

1 **Assimilating summer sea-ice thickness observations**
2 **improves Arctic sea-ice forecast**

3 **Ruizhe Song^{1,2,3,4}, Longjiang Mu², Svetlana N. Loza^{3,5}, Frank Kauker³,**
4 **Xianyao Chen^{1,2}**

5 ¹Frontier Science Center for Deep Ocean Multispheres and Earth System and Physical Oceanography
6 Laboratory, Ocean University of China, Qingdao, China

7 ²Laoshan Laboratory, Qingdao, China

8 ³Alfred Wegener Institute, Helmholtz Centre for Polar and Marine Research, Bremerhaven, Germany

9 ⁴Academy of the Future Ocean, Ocean University of China, Qingdao, China

10 ⁵Shirshov Institute of Oceanology, Russian Academy of Sciences, Moscow, Russia

11 **Key Points:**

- 12 • Assimilating summer CryoSat-2 sea-ice thickness (SIT) observations makes more
13 skillful Arctic ice-edge forecasts on multiple time scales.
14 • The long-term SIT forecasts improve with the assimilation of summer CryoSat-
15 2 SIT observations.
16 • Further refinement is needed for summer CryoSat-2 SIT observations.

Abstract

Accurate Arctic sea-ice forecasting for the melt season is still a major challenge because of the lack of reliable pan-Arctic summer sea-ice thickness (SIT) data. A new summer CryoSat-2 SIT observation dataset based on an artificial intelligence algorithm may alleviate this situation. We assess the impact of this new dataset on the initialization of sea-ice forecasts in the melt seasons of 2015 and 2016 in a coupled sea ice-ocean model with data assimilation. We find that the assimilation of the summer CryoSat-2 SIT observations can reduce the summer ice-edge forecast error. Further, adding SIT observations to an established forecast system with sea-ice concentration assimilation leads to more realistic short-term summer ice-edge forecasts in the Arctic Pacific sector. The long-term Arctic-wide SIT prediction is also improved. In spite of remaining uncertainties, summer CryoSat-2 SIT observations have the potential to improve Arctic sea-ice forecast on multiple time scales.

Plain Language Summary

Arctic sea ice is rapidly declining due to global warming, especially in summer. Accurate sea-ice forecasting is important to understand the potential influence of these changes and devise effective responses. The performance of sea-ice forecasts highly depends on the accuracy of the initial sea-ice states. So refining the initial conditions of sea-ice forecasts with satellite observations is a common way to reduce forecast errors. However, obtaining reliable summer pan-Arctic satellite sea-ice thickness (SIT) data is challenging due to complex ice-surface conditions in summer. A new artificial-intelligence-based summer SIT satellite data product may improve initial SIT states. We integrate this dataset into a sea-ice forecast system to evaluate its impact on forecast skill. We find that the new summer satellite SIT data can reduce short-term ice-edge location forecast errors and benefit long-term SIT forecasts.

1 Introduction

Arctic sea ice is declining at unprecedented speed (Rothrock et al., 1999; Comiso et al., 2008; Kwok & Rothrock, 2009; Stroeve et al., 2012), which would pose challenges to climatic and ecological stakeholders (Landrum & Holland, 2020). The Arctic Passage, opening up with the gradually melting summer sea ice, calls for accurate Arctic sea-ice prediction from daily to seasonal scales for safe navigation (Jung et al., 2016).

Accurate initialization of sea-ice state is vital for predicting Arctic sea ice (e.g., Blanchard-Wrigglesworth et al., 2011; Guemas et al., 2016; Xie et al., 2016; Dirkson et al., 2017; Bushuk et al., 2022). The assimilation of sea-ice concentration (SIC) has improved the short-term sea-ice forecasts greatly as documented in the literature, and is now widely used at forecasting centers (e.g., Hebert et al., 2015; Lemieux et al., 2015). Sea-ice thickness (SIT) persists longer, therefore assimilation of SIT raises long-term sea-ice forecast skills even higher (Day, Hawkins, & Tietsche, 2014; Shu et al., 2021; Mu et al., 2022).

However, the potential impacts of summer SIT observations on sea-ice forecasts are not examined comprehensively yet due to a lack of data. An effective retrieval method for the remotely sensed SIT from May to September was desired (Laxon et al., 2013; Ricker et al., 2014). The complex summer ice-surface conditions restrict the application of classical algorithms designed for winter conditions. For instance, melt ponds which occupy a huge fraction of the sea-ice surface in the melt seasons (Maykut et al., 1992) complicate the classification algorithms (Lee et al., 2018; Tilling et al., 2019) and introduce large uncertainties due to increased moisture in the snow (Drinkwater, 1991). On the other hand, in-situ Arctic SIT observations are rather scarce and localized, which can be hardly used for assimilation due to their limited spatial representation within a relatively large model grid cell.

64 In a recent study, Dawson et al. (2022) presented the first estimate of pan-Arctic summer
65 sea-ice freeboard from radar altimeter by using a 1D convolutional neural network (CNN)
66 to distinguish ice leads from melt ponds. Landy et al. (2022) converted summer CryoSat-2
67 radar freeboard to SIT and applied further corrections. The spring predictability barrier of
68 the Arctic sea ice (e.g., Day, Tietsche, & Hawkins, 2014; Bushuk et al., 2017) suggests that
69 sea-ice forecast should benefit from the initialization with SIT in the melt season (Bushuk et
70 al., 2020). Therefore, it presents an opportunity to explore the extent to which the summer
71 SIT observation could improve the real-time forecast skill. Min et al. (2023) demonstrated
72 that assimilation of summer SIT corrects the overestimation in the Combined Model and
73 Satellite Thickness (CMST; Mu et al., 2018b) product. Y.-F. Zhang et al. (2023) found
74 that the assimilation of May to August CryoSat-2 SIT anomalies improves local SIC and
75 sea-ice extent (SIE) forecasts in September. However, the influence of assimilating summer
76 CryoSat-2 SIT observations on short-term sea-ice forecast in summer and on long-term
77 forecast extending beyond September still needs further investigation.

78 In this study, we focus on the impact of summer SIT observations on the daily and
79 seasonal forecast skills of a sea-ice prediction modelling system. In particular, we perform
80 a series of short- and long-term ensemble sea-ice forecasts where the sea ice-ocean initial
81 state is constrained by the summer CryoSat-2 SIT or where these data are not used. The
82 benefits and challenges of using these new SIT data are evaluated and critically discussed
83 using independent sea-ice data.

84 **2 Data and Methods**

85 **2.1 The coupled sea ice-ocean model**

86 We use a regional coupled sea ice-ocean model driven by atmospheric forecasts to con-
87 figure the sea ice-ocean forecast system. The model is based on the Massachusetts Institute
88 of Technology general circulation model (MITgcm; Marshall et al., 1997) and covers the
89 pan-Arctic region with a horizontal resolution of around 18 km as in Losch et al. (2010).
90 The sea-ice model uses a viscous-plastic rheology (Hibler III, 1979; J. Zhang & Hibler III,
91 1997) and a zero-layer thermodynamic formulation without heat capacity (Semtner, 1976;
92 Parkinson & Washington, 1979). The readers are referred to Losch et al. (2010) and Nguyen
93 et al. (2011) for more details on the model.

94 **2.2 Data assimilation and forecast**

95 The summer data assimilation system is initialized from restart files generated by CMST
96 (Mu et al., 2018b) simulation with 11 ensemble members. CMST combines model physics
97 with information from remote-sensed SIT and SIC observations. It successfully reproduces
98 the spatio-temporal sea-ice variations (Mu et al., 2018b). In this study, the summer data
99 assimilation and forecast strategy follows Mu et al. (2019). All observations and corre-
100 sponding uncertainties are interpolated onto the 18-km model grid for assimilation. After
101 the assimilation of sea-ice observations using a Local Error Subspace Transform Kalman
102 Filter (Nerger et al., 2012) coded within the Parallel Data Assimilation Framework (Nerger
103 et al., 2005), the ensemble sea-ice forecasts start from the new analyses and are integrated
104 forced by the atmospheric forecasts (cf. Section 2.3). More details on the data assimilation
105 and forecast system are given in Supporting Information.

106 The 80-km-resolution CryoSat-2 summer SIT data set is derived from local variations
107 in the CryoSat-2 radar echo response using a deep learning method (Dawson et al., 2022;
108 Landy et al., 2022). This is the first estimate of pan-Arctic summer SIT from satellite
109 observations. The summer SIT is assimilated into the system on a daily basis using the
110 observations linearly interpolated between two biweekly (twice per month) records. How-
111 ever, the roughness-induced electromagnetic range bias on the heavily-deformed ice in the
112 coast regions leads to significant SIT underestimate north of the CAA and Greenland in late

113 summer (Landy et al., 2022). Practically we set the observation uncertainties higher than
114 the original values over thick ice regions, while still using the provided errors over thin ice
115 regions (Supporting Information). The SIC data used in the assimilation are computed at
116 the French Research Institute for Exploitation of the Sea (IFREMER) based on the 85-GHz
117 SSM/I and SSM/IS channels with a resolution of 12.5 km (Kaleschke et al., 2001; Spreen
118 et al., 2008; Kern et al., 2010). The uncertainty of the SIC observation is set to a constant
119 value of 0.25 following Yang, Losa, Losch, Jung, and Nerger (2015).

