

# A STUDY OF GLACIAL-INTERGLACIAL VARIATIONS OF THE MARINE STABLE CARBON ISOTOPE RECORD USING A NON-REDFIELD BIOGEOCHEMICAL MODEL



**PAL MOD**

GERMAN CLIMATE MODELING INITIATIVE



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**Martin Butzin, Peter Köhler, Christoph Völker**  
**Alfred-Wegener-Institut Helmholtz Zentrum für Polar- und Meeresforschung, Bremerhaven**  
**Martin.Butzin@awi.de**

## Motivation

Photosynthetic plankton has a lower  $^{13}\text{C}$  concentration than the aqueous  $\text{CO}_2$  from which it was formed. The isotopic discrimination during photosynthesis increases the  $^{13}\text{C}:^{12}\text{C}$  ratio (or  $\delta^{13}\text{C}$ ) of dissolved inorganic carbon (DIC). Here, we investigate the effect of two different parametrizations of biogenic fractionation on the carbon-isotopic signature of DIC under present and glacial climate conditions.

## Biogenic carbon fractionation

We consider two parametrizations of biogenic fractionation. Rau (1994) found that the isotopic depletion of phytoplankton  $\delta^{13}\text{C}_p$  increases with the availability of aqueous carbon dioxide  $\text{CO}_2^*$ :

$$\delta^{13}\text{C}_p = -a_0\text{CO}_2^* - b_0$$

Laws et al. (1997) found that  $\delta^{13}\text{C}_p$  also depends on the isotopic composition of  $\text{CO}_2^*$  (depending on temperature and  $\text{CO}_3^{2-}$ ) and photosynthesis  $\mu$ :

$$\delta^{13}\text{C}_p = \delta^{13}\text{C}\text{CO}_2^* - (a_1\text{CO}_2^* + b_1\mu) / (a_2\text{CO}_2^* + b_2\mu)$$

## Conclusion / Summary / Outlook

Different parametrizations of biogenic fractionation lead to discernable changes in the carbon-isotopic composition of DIC. The differences are seen in the entire water column. This may be particularly an issue in future model - data comparisons for the glacial ocean.

