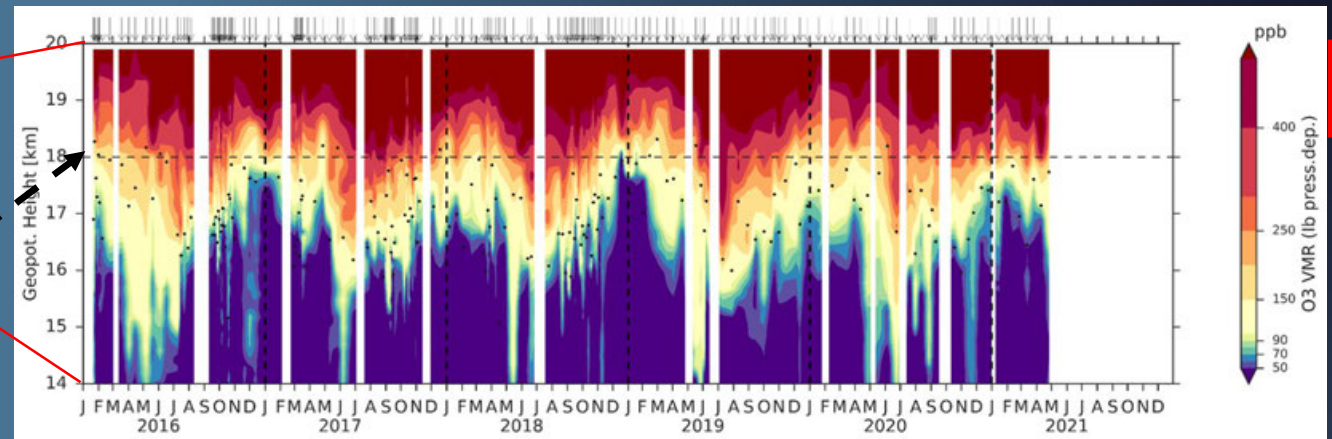




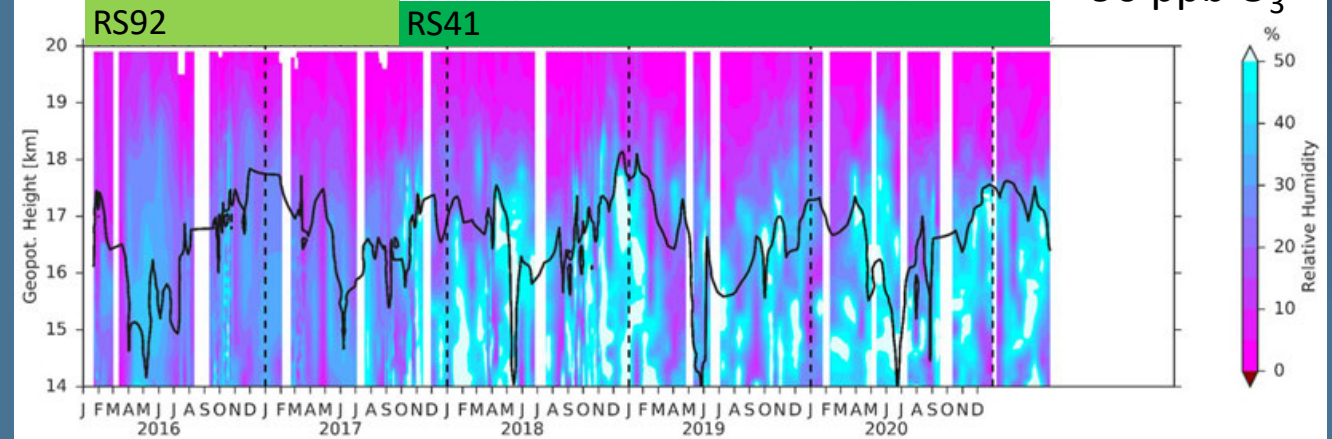
300m Layer around 18.15km



Müller 2020



90 ppb O₃