120 The short-term ensemble assimilation and forecast experiments are driven by the 174-
121 hour atmospheric ensemble forecasts from the United Kingdom Met Office (UKMO) En-
122 semble Prediction System (EPS; Bowler et al., 2008). For the long-term prediction, the
123 ensemble members are driven by deterministic atmospheric forcing (single member). The
124 atmospheric forecasts from the NCEP Climate Forecast System Version 2 (CFSv2; Saha et
125 al., 2014) are used for the 9-month long-term forecasts, while the ECMWF Reanalysis v5
126 (ERA5; Hersbach et al., 2020) is used as the atmospheric forcing during the data assimila-
127 tion to minimize the potential error caused by deviations of atmospheric forcing during this
128 period.

129 2.3 Experiment design

130 In order to investigate the potential impact of the CryoSat-2 summer SIT on sea-ice
131 forecasts, this study designs both short-term (7 days) and long-term (270 days) forecasts
132 (Table. 1). These experiments are conducted over different months. The short-term experi-
133 ments in 2015, which cover the melt season, start from the CMST restart files on May 1, May
134 31, June 30, July 30, and August 29, respectively. Each forecast experiment lasts for 30 days
135 and on each day a 7-day sea-ice forecast is run using the atmospheric forcing from the daily
136 UKMO ensemble forecasts. No data assimilation is applied in the control run of the short-
137 term forecasts (Short-CTRL). The Short-SIT experiments assimilate only the CryoSat-2
138 summer SIT data, and the Short-SIC experiments assimilate only the SSMI/SSMIS SIC
139 data, while both data sets are assimilated in the Short-SICSIT experiments. For the 2016
140 experiments, only the start dates are changed to match the available restart files from CMST
141 (Table. 1).

142 The long-term forecast experiments are designed to diagnose the persistence of the
143 assimilated CryoSat-2 summer SIT over the months from the melt season to the freezing
144 season. The Long-SIT, Long-SIC and Long-SICSIT experiments with data assimilation
145 start each summer month from CMST restart files. Unlike the incremental analysis update
146 approach, the state vector is updated each day directly in the next 15 days to assimilate
147 observations. Over that period, ERA5 atmospheric reanalysis forcing is used. Then, the
148 270-day sea-ice forecasts start from the sea-ice analysis restart files and are forced by the
149 CFSv2 operational atmospheric forecasts. No data assimilation is performed in the Long-
150 CTRL experiments. The forecast start dates are listed in Table 1.

151 2.4 Verification

152 Airborne electromagnetic SIT observations north of Greenland from AWI IceBird cam-
153 paigns in July and August 2016 are employed for comparison with the assimilation results.
154 Locations of these observations are indicated in Figure S1. The integrated ice-edge error
155 (IIEE; Goessling et al., 2016) is used to quantify the skill of the short-term ice-edge fore-
156 casts. It measures the discrepancy between the forecasted and observed SIE. The reference
157 observation used in this study is the 25-km-resolution NOAA/NSIDC Climate Data Record
158 (CDR) of Passive Microwave Sea Ice Concentration Version 4 (Meier et al., 2021).

159 To validate the skill of the long-term sea-ice forecast, we compute the IIEE and the
160 RMSD of SIT against various other products and in-situ observations. The IIEE is still com-
161 puted using the NOAA/NSIDC CDR data. The RMSDs of SIT are computed with respect

Table 1. Summary of forecast experiments design. The number in the parenthesis represents the size of atmospheric forcing ensemble. Short: short-term forecast. Long: long-term forecast. SIC: sea-ice concentration. SIT: sea-ice thickness.

Experiment	Assimilated data	Forecast duration (days)	Atmospheric forcing during assimilation	Atmospheric forcing during forecast	Forecast start date
Short-CTRL	/	7	UKMO (11)	UKMO (11)	Daily forecast starting from 01 May 2015,
Short-SIT	CryoSat-2 SIT	7	UKMO (11)	UKMO (11)	31 May 2015, 30 Jun 2015, 30 Jul 2015,
Short-SIC	SSMI/SSMIS SIC	7	UKMO (11)	UKMO (11)	29 Aug 2015, 25 Apr 2016, 25 May 2016,
Short-SICSIT	SSMI/SSMIS SIC and CryoSat-2 SIT	7	UKMO (11)	UKMO (11)	24 Jun 2016, 24 Jul 2016, 23 Aug 2016.
Long-CTRL	/	270	ERA5 (1)	CFSv2 (1)	16 May 2015, 15 Jun 2015,
Long-SIT	CryoSat-2 SIT	270	ERA5 (1)	CFSv2 (1)	15 Jul 2015, 14 Aug 2015, 13 Sep 2015,
Long-SIC	SSMI/SSMIS SIC	270	ERA5 (1)	CFSv2 (1)	10 May 2016, 09 Jun 2016,
Long-SICSIT	SSMI/SSMIS SIC and CryoSat-2 SIT	270	ERA5 (1)	CFSv2 (1)	09 Jul 2016, 08 Aug 2016, 07 Sep 2016.

162 to the 25-km-resolution CS2SMOS products (Ricker et al., 2017) when they are available be-
 163 tween October and the following April. Both NOAA/NSIDC CDR and CS2SMOS data are
 164 interpolated onto the 18-km grid to calculate the IIEE and RMSD. Note that CS2SMOS is a
 165 merged product using winter Cryosat-2 and Soil Moisture Ocean Salinity (SMOS) SIT. The
 166 SIT observations derived from upward-looking sonar moorings maintained by the Beaufort
 167 Gyre Exploration Program (BGEP) are used for the forecast evaluation. The three moorings
 168 BGEP-A, BGEP-B, and BGEP-D, which provide year-round sea-ice draft observations, are
 169 located at (75.0°N, 150.0°W), (78.0°N, 150.0°W) and (74.0°N, 140.0°W), respectively (Figure
 170 S1). The draft is converted to SIT by multiplying it by a constant factor of 1.1 as in Nguyen
 171 et al. (2011).

172 3 Result

173 3.1 Short-term ice-edge forecast

174 SIT from CryoSat-2 and the short-term experiments in 2015 is shown in Figure 1.
 175 The spatially averaged SIT differences between Short-SIT and Short-CTRL from May to
 176 September 2015 are 0.10 m, -0.06 m, -0.37 m, -0.37 m and -0.39 m, respectively. Overall,
 177 the SIT differences are smallest in May and June, when the assimilation of the summer
 178 CryoSat-2 observations reduces the SIT in the Pacific sector and increases it in the Atlantic
 179 sector (regions shown in Figure S1). Along with higher uncertainties in the CryoSat-2 SIT
 180 observations due to strong ice melting in July, August and September, a remarkable SIT
 181 reduction over the multi-year ice regions (regions shown in Figure S2) is found. SIT is also
 182 reduced in most of the marginal ice zones, especially in the Beaufort Sea and the Chukchi
 183 Sea. In our experiment, SIC assimilation, however, has only limited impact on SIT near
 184 the ice edge due to reliable restart states from CMST system that already has assimilated
 185 SIC observations. The absolute spatially averaged SIT differences between Short-SIC and
 186 Short-CTRL are minor, within 0.04 m. In the sea ice interior region with SIC close to 1.0
 187 far from the ice edge, SIC assimilation can hardly further improve the SIT there by means
 188 of the covariance matrix, due to a narrow SIC ensemble spread. Similar results are also
 189 found in 2016 (Figure S3).

190 Assimilating summer CryoSat-2 SIT in Short-SIT gives rise to a more reasonable SIT
 191 probability density distribution along the trajectories of the IceBird campaigns north of
 192 Greenland (Figure S4), particularly for the modal SIT. Overestimation in CMST as indi-
 193 cated by Short-CTRL is significantly reduced. The median SIT difference against IceBird
 194 observations is mitigated in Short-SIT (-0.42 m), while it is -0.71 m and 0.98 m for CryoSat-2
 195 and Short-SIC, respectively. Short-SIT removes ice thicker than 3 m, resulting in a lower
 196 median than IceBird observations. Compared to observations from BGEP moorings (Fig-
 197 ure S5), the assimilation of summer CryoSat-2 SIT leads to a further underestimated SIT
 198 particularly in May, but corrects the SIT overestimation in late summer.

