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Your complimentary te upon the hydrolgy II polar glacier bre (79°N, Svalbard)







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The Lovén FLOWS program will be conducted over a 4-years period including IPY.

The chiective is to set up a real scale budged above to se The objective is to set up a real scale hydrological observatory of a representative polar glacier basin of Spitsbergen, by a survey in continuous of both (i) the fluxes and their spatial and temporal dynamics and (ii) some meteorological parameters (T°C, Pmm). One of the originality of this project is the setting up of a real in situ sensing system based on digital cameras to survey the basin.

The AustreLovén glacier basin(10 km2) has been selected to be an observatory of streamflow dynamics in western Spitsbergen. The se lection is based upon 4 criteria:

1. It is representative of the small continental glaciers of western Spitsbergen;



2. The outflows of the basin are concentrated into two well-defined rivers crossing a calcareous bar in twocanyons. This situation exceptional in polar regions allows the measurement of the total out-

3. The glacier displays a rapid response to recent



The Loven East glacier has retreated extensive especially during the last sixty years due global warming.

4. The glacier shows a high decrease in thickness in lower part while its thickness displays few variation at high altitude

water

and

• the hydrological way (rainfall, snow, flow

budget a

mass balance is

studied by 2 ap-

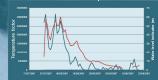


Hydrological data

area The water level of the outlet of the Austre Loven glacier catchment is more dependant upon air temperature conditions (A) than upon rainfall or snow fall amount (B)



However the water level curve differs to that of air temperature (T factor is an indicated of air T integrated over the whole basin)





Previous studies have shown a very important fluvioglacial dynamics, characterised by a high discontinuity of runoff in space



The in situ sensing is helpful to follow the fluvio-glacial dynamics in space and time, at different scales (km2, m2, dm2):

- organisation of the runoff.
- constitution and evolution of the icing and the subglacial flow,



dynamics of flows and hydro-morphologic evolution in the moraine

- front retreat,
- closing up of the subglacial channel,
- snow distibution



The runoff water during summer is rather diluted due to a high proportion of melting water.

At the beginning of winter, the

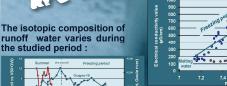
and ation of water

increases progressively independantly of T or P.

summer, runoff water gets low mineralisation (less than 200 μS/cm) with variable pН values (7-8).

During the beginning of winter (freezing period), the chemical composition of water evolves progres-sively between 2 end-

members : (1) melting water and (2) sub-glacial water (pH of 8 and EC of 800 μS/cm).





Summer: the enrichment in O-18 ind water changes through the thawiur Winter: the increasing proportion of sub-glacial water explains the depletion in O-18 of water.