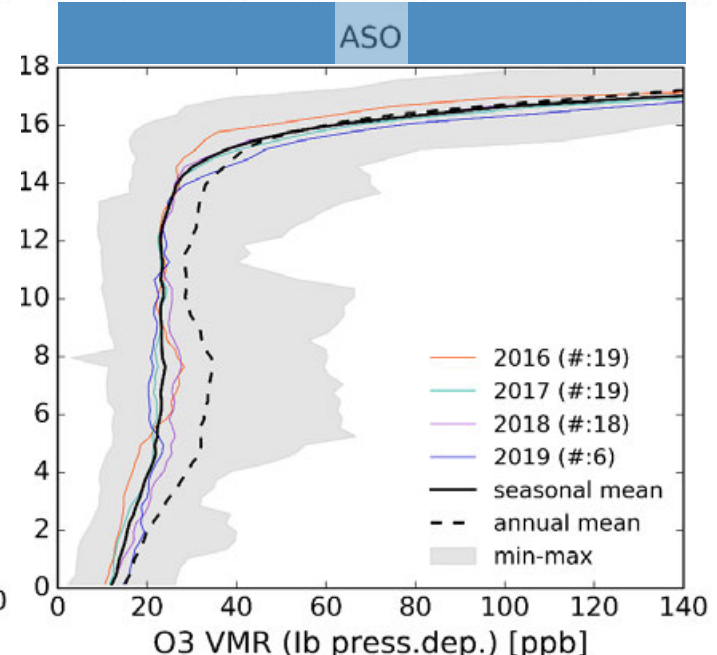
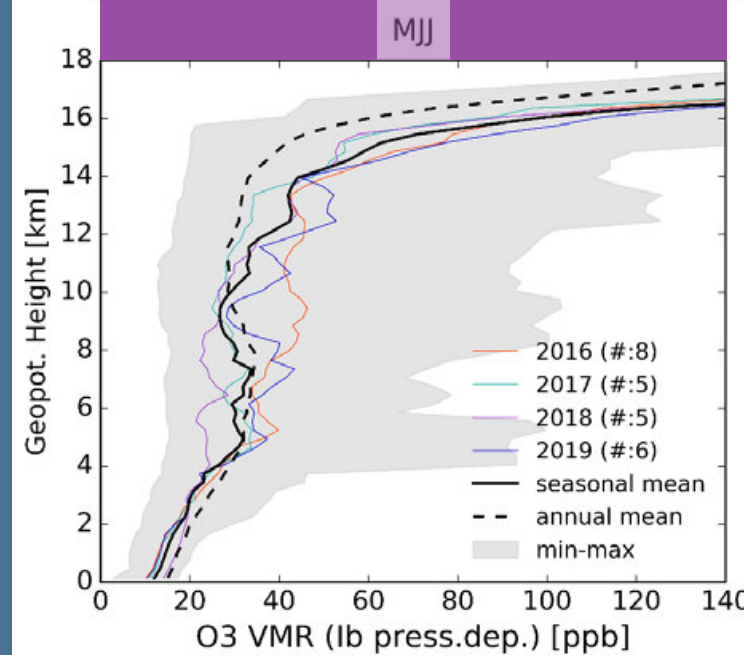
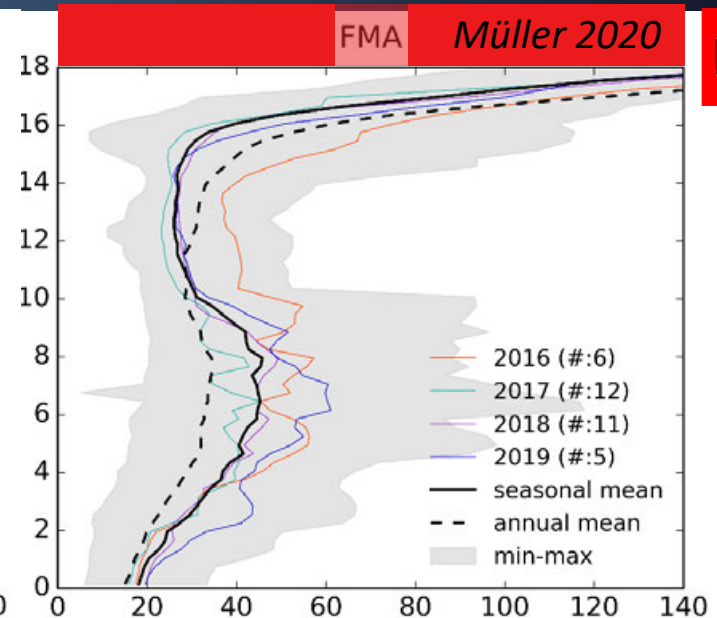
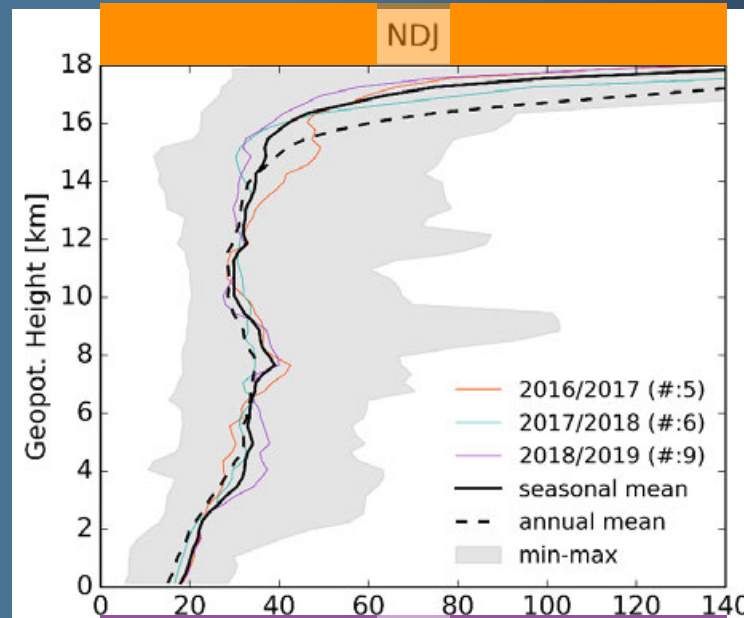