199 SIT assimilation has an important impact on SIC simulations through the physical
 200 connection between thickness and concentration over thin ice areas (Xie et al., 2016; Mignac
 201 et al., 2022). Short-term forecast of ice edge, defined as the 15% SIC isoline, can be strongly
 202 influenced by SIT assimilation. Figure 2 shows the IIEE difference in the Pacific sector
 203 and Atlantic sector (regions shown in Figure S1). IIEE in each forecast experiment is
 204 given in Figure S6. The observed SIC used as the reference for the IIEE calculation is
 205 the NOAA/NSIDC SIC CDR. The difference in the ice-edge position between forecasts and
 206 observations in 2015 and 2016 is displayed in Figure S7 and Figure S8.

207 The impact of CryoSat-2 SIT assimilation on ice-edge forecasts varies with time and
 208 region. Compared to Short-CTRL, IIEE in Short-SIT is strongly reduced in most times and
 209 both sectors (Figure 2). The ice-edge position in the forecasts is consistently overestimated
 210 in Short-CTRL. Assimilation of the summer SIT reduces the SIT of the forecasts near the

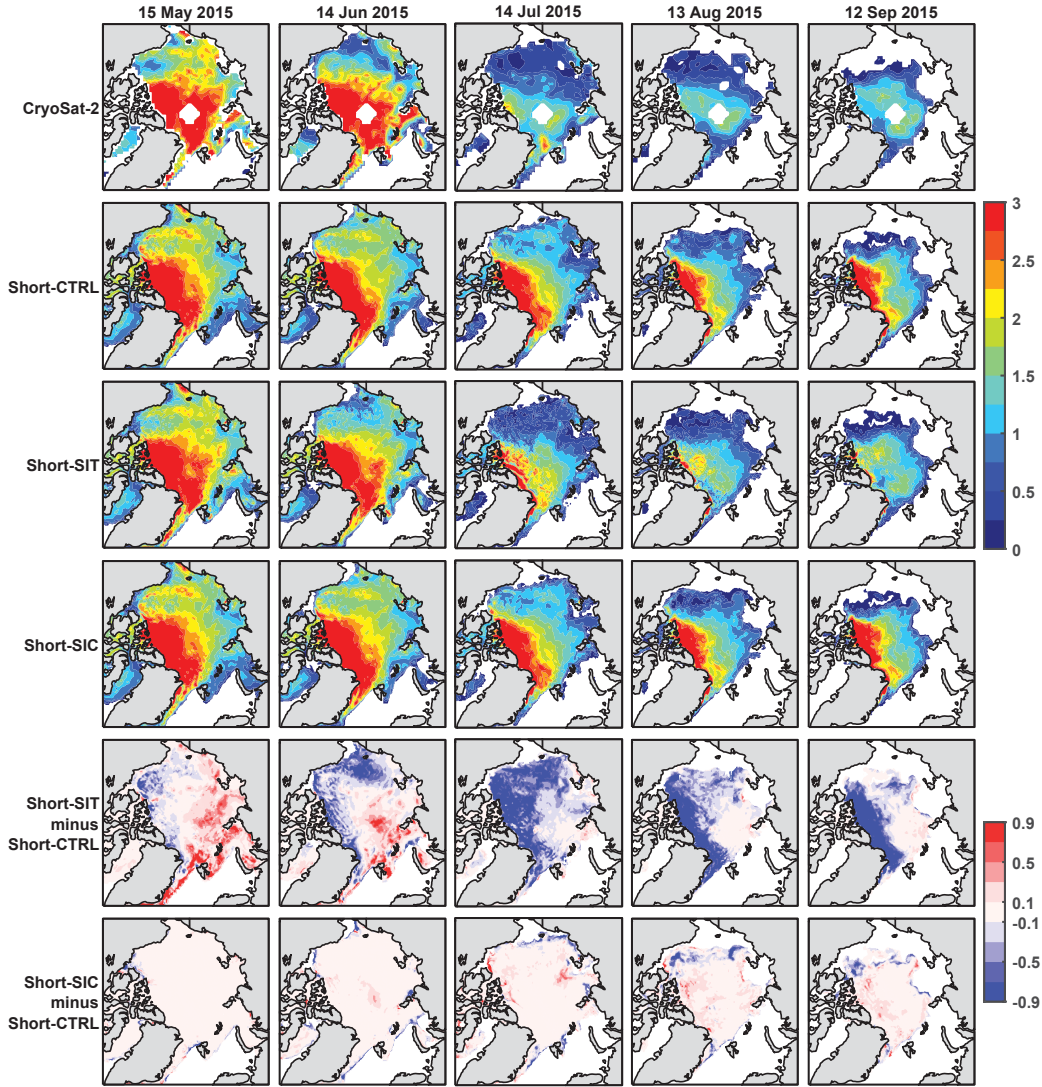


Figure 1. CryoSat-2 SIT (m) used for assimilation, SIT analysis from short-term experiments, and their differences between experiments on the 15th day from the model start date in 2015.

211 ice edge, resulting in a better agreement between the ice-edge forecasts and the ice-edge
 212 observations from the satellite compared with Short-CTRL (Figure S7 and Figure S8).

213 In the Pacific sector, only a slight improvement in IIEE is observed in May and June
 214 for Short-SIT compared to Short-CTRL (Figure 2). However, in July, especially in 2015,
 215 IIEE increases and the forecast skill degrades. This can be attributed to the fact that
 216 the melt-pond fraction starts to increase in June and reaches its maximum in July (Feng et al.,
 217 2022). In particular, 2015 was the peak year for observed melt-pond fraction in the Beaufort
 218 Sea between 2000-2021 (Xiong & Ren, 2023). The presence of excessive melt-pond fraction
 219 in this region may lead to more misclassification of ice leads and melt ponds in the CryoSat-2
 220 sea-ice freeboard retrieval using the CNN model, which affects the SIT analysis in the Pacific
 221 sector. Therefore, the underestimated SIT erroneously leads to a large ice-edge error in July
 222 of the Short-SIT experiments. This warrants further refinement of the artificial intelligence
 223 algorithm used for summer CryoSat-2 SIT retrieval. In late summer, the assimilation of
 224 CryoSat-2 SIT observations in Short-SIT leads to more skillful ice-edge forecasts, resulting

225 in a statistically significant average reduction in IIEE of about $2.1 \times 10^5 \text{ km}^2$. For example,
 226 the assimilation of SIT allows the model to predict an ice-free "cave" inside the Beaufort
 227 Sea in August 2015, while it is completely covered by sea ice in Short-CTRL and still with
 228 a connected strip of ice in Short-SIC (Figure S7). Furthermore, the ice-edge forecasts in
 229 the Atlantic sector are also improved for Short-SIT compared to Short-CTRL, especially in
 230 June (about $0.8 \times 10^5 \text{ km}^2$) and July (more than $0.9 \times 10^5 \text{ km}^2$).

231 We further investigate the influences of SIC assimilation together with summer SIT
 232 assimilation on the ice-edge forecasts, considering the more important role of SIC observa-
 233 tions on summer sea-ice forecasts as documented in the literature (e.g., Posey et al., 2015;
 234 Yang, Losa, Losch, Liu, et al., 2015). Forecasts from the Short-SICSIT experiments are also
 235 compared to the Short-SIC experiments, which performs SIC assimilation only.

236 In the Pacific sector, the additional SIT assimilation tends to yield more favorable ice-
 237 edge forecasts compared to Short-SIC (Figure 2). Similar to the IIEE differences between
 238 Short-SIT and Short-CTRL, the improvement in May and June between Short-SICSIT and
 239 Short-SIC is relatively small (only $3.0 \times 10^3 \text{ km}^2$ on average). In July, IIEE becomes smaller
 240 in 2015 but larger in 2016 relative to Short-SIC. In late summer, the analysis of summer
 241 SIT observations significantly reduces the IIEE, bringing the ice-edge forecasts closer to
 242 the observations. In the Atlantic Sector, Short-SICSIT tends to give rise to larger IIEE,
 243 resulting in more detrimental effects, particularly noticeable in May and June (Figure 2).
 244 Nevertheless, these mean IIEE differences are still in the range of $\pm 0.5 \times 10^5 \text{ km}^2$, which is
 245 much smaller than the changes between Short-SIT and Short-CTRL. In the Atlantic sector,
 246 Short-SIC is already close to the observations due to a reasonable CMST SIT estimate north
 247 of the Svalbard and Novaya Zemlya, so further improvements are rather limited.

248 Note that, as illustrated by the solid lines representing the mean IIEE differences in
 249 Figure 2, the impact of the summer CryoSat-2 SIT assimilation becomes more obvious with
 250 increasing lead time in Short-SICSIT. The improvements of Short-SICSIT relative to Short-
 251 SIC increase as forecast progressing, while the deteriorations of IIEE become smaller, with
 252 the exception of the June 2016 forecasts.

