# A quantitative analysis of the reactions involved in stratospheric polar ozone depletion

Ingo Wohltmann, Ralph Lehmann, Markus Rex

Alfred Wegener Institute for Polar and Marine Research, Potsdam

# **Introduction**

- $\rightarrow$  Quantitative analysis of chemical reactions involved in stratospheric polar ozone depletion
- $\rightarrow$  Quantitative estimates of importance of single reactions or reaction cycles are rare. No comprehensive and quantitative study of the reaction rates averaged over polar vortex under conditions of heterogeneous chemistry so far.
- $\triangleright$  Mixing ratios and reaction rates obtained from ATLAS Chemistry and Transport Model driven by ECMWF ERA Interim reanalysis data.
- $\triangleright$  One Arctic winter (2004/2005) and one Antarctic winter (2006) in a layer in the lower stratosphere around 54 hPa.

# Method

- ATLAS is a global Lagrangian (trajectory-based) Chemistry and Transport Model including gas phase stratospheric chemistry, heterogenous chemistry on PSCs and a particle-based denitrification module. Includes 47 chemical species and more than 180 reactions.
- ▶ Vortex edge assumed at 36 PVU modified PV. Average only over air parcels with less than 30% extra-vortex air mixed in (do not mix up different chemical regimes).
- ▶ Reaction rates are changes over 24 h (one diurnal cycle)
- $\triangleright$  Daytime averages: less than 80° solar zenith angle, nighttime averages: more than 100° solar zenith angle.
- $\triangleright$  Days without sufficient data for averaging are not shown (grey bars).



Vortex averaged partitioning of inorganic chlorine species Cly. Left: Arctic, Right: Antarctic, Top: Day, Bottom: Night.



Vortex averaged partitioning of NO<sub>x</sub>. Left: Arctic, Right: Antarctic, Top: Day, Bottom: Night.

Vortex averaged reaction rates of reactions involving HCl. Green line: Net change of HCl. Left: Arctic, Right: Antarctic. Deactivation into HCl due to low ozone and denitrified conditions in SH and change in activation from  $CIONO<sub>2</sub> + HCI$  in early winter to  $HOCI + HCI$  later in SH visible.

### ClONO2 production and loss reactions



Vortex averaged reaction rates of reactions involving CIONO<sub>2</sub>. Green line: Net change of CIONO<sub>2</sub>. Left: Arctic, Right: Antarctic. Deactivation into  $CIONO<sub>2</sub>$  in NH when  $NO<sub>x</sub>$  is produced by shift in  $CIONO<sub>2</sub>$  equilibrium in NH, later  $CIONO<sub>2</sub>$  depleted to form HCl.

#### **Ozone** Catalytic ozone depletion cycles −60 −50 −40 −30 −20 −10 0 10 NH 2005 54 hPa [ppb/day] ClO−ClO cycle (Cl2O2+hv) BrO−CIO cycle (BrO+ClO O cycle with Cl (O+ClO) cycle with  $NO (O+NO)$ −60 −50 −40 −30 −20 −10 0 10 SH 2006 54 hPa [ppb/day]

Vortex averaged reaction rates of reactions changing (an extended)  $HO_{x}$  (OH + HO<sub>2</sub> + H + HOCl + HOBr + HO<sub>2</sub>NO<sub>2</sub>). Green line: Net change of extended  $HO_x$ . Left: Arctic, Right: Antarctic.



Vortex averaged partitioning of NOy. Left: Arctic, Right: Antarctic. Change is dominated by denitrification.

> Vortex averaged partitioning of  $HO_x$  (only day, no  $HO_x$  at night). Left: Arctic, Right: Antarctic.

### NOx production and loss reactions



Vortex averaged reaction rates of reactions changing (an extended)  $NO_{x}$  (NO + NO<sub>2</sub> + NO<sub>3</sub> + 2 N<sub>2</sub>O<sub>5</sub> + CIONO<sub>2</sub> + BrONO<sub>2</sub> +  $HO_2NO_2$ ). Green line: Net change of extended  $NO_x$ . Left: Arctic, Right: Antarctic.

### Partitioning NOx









Vortex averaged reaction rates of reactions changing NO to illustrate NO<sup>x</sup> partitioning. Left: Arctic, Right: Antarctic.





Vortex averaged partitioning of inorganic bromine species Bry. Left: Arctic, Right: Antarctic, Top: Day, Bottom: Night.

- $HOCl + HCl$  accounts for 70% of activation of HCl in SH, and for 30% of activation in NH.
- $\triangleright$  CIO<sub>x</sub> peaks at 2.0–2.5 ppb. 70% of CIO<sub>x</sub> present as CIO during daytime.
- $H_4$  HO<sub>x</sub> peaks at 4 ppt. HO<sub>x</sub> mainly produced from CH<sub>4</sub> oxidation in SH. In NH, production by  $HNO<sub>3</sub>$ , CH<sub>4</sub> and H<sub>2</sub>O comparable. Partitioning  $OH/HO<sub>2</sub>$  results in 20%–40% OH in SH and in 10%–20% OH in NH.
- $\triangleright$  NO<sub>x</sub> smaller than 2 ppb in NH and smaller than 0.75 ppb in SH due to denitrified conditions there. Partitioning  $NO/NO<sub>2</sub>$  during daytime 80%–90% NO in SH and 20%–40% NO in the NH. Higher NO caused by much lower ozone in SH.
- $\sim 60\%-80\%$  of NO<sub>x</sub> production in spring caused by the HNO<sub>3</sub> + OH, remainder caused by  $HNO<sub>3</sub> + h\nu$ . Deactivation of  $ClO<sub>x</sub>$  by formation of  $CIONO<sub>2</sub>$  in NH by shift in equilibrium between CIO,  $NO<sub>2</sub>$  and  $CIONO<sub>2</sub>$ , which in turn is caused by the production of  $NO<sub>x</sub>$ .

−70

O cycle with OH (O+HO2)

1 May 1 Jun 1 Jul 1 Aug 1 Sep 1 Oct 1 Nov

−70

Vortex averaged net loss rates of odd oxygen by different catalytic cycles. Red line: Net loss rate of ozone. Contribution of different cycles shown by reaction rates of rate limiting step. Left: Arctic, Right: Antarctic.

# Hydrogen

### HOx production and loss reactions

1 Dec 1 Jan 1 Feb 1 Mar



### Partitioning HOx



### HOx partitioning reactions



Vortex averaged reaction rates of reactions changing  $HO<sub>2</sub>$  to illustrate HO<sub>x</sub> partitioning. Left: Arctic, Right: Antarctic.

## Some selected findings

- $\triangleright$  CIO dimer cycle contributes 50% to vortex averaged ozone loss (54 hPa) in both hemispheres. CIO–BrO cycle contributes 40%.
- $\blacktriangleright$  In southern hemisphere, clear shift from chlorine activation by  $CIONO<sub>2</sub> + HCI$  in early winter to  $HOCI + HCI$  later in winter.