- Greatest amplitude of the TTL O₃ cycle occurs around 18 km
- Feedback during **El Nino 2016** in the UT: suppressed convection with less uplift of ozone-poor air from the ground or wash-out of O₃ precursors

(II) Seasons

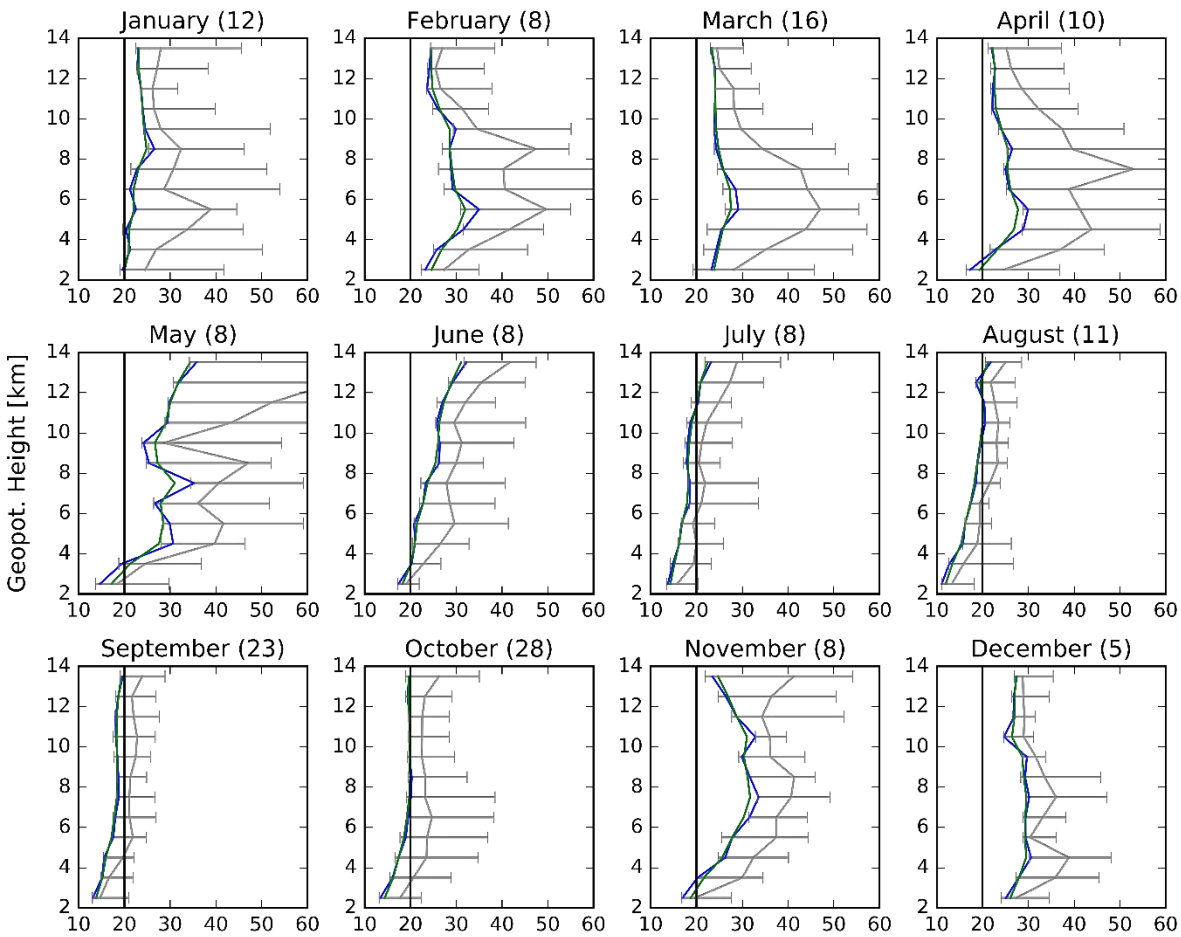


- November-January: NDJ, February-April: FMA, ..., chosen due to similar profile shapes
- Low O₃ (<25ppb) and enhanced mid-trop. O₃ (>50 ppb) observed in all seasons
- greatest anomalies from the annual mean: **FMA** & **ASO**
- Air masses deviating from a low O₃ background signal occur as **filaments or layers**, predominantly in the **5-10 km** layer, disguised in the averaged belly of the ,S'