253 3.2 Long-term sea-ice forecast

254 The Long-SIT experiments with summer CryoSat-2 SIT assimilation provides signifi-
 255 cant benefits for ice-edge and thickness forecasts against Long-CTRL. Reductions in IIEE
 256 are found for the first 30 days in May, June and August of 2015 and 2016 (Figure S9a, b).
 257 For experiments also initialized with SIC constraints, the IIEEs are reduced for most of the
 258 time during these three months, but not overall (Figure 3a, b). For the forecast initialized
 259 in July, the CryoSat-2 SIT assimilation is generally detrimental and only effective for a few
 260 days due to the underestimated thickness uncertainties caused by melt ponds. In Septem-
 261 ber, improvements in ice-edge forecasts without SIC assimilation are seen for the first three
 262 weeks in 2015, and two weeks in 2016 (Figure S9a, b). The assimilation of SIC reduces such
 263 benefit (Figure 3a, b), which is not surprised.

264 With respect to the CS2SMOS SIT product, the predicted Arctic-wide thickness is also
 265 improved (Figure 3c, d; Figure S9c, d), except for the forecast starting in July 2016, which
 266 degrades after 140 days. The summer CryoSat-2 SIT mitigates the SIT overestimation in
 267 the Beaufort Sea in Long-CTRL and Long-SIC (not shown). The improvements are most
 268 pronounced in October, when the freezing season begins, and decrease exponentially with
 269 time until the forecast system falls into the control of the internal variability. This superior
 270 skill may even persist throughout the freezing season, similar to the previous findings on an
 271 optimal winter SIT initialization improving the predictive skill of summer sea ice (Blockley
 272 & Peterson, 2018). Consistent with the performance of the short-term forecasts in section
 273 3.1, the reduction of SIT RMSD in 2015 is more significant than that in 2016. When SIC
 274 assimilation is absent, the effect of SIT initialization on ice-edge forecasts is more pronounced

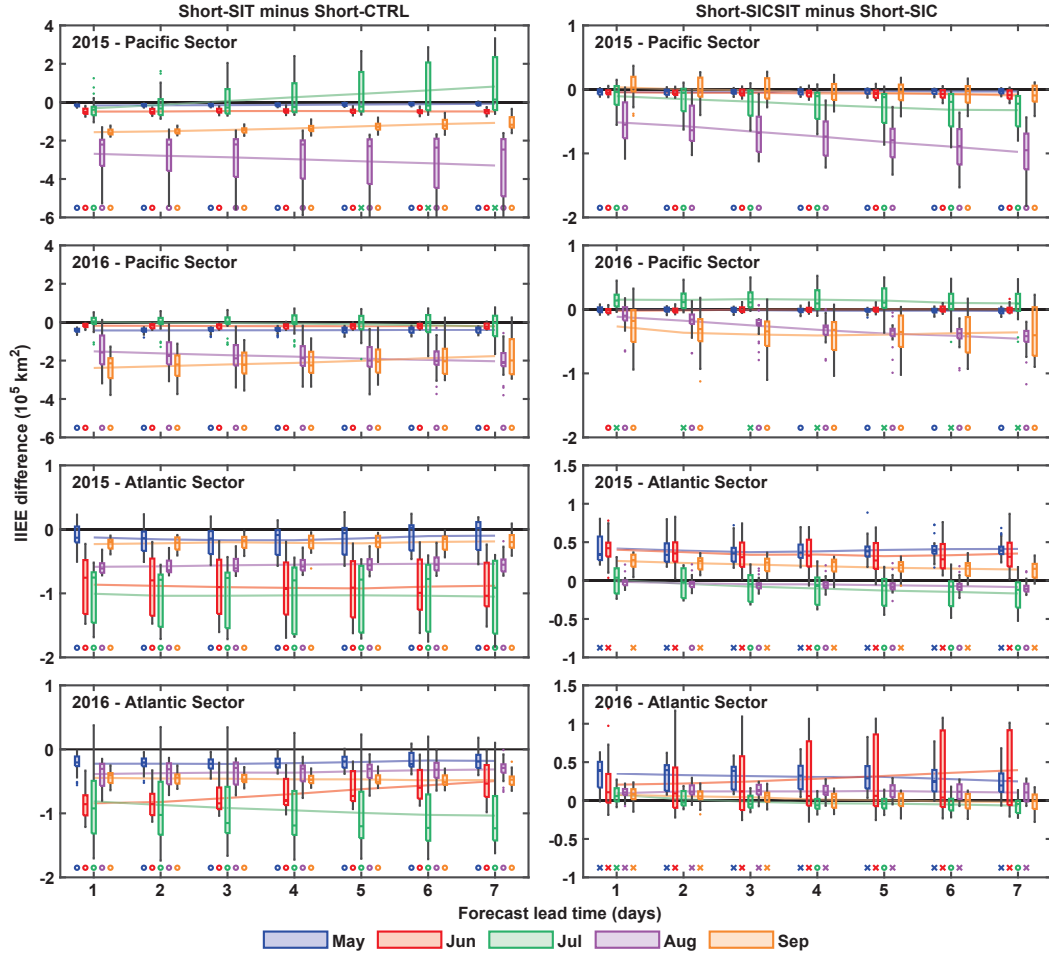


Figure 2. Box plot of the IIEE difference (10^5 km^2) between Short-SIT and Short-CTRL (left), together with that between Short-SICSIT and Short-SIC (right) in the 7-day sea-ice forecasts. The IIEE in the box plot is calculated after 7 days of assimilation when the summer CryoSat-2 SIT is fully assimilated. Box colors indicate different months. Box sizes indicate IIEE difference between the lower and upper quartiles. Outliers denote values more than 1.5 interquartile range from the top or bottom of the colored box. The outer edges of the black lines denote the minimum and maximum values that are not outliers. Solid-colored lines show the mean IIEE difference at each lead time. A positive value indicates an increase in IIEE, when SIT is assimilated, while a negative value indicates a decrease in the IIEE. Markers at the bottom of each panel indicate increases (cross) and decreases (circle) in IIEE that pass the Student’s T-test at the 95% confidence level. Negative values indicate better forecast skills. Note that different subfigures use different y-axis scales.

275 (Figure S9). However, the skill of the long-term SIT forecasts remains nearly unchanged
 276 regardless of whether SIC assimilation is included.

277 We also examine the performance of the long-term SIT forecasts at the BGEP sites
 278 (Figure S5). In general, significant improvements in the SIT forecasts are found in Long-
 279 SICSIT initialized in July, August and September of 2015. The differences between Long-
 280 SICSIT and Long-SIC in 2016 are limited, not exceeding 30 cm most of the time. The
 281 forecasts tend to overestimate SIT in the early freezing season in the Beaufort Sea. To check

282 if these biases are caused by the growing errors in the long-term atmospheric forecasts, we
 283 performed additional forecast experiments in 2015 with the same configuration as Long-
 284 CTRL, except that the CFSv2 atmospheric forecast is replaced by the ERA5 reanalysis for
 285 the atmospheric forcing. The ERA5 driven simulations show a similar overestimation of
 286 SIT in the Beaufort Sea (not shown). The anticyclonic wind in the Beaufort Gyre pushes
 287 excessively thick ice from the multi-year ice region north of the CAA into the Beaufort Sea.
 288 This suggests that the overestimation is not mainly due to biases in the atmospheric forcing
 289 but imperfect model parameterizations and initial ice-ocean conditions.

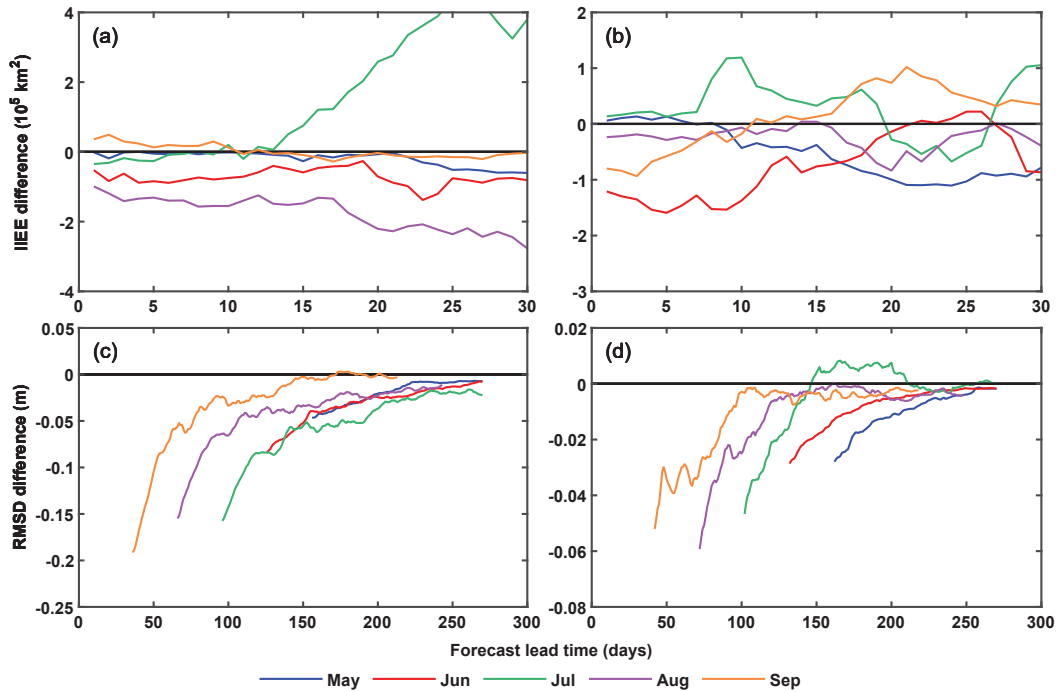


Figure 3. The difference of the IIEE (10^5 km^2) in 2015 (a) and in 2016 (b), and the difference of the SIT RMSD (m) in 2015 (c) and in 2016 (d) between Long-SICSIT and Long-SIC forecasts initialized from May to September (Long-SICSIT minus Long-SIC). The RMSD of the SIT is computed with respect to the CS2SMOS product available from October to April for (c) and (d). Negative values indicate better forecast skill. Note that different subfigures use different y-axis scales.

290 4 Summary

291 This study examines the impact of summer CryoSat-2 SIT assimilation on short- and
 292 long-term sea-ice forecasts in 2015 and in 2016. Compared to the experiments without
 293 any data assimilation, the ice-edge forecasts with summer CryoSat-2 SIT assimilation are
 294 improved. When the summer CryoSat-2 SIT data are assimilated together with SIC data,
 295 the effects on the ice-edge forecast skill are rather dependent on the time when the forecast is
 296 initialized and are spatially highly variable. In the Pacific sector, the combined assimilation
 297 of summer SIT and SIC observations leads to more realistic summer ice-edge forecasts with
 298 a one-week lead time.

299 The long-term sea-ice forecasts show significant reductions in both IIEE and RMSD
 300 of the SIT, except for those initialized in July, when the summer CryoSat-2 SIT has large
 301 uncertainties. The improvement in ice-edge forecasts can last up to about 30 days, while for

302 the SIT forecasts the benefits can last for more than 3 months. This result demonstrates
303 that, although the atmospheric forecasts used to drive the model can evolve freely after
304 about one month, the SIT initialization in summer remains a primary factor in predicting
305 long-term SIT variations. An extended study covering all available years of the CryoSat-2
306 dataset may concrete the conclusion.

307 However, limitations of the summer CryoSat-2 SIT data product still remain. The
308 deep learning algorithm used has a certain degree of uncertainty in classifying ice leads and
309 melt ponds, especially when the melt-pond fraction is large. The underestimation in the
310 summer CryoSat-2 SIT from July to September in the coastal regions north of the CAA and
311 Greenland requires further work on the sea-ice freeboard and thickness retrieval algorithm
312 or exploration of new correction schemes to improve their reliability and accuracy. Further-
313 more, it is still an open question how this product should be used for real-time Arctic sea-ice
314 forecasting, since its uncertainty currently does not account for all the algorithm errors, and
315 possible representation errors (Janjić et al., 2018) should be considered accurately.

316 5 Open Research

317 The ensemble mean Arctic sea-ice thickness (SIT) and sea-ice concentration (SIC) fore-
318 cast data used in the study can be downloaded at Song et al. (2024). The file size of
319 the forecast results with all ensemble members exceeds 50GB and can be made available
320 upon request through contact. The CMST SIT estimate is available at Mu et al. (2018a).
321 The summer CryoSat-2 SIT observations can be downloaded from Landy and Dawson
322 (2022). The SSMI/SSMIS SIC data is available from Kern et al. (2024). The UKMO
323 atmospheric ensemble forecasts are available in the THORPEX Interactive Grand Global
324 Ensemble (TIGGE; Bougeault et al., 2010) archive ([https://apps.ecmwf.int/datasets/
325 data/tigge](https://apps.ecmwf.int/datasets/data/tigge)). The hourly ERA5 reanalysis is available at Hersbach et al. (2023). The CFSv2
326 atmospheric forecasts are available at [https://www.ncei.noaa.gov/products/weather-
328 climate-models/climate-forecast-system](https://www.ncei.noaa.gov/products/weather-327-climate-models/climate-forecast-system). The NOAA/NSIDC SIC CDR data is avail-
329 able at Meier et al. (2021). The CS2SMOS data is available at [https://www.meereisportal
331 .de](https://www.meereisportal330.de). Mooring observations from BGEP are downloaded from [https://www2.whoi.edu/
site/beaufortgyre](https://www2.whoi.edu/site/beaufortgyre). The EASE-Grid Sea Ice Age, Version 4 (Tschudi et al., 2019) is
available at <https://nsidc.org/data/nsidc-0611/versions/4>.

332 Acknowledgments

333 We thank two anonymous reviewers for their constructive comments. We thank Thomas
334 Krumpfen for providing sea ice thickness observations from AWI IceBird campaigns. This
335 study is supported by the National Key R&D Program of China under Grant 2019YFA0607000,
336 the National Natural Science Foundation of China (42176235) and the computing resources
337 supported by the Laoshan Laboratory (LSKJ202202301, LSKJ202300303). Contribution of
338 Svetlana N. Loza was supported by the Federal Ministry of Education and Research of Ger-
339 many in the framework of the Seamless Sea Ice Prediction project (SSIP, Grant 01LN1701A)
340 and partly made in the framework of the state assignment of SIO RAS (theme FMWE-2024-
341 0028).

342 References

- 343 Blanchard-Wrigglesworth, E., Bitz, C. M., & Holland, M. M. (2011, 09). Influence of
344 initial conditions and climate forcing on predicting arctic sea ice. *Geophysical Research
345 Letters*, *38*, L18503. doi: 10.1029/2011GL048807
- 346 Blockley, E. W., & Peterson, K. A. (2018). Improving met office seasonal predictions
347 of arctic sea ice using assimilation of cryosat-2 thickness. *The Cryosphere*, *12*(11),
348 3419-3438. doi: 10.5194/tc-12-3419-2018
- 349 Bougeault, P., Toth, Z., Bishop, C., Brown, B., Burridge, D., Chen, D. H., ... Worley,