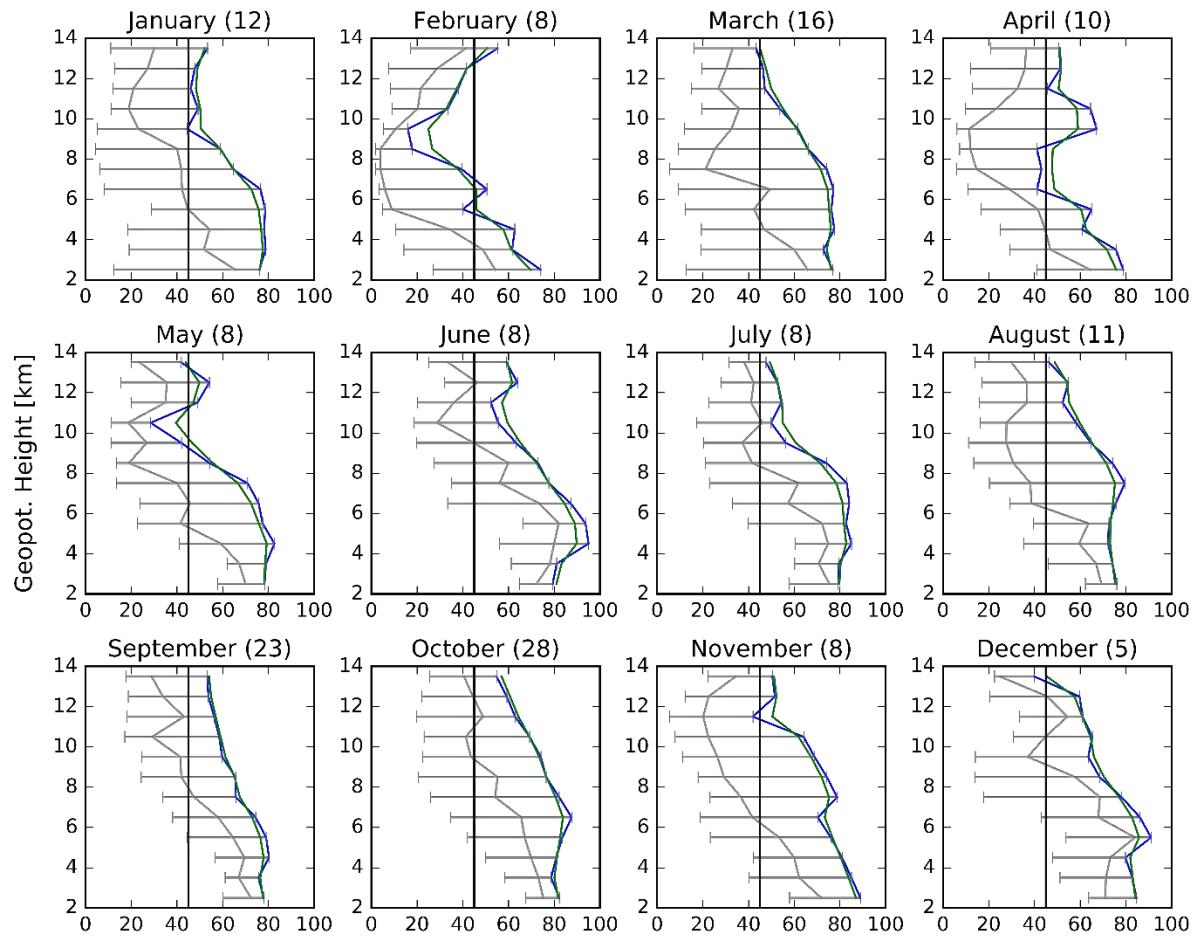
Monthly Quantiles



O3 VMR [ppb]