- 350 S. (2010). The thorpex interactive grand global ensemble. *Bulletin of the American*
351 *Meteorological Society*, *91*(8), 1059-1072. doi: 10.1175/2010BAMS2853.1
- 352 Bowler, N. E., Arribas, A., Mylne, K. R., Robertson, K. B., & Beare, S. E. (2008). The
353 mogreps short-range ensemble prediction system. *Quarterly Journal of the Royal*
354 *Meteorological Society*, *134*(632), 703-722. doi: 10.1002/qj.234
- 355 Bushuk, M., Msadek, R., Winton, M., Vecchi, G. A., Gudgel, R., Rosati, A., & Yang, X.
356 (2017). Skillful regional prediction of arctic sea ice on seasonal timescales. *Geophysical*
357 *Research Letters*, *44*(10), 4953-4964. doi: 10.1002/2017GL073155
- 358 Bushuk, M., Winton, M., Bonan, D. B., Blanchard-Wrigglesworth, E., & Delworth, T. L.
359 (2020). A mechanism for the arctic sea ice spring predictability barrier. *Geophysical*
360 *Research Letters*, *47*(13), e2020GL088335. doi: 10.1029/2020GL088335
- 361 Bushuk, M., Zhang, Y., Winton, M., Hurlin, B., Delworth, T., Lu, F., ... Zeng, F. (2022,
362 07). Mechanisms of regional arctic sea ice predictability in two dynamical seasonal
363 forecast systems. *Journal of Climate*, *35*, 4207-4231. doi: 10.1175/JCLI-D-21-0544.1
- 364 Comiso, J. C., Parkinson, C. L., Gersten, R., & Stock, L. (2008). Accelerated decline in
365 the arctic sea ice cover. *Geophysical Research Letters*, *35*(1), L01703. doi: 10.1029/
366 2007GL031972
- 367 Dawson, G., Landy, J., Tsamados, D. M., Komarov, A. S., Howell, S., Heorton, H., &
368 Krumpfen, T. (2022, 01). A 10-year record of arctic summer sea ice freeboard from
369 cryosat-2. *Remote Sensing of Environment*, *268*, 112744. doi: 10.1016/j.rse.2021
370 .112744
- 371 Day, J. J., Hawkins, E., & Tietsche, S. (2014). Will arctic sea ice thickness initialization
372 improve seasonal forecast skill? *Geophysical Research Letters*, *41*, 7566-7575. doi:
373 10.1002/2014GL061694
- 374 Day, J. J., Tietsche, S., & Hawkins, E. (2014). Pan-arctic and regional sea ice predictability:
375 initialization month dependence. *Journal of Climate*, *27*(12), 4371-4390. doi: 10.1175/
376 JCLI-D-13-00614.1
- 377 Dirkson, A., Merryfield, W. J., & Monahan, A. H. (2017). Impacts of sea ice thickness
378 initialization on seasonal arctic sea ice predictions. *Journal of Climate*, *30*, 1001-
379 1017. doi: 10.1175/JCLI-D-16-0437.1
- 380 Drinkwater, M. R. (1991). K_u band airborne radar altimeter observations of marginal sea
381 ice during the 1984 marginal ice zone experiment. *Journal of Geophysical Research:*
382 *Oceans*, *96*(C3), 4555-4572. doi: 10.1029/90JC01954
- 383 Feng, J., Zhang, Y., Cheng, Q., & Tsou, J. Y. (2022). Pan-arctic melt pond fraction trend,
384 variability, and contribution to sea ice changes. *Global and Planetary Change*, *217*,
385 103932. doi: 10.1016/j.gloplacha.2022.103932
- 386 Goessling, H. F., Tietsche, S., Day, J. J., Hawkins, E., & Jung, T. (2016). Predictability of
387 the arctic sea-ice edge. *Geophysical Research Letters*, *43*, 1642-1650. doi: 10.1002/
388 2015GL067232
- 389 Guemas, V., Blanchard-Wrigglesworth, E., Chevallier, M., Day, J. J., Déqué, M., Doblus-
390 Reyes, F. J., ... Tietsche, S. (2016). A review on arctic sea ice predictability and
391 prediction on seasonal-to-decadal timescales. *Quarterly Journal of the Royal Meteorological*
392 *Society*, *142*, 546-561. doi: 10.1002/qj.2401
- 393 Hebert, D. A., Allard, R. A., Metzger, E. J., Posey, P. G., Preller, R. H., Wallcraft, A. J.,
394 ... Smedstad, O. M. (2015, 11). Short-term sea ice forecasting: An assessment of ice
395 concentration and ice drift forecasts using the u.s. navy's arctic cap nowcast/forecast
396 system. *Journal of Geophysical Research: Oceans*, *120*, 8327-8345. doi: 10.1002/
397 2015JC011283
- 398 Hersbach, H., Bell, B., Berrisford, P., Biavati, G., Horányi, A., Muñoz Sabater, J., ...
399 Thépaut, J.-N. (2023). *Era5 hourly data on single levels from 1940 to present* [dataset].
400 Copernicus Climate Change Service (C3S) Climate Data Store (CDS). Retrieved from
401 [https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-
402 -single-levels?tab=overview](https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-single-levels?tab=overview) doi: 10.24381/cds.adbb2d47
- 403 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., ...
404 Thépaut, J.-N. (2020). The era5 global reanalysis. *Quarterly Journal of the Royal*

- 405 *Meteorological Society*, 1999–2049. doi: 10.1002/qj.3803
- 406 Hibler III, W. D. (1979). A dynamic thermodynamic sea ice model. *Journal of Physical*
 407 *Oceanography*, *9*, 815-846. doi: 10.1175/1520-0485(1979)009<0815:ADTSIM>2.0.CO;
 408 2
- 409 Janjić, T., Bormann, N., Bocquet, M., Carton, J. A., Cohn, S. E., Dance, S. L., ... Weston,
 410 P. (2018). On the representation error in data assimilation. *Quarterly Journal of the*
 411 *Royal Meteorological Society*, *144*(713), 1257-1278. doi: 10.1002/qj.3130
- 412 Jung, T., Gordon, N. D., Bauer, P., Bromwich, D. H., Chevallier, M., Day, J. J., ... Yang,
 413 Q. (2016). Advancing polar prediction capabilities on daily to seasonal time scales.
 414 *Bulletin of the American Meteorological Society*, *97*, 160113112747009. doi: 10.1175/
 415 BAMS-D-14-00246.1
- 416 Kaleschke, L., Lüpkes, C., Vilna, T., Haarpaintner, J., Borchert, A., Hartmann, J., &
 417 Heygster, G. (2001). Ssm/i sea ice remote sensing for mesoscale ocean-atmosphere
 418 interaction analysis. *Canadian Journal of Remote Sensing*, *27*, 526-537. doi: 10.1080/
 419 07038992.2001.10854892
- 420 Kern, S., Kaleschke, L., Girard-Ardhuin, F., Spreen, G., & Beitsch, A. (2024). *Global daily*
 421 *gridded 5-day median-filtered, gap-filled asi algorithm ssmi-ssmis sea ice concentration*
 422 *data* [dataset]. Integrated Climate Data Center. Retrieved from [https://www.cen.uni-](https://www.cen.uni-hamburg.de/en/icdc/data/cryosphere/seaiceconcentration-asi-ssmi.html)
 423 [hamburg.de/en/icdc/data/cryosphere/seaiceconcentration-asi-ssmi.html](https://www.cen.uni-hamburg.de/en/icdc/data/cryosphere/seaiceconcentration-asi-ssmi.html)
- 424 Kern, S., Kaleschke, L., & Spreen, G. (2010). Climatology of the nordic (irminger, greenland,
 425 barents, kara and white/pechora) seas ice cover based on 85 ghz satellite microwave
 426 radiometry: 1992–2008. *Tellus A*, *62*, 411-434. doi: 10.3402/tellusa.v62i4.15709
- 427 Kwok, R., & Rothrock, D. A. (2009). Decline in arctic sea ice thickness from submarine and
 428 icesat records: 1958-2008. *Geophysical Research Letters*, *36*, L15501. doi: 10.1029/
 429 2009GL039035
- 430 Landrum, L., & Holland, M. M. (2020, 12). Extremes become routine in an emerging new
 431 arctic. *Nature Climate Change*, *10*, 1-8. doi: 10.1038/s41558-020-0892-z
- 432 Landy, J. C., & Dawson, G. J. (2022). *Year-round arctic sea ice thickness from cryosat-*
 433 *2 baseline-d level 1b observations 2010-2020 (version 1.0)* [dataset]. NERC EDS UK
 434 Polar Data Centre. Retrieved from [https://data.bas.ac.uk/full-record.php?id=](https://data.bas.ac.uk/full-record.php?id=GB/NERC/BAS/PDC/01613)
 435 [GB/NERC/BAS/PDC/01613](https://data.bas.ac.uk/full-record.php?id=GB/NERC/BAS/PDC/01613) doi: 10.5285/d8c66670-57ad-44fc-8fef-942a46734ecb
- 436 Landy, J. C., Dawson, G. J., Tsamados, M., Bushuk, M., Stroeve, J. C., Howell, S. E. L.,
 437 ... Aksenov, Y. (2022). A year-round satellite sea-ice thickness record from cryosat-2.
 438 *Nature*, *609*, 1-6. doi: 10.1038/s41586-022-05058-5
- 439 Laxon, S. W., Giles, K. A., Ridout, A. L., Wingham, D. J., Willatt, R., Cullen, R., ...
 440 Davidson, M. (2013). Cryosat-2 estimates of arctic sea ice thickness and volume.
 441 *Geophysical Research Letters*, *40*(4), 732-737. doi: 10.1002/grl.50193
- 442 Lee, S., Kim, H.-C., & Im, J. (2018). Arctic lead detection using a waveform mixture
 443 algorithm from cryosat-2 data. *The Cryosphere*, *12*(5), 1665-1679. doi: 10.5194/
 444 tc-12-1665-2018
- 445 Lemieux, J.-F., Beaudoin, C., Dupont, F., Roy, F., Smith, G. C., Shlyaeva, A., ... Ferry, N.
 446 (2015, 03). The regional ice prediction system (rips): Verification of forecast sea ice
 447 concentration. *Quarterly Journal of the Royal Meteorological Society*, *142*, 632-643.
 448 doi: 10.1002/qj.2526
- 449 Losch, M., Menemenlis, D., Campin, J.-M., Heimbach, P., & Hill, C. (2010). On the
 450 formulation of sea-ice models. part 1: Effects of different solver implementations and
 451 parameterizations. *Ocean Modelling*, *33*(1), 129-144. doi: 10.1016/j.ocemod.2009.12
 452 .008
- 453 Marshall, J., Adcroft, A., Hill, C., Perelman, L., & Heisey, C. (1997). A finite-volume,
 454 incompressible navier stokes model for studies of the ocean on parallel computers.
 455 *Journal of Geophysical Research*, *102*, 5753-5766. doi: 10.1029/96JC02775
- 456 Maykut, G. A., Grenfell, T. C., & Weeks, W. (1992). On estimating spatial and temporal
 457 variations in the properties of ice in the polar oceans. *Journal of Marine Systems*, *3*,
 458 41-72. doi: 10.1016/0924-7963(92)90030-C
- 459 Meier, W. N., Fetterer, F., Windnagel, A. K., & Stewart, J. S. (2021). *Noaa/nsidc cli-*

- 460 *mate data record of passive microwave sea ice concentration, version 4* [dataset]. National
 461 Snow and Ice Data Center. Retrieved from [https://nsidc.org/data/G02202/](https://nsidc.org/data/G02202/versions/4)
 462 [versions/4](https://nsidc.org/data/G02202/versions/4) doi: 10.7265/efmz-2t65
- 463 Mignac, D., Martin, M., Fiedler, E., Blockley, E., & Fournier, N. (2022, 02). Improving the
 464 met office’s forecast ocean assimilation model (foam) with the assimilation of satellite-
 465 derived sea-ice thickness data from cryosat-2 and smos in the arctic. *Quarterly Journal*
 466 *of the Royal Meteorological Society*, *148*, 1-24. doi: 10.1002/qj.4252
- 467 Min, C., Yang, Q., Luo, H., Chen, D., Krumpfen, T., Mamnun, N., ... Nerger, L. (2023).
 468 Improving arctic sea-ice thickness estimates with the assimilation of cryosat-2 summer
 469 observations. *Ocean-Land-Atmosphere Research*, *2*, 0025. doi: 10.34133/olar.0025
- 470 Mu, L., Liang, X., Yang, Q., Liu, J., & Zheng, F. (2019). Arctic ice ocean prediction system:
 471 evaluating sea ice forecasts during xuelong’s first trans-arctic passage in summer 2017.
 472 *Journal of Glaciology*, 1-9. doi: 10.1017/jog.2019.55
- 473 Mu, L., Losch, M., Yang, Q., Ricker, R., Losa, S. N., & Nerger, L. (2018a). *The arc-*
 474 *tic combined model and satellite sea ice thickness (cmst) dataset* [dataset]. PAN-
 475 GAEA. Retrieved from <https://doi.org/10.1594/PANGAEA.891475> doi: 10.1594/
 476 PANGAEA.891475
- 477 Mu, L., Losch, M., Yang, Q., Ricker, R., Losa, S. N., & Nerger, L. (2018b). Arctic-
 478 wide sea ice thickness estimates from combining satellite remote sensing data and a
 479 dynamic ice-ocean model with data assimilation during the cryosat-2 period. *Journal*
 480 *of Geophysical Research: Oceans*, *123*, 7763-7780. doi: 10.1029/2018JC014316
- 481 Mu, L., Nerger, L., Streffing, J., Tang, Q., Niraula, B., Zampieri, L., ... Goessling,
 482 H. F. (2022). Sea-ice forecasts with an upgraded awi coupled prediction sys-
 483 tem. *Journal of Advances in Modeling Earth Systems*, *14*(12), e2022MS003176. doi:
 484 10.1029/2022MS003176
- 485 Nerger, L., Hibler III, W. D., & SCHRÖTER, J. (2005). Pdaf-the parallel data assimilation
 486 framework: Experiences with kalman filtering. *World Scientific*, 63–83. doi: 10.1142/
 487 9789812701831.0006
- 488 Nerger, L., Janjić, T., Schröter, J., & Hiller, W. (2012). A regulated localization scheme for
 489 ensemble-based kalman filters. *Quarterly Journal of the Royal Meteorological Society*,
 490 *138*(664), 802-812. doi: 10.1002/qj.945
- 491 Nguyen, A. T., Menemenlis, D., & Kwok, R. (2011). Arctic ice-ocean simulation with opti-
 492 mized model parameters: Approach and assessment. *Journal of Geophysical Research:*
 493 *Oceans*, *116*, C04025. doi: 10.1029/2010JC006573
- 494 Parkinson, C. L., & Washington, W. M. (1979). A large-scale numerical model of sea ice.
 495 *Journal of Geophysical Research*, *84*, 311-337. doi: 10.1029/JC084iC01p00311
- 496 Posey, P., Metzger, E., Wallcraft, A., Hebert, D., Allard, R., Smedstad, O., ... Helfrich,
 497 S. (2015, 08). Improving arctic sea ice edge forecasts by assimilating high horizontal
 498 resolution sea ice concentration data into the us navy’s ice forecast systems. *The*
 499 *Cryosphere*, *9*, 1735-1745. doi: 10.5194/tc-9-1735-2015
- 500 Ricker, R., Hendricks, S., Helm, V., Skourup, H., & Davidson, M. (2014). Sensitivity of
 501 cryosat-2 arctic sea-ice freeboard and thickness on radar-waveform interpretation. *The*
 502 *Cryosphere*, *8*, 1607-1622. doi: 10.5194/tc-8-1607-2014
- 503 Ricker, R., Hendricks, S., Kaleschke, L., Tian-Kunze, X., King, J., & Haas, C. (2017). A
 504 weekly arctic sea-ice thickness data record from merged cryosat-2 and smos satellite
 505 data. *The Cryosphere*, *11*(4), 1607–1623. doi: 10.5194/tc-11-1607-2017
- 506 Rothrock, D. A., Yu, Y., & Maykut, G. A. (1999). Thinning of the arctic sea-ice cover.
 507 *Geophysical Research Letters*, *26*, 3469-3472. doi: 10.1029/1999GL010863
- 508 Saha, S., Moorthi, S., Wu, X., Wang, J., Nadiga, S., Tripp, P., ... Becker, E. (2014).
 509 The ncep climate forecast system version 2. *Journal of Climate*, *27*, 2185–2208. doi:
 510 10.1175/JCLI-D-12-00823.1
- 511 Semtner, A. J. (1976). A model for thermodynamic growth of sea ice in numeri-
 512 cal investigations of climate. *Journal of Physical Oceanography*, *6*, 379-389. doi:
 513 10.1175/1520-0485(1976)006<0379:AMFTTG>2.0.CO;2
- 514 Shu, Q., Qiao, F., Liu, J., Song, Z., Chen, Z., Zhao, J., ... Song, Y. (2021). Arctic sea ice

- 515 concentration and thickness data assimilation in the fio-esm climate forecast system.
516 *Acta Oceanologica Sinica*, 40, 65-75. doi: 10.1007/s13131-021-1768-4
- 517 Song, R., Mu, L., Kauker, F., Loza, S., & Chen, X. (2024). *Forecast data for the paper:*
518 *"assimilating summer sea-ice thickness observations improves arctic sea-ice forecast"*
519 [dataset]. Zenodo. Retrieved from <https://doi.org/10.5281/zenodo.11352863> doi:
520 10.5281/zenodo.11352863
- 521 Spreen, G., Kaleschke, L., & Heygster, G. (2008). Sea ice remote sensing using amsr-e 89-ghz
522 channels. *Journal of Geophysical Research*, 113, C02S03. doi: 10.1029/2005JC003384
- 523 Stroeve, J. C., Serreze, M. C., Holland, M. M., Kay, J. E., Malanik, J., & Barrett, A. P.
524 (2012, 02). The arctic's rapidly shrinking sea ice cover: A research synthesis. *Climatic*
525 *Change*, 110, 1005-1027. doi: 10.1007/s10584-011-0101-1
- 526 Tilling, R., Ridout, A., & Shepherd, A. (2019). Assessing the impact of lead and floe
527 sampling on arctic sea ice thickness estimates from envisat and cryosat-2. *Journal of*
528 *Geophysical Research: Oceans*, 124, 7473-7485. doi: 10.1029/2019JC015232
- 529 Tschudi, M., Meier, W. N., J. S. Stewart, C. F., & Maslanik., J. (2019). *Ease-grid sea ice*
530 *age, version 4* [dataset]. NASA National Snow and Ice Data Center Distributed Active
531 Archive Center. Retrieved from <https://nsidc.org/data/NSIDC-0611/versions/4>
532 doi: 10.5067/UTAV7490FEPB
- 533 Xie, J., Counillon, F., Bertino, L., Tian-Kunze, X., & Kaleschke, L. (2016). Benefits of
534 assimilating thin sea ice thickness from smos into the topaz system. *The Cryosphere*,
535 10, 2745-2761. doi: 10.5194/tc-10-2745-2016
- 536 Xiong, C., & Ren, Y. (2023). Arctic sea ice melt pond fraction in 2000-2021 derived by
537 dynamic pixel spectral unmixing of modis images. *ISPRS Journal of Photogrammetry*
538 *and Remote Sensing*, 197, 181-198. doi: 10.1016/j.isprsjprs.2023.01.023
- 539 Yang, Q., Losa, S. N., Losch, M., Jung, T., & Nerger, L. (2015). The role of atmospheric
540 uncertainty in arctic summer sea ice data assimilation and prediction. *Quarterly*
541 *Journal of the Royal Meteorological Society*, 141, 2314-2323. doi: 10.1002/qj.2523
- 542 Yang, Q., Losa, S. N., Losch, M., Liu, J., Zhang, Z., Nerger, L., & Yang, H. (2015).
543 Assimilating summer sea-ice concentration into a coupled ice-ocean model using a
544 lseik filter. *Annals of Glaciology*, 56(69), 38-44. doi: 10.3189/2015AoG69A740
- 545 Zhang, J., & Hibler III, W. D. (1997). On an efficient numerical method for modeling sea ice
546 dynamics. *Journal of Geophysical Research*, 102, 8691-8702. doi: 10.1029/96JC03744
- 547 Zhang, Y.-F., Bushuk, M., Winton, M., Hurlin, B., Gregory, W., Landy, J., & Jia, L.
548 (2023). Improvements in september arctic sea ice predictions via assimilation of sum-
549 mer cryosat-2 sea ice thickness observations. *Geophysical Research Letters*, 50(24),
550 e2023GL105672. doi: 10.1029/2023GL105672

Figure 1.