- 20%
- 20% smoothed
- Median

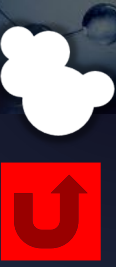
Monthly Quantiles



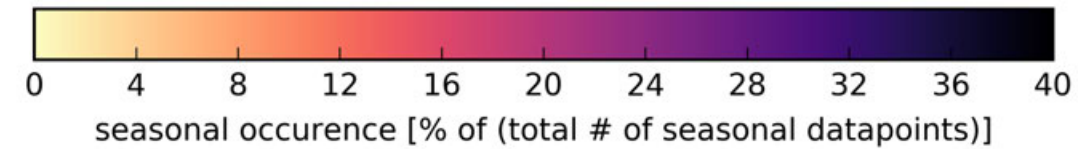
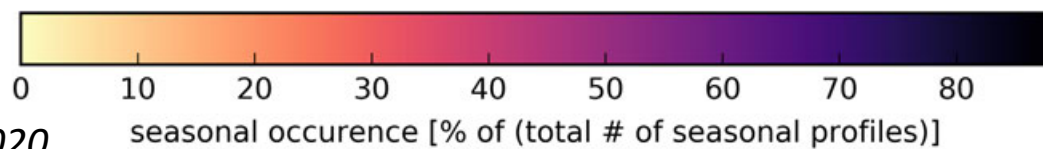
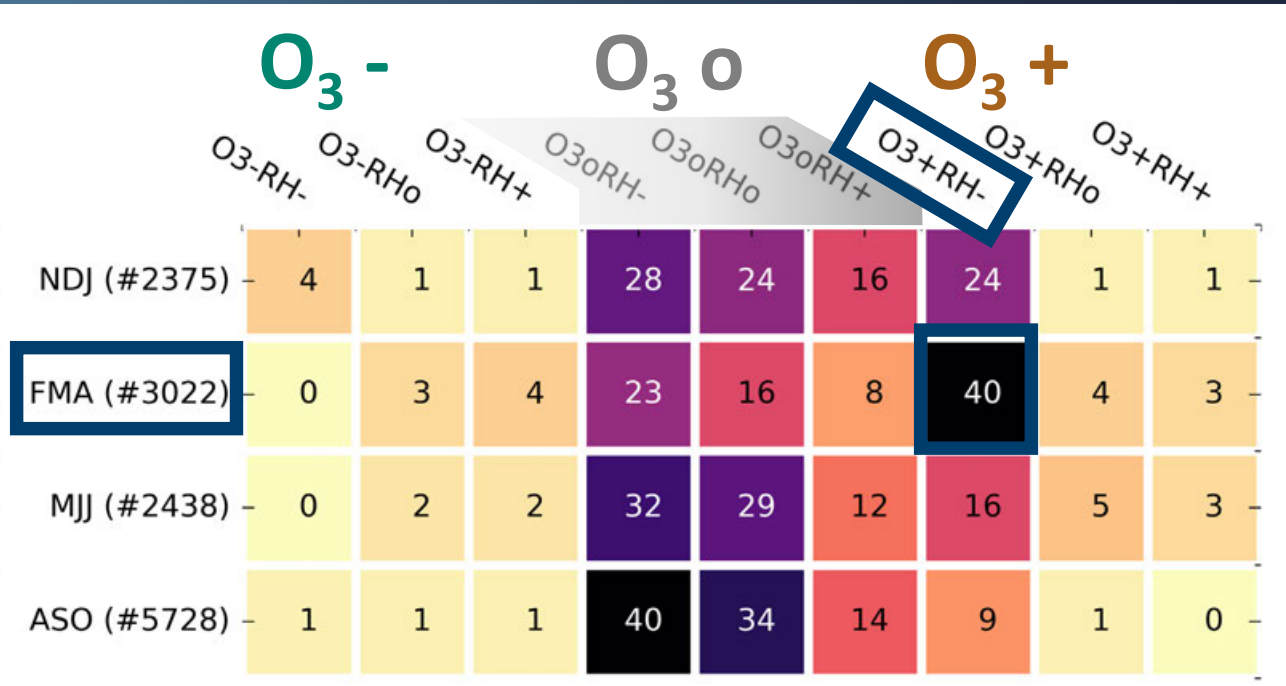
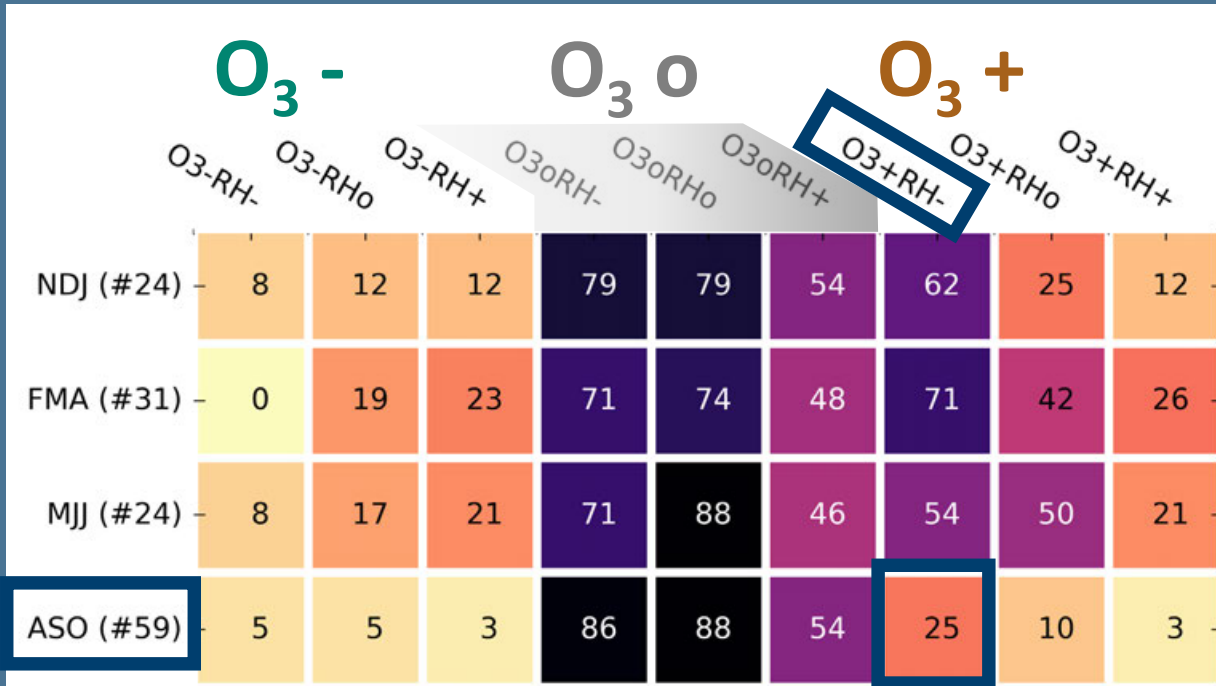
RH [%]

- 83.3%
- 83.3% smoothed
- Median

(III) Seasonal Occurrence of O₃/RH groups



Heatmaps for the seasonal occurrence of air masses for all nine anomaly groups, full time series, between 5-10 km altitude



Müller 2020

Example: O₃+RH- air masses occur in 25% of all ASO profiles.

Example: O₃+RH- air masses make up for 40% of all datapoints observed in FMA.