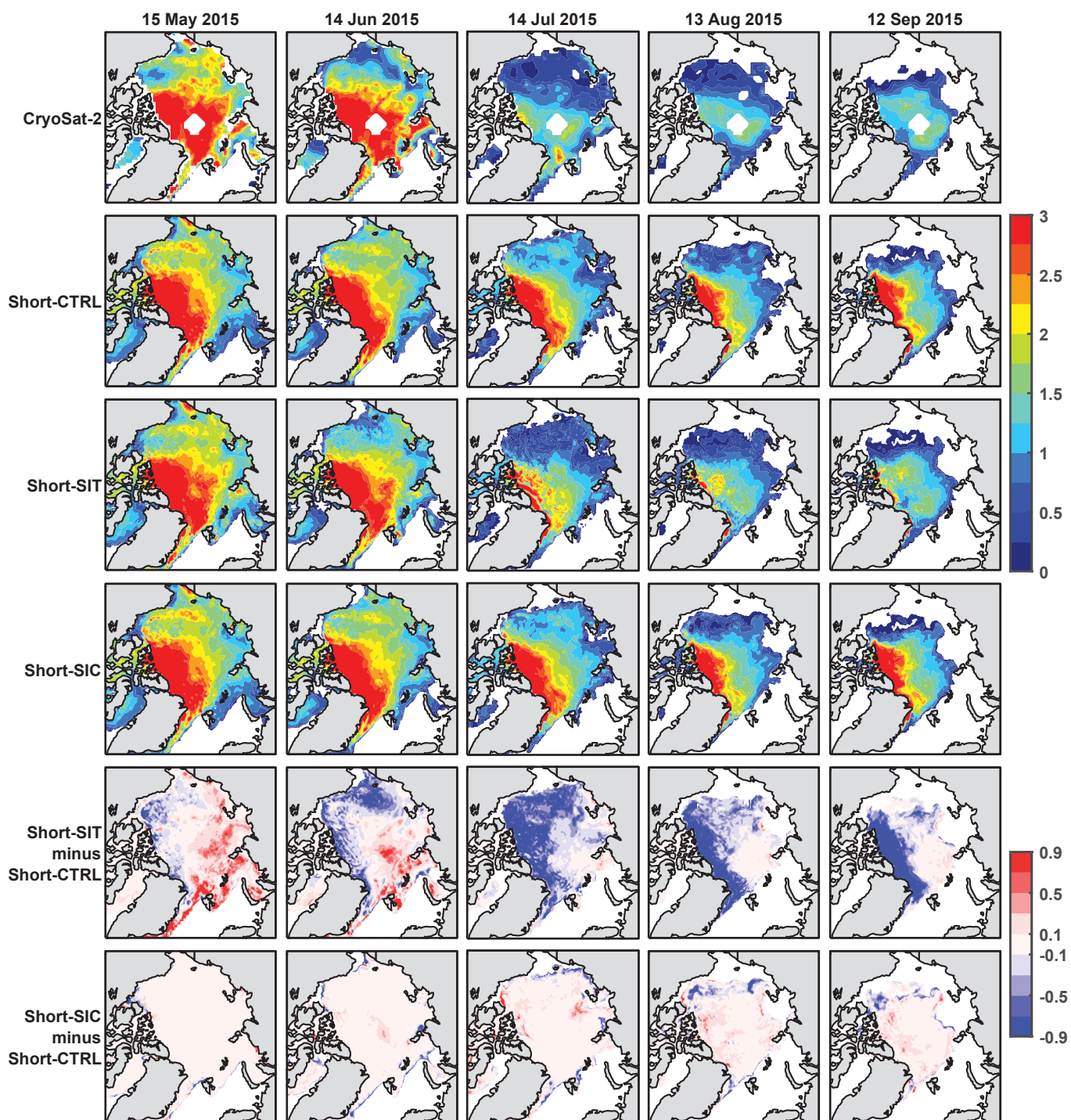


Figure 2.

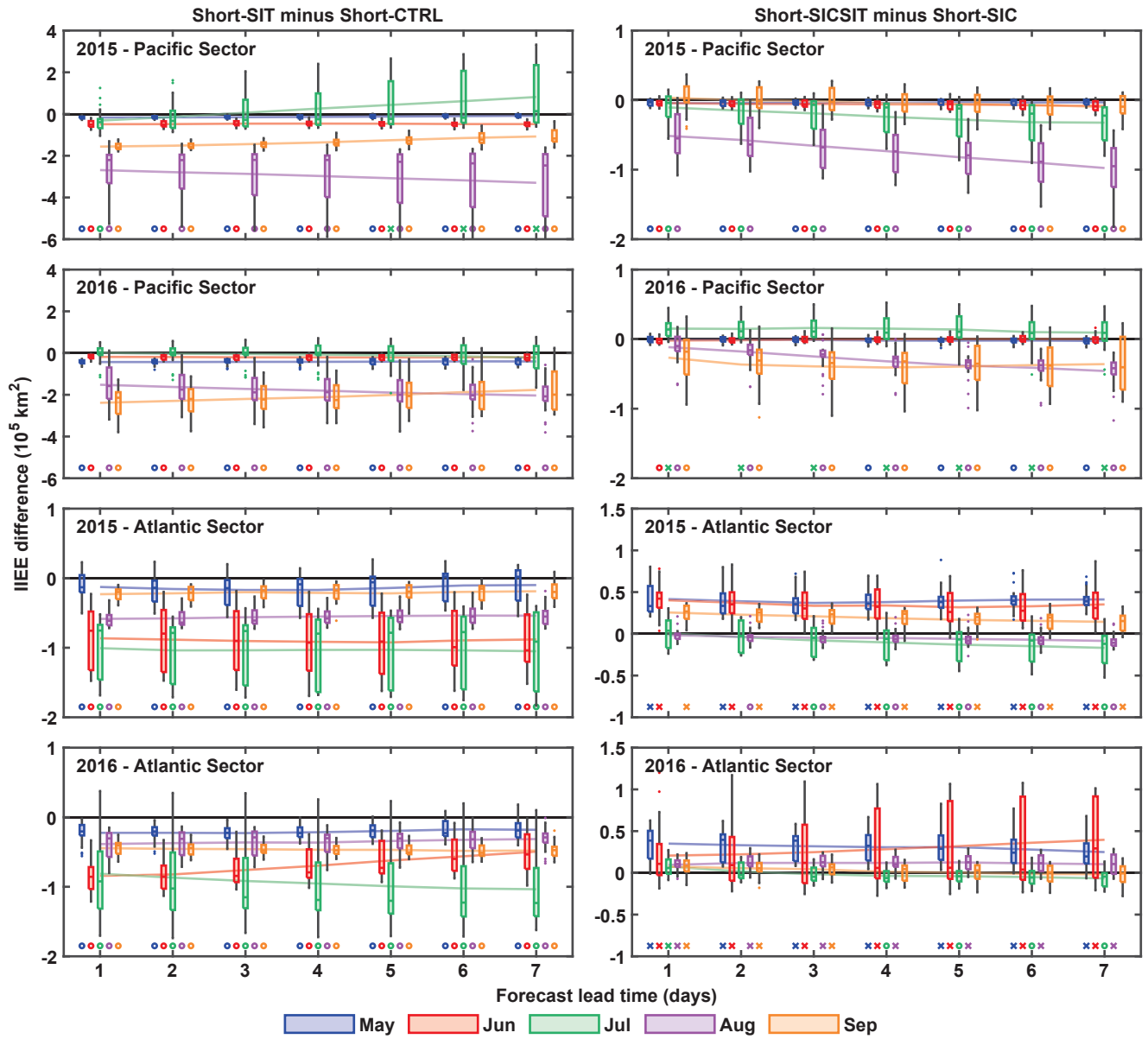


Figure 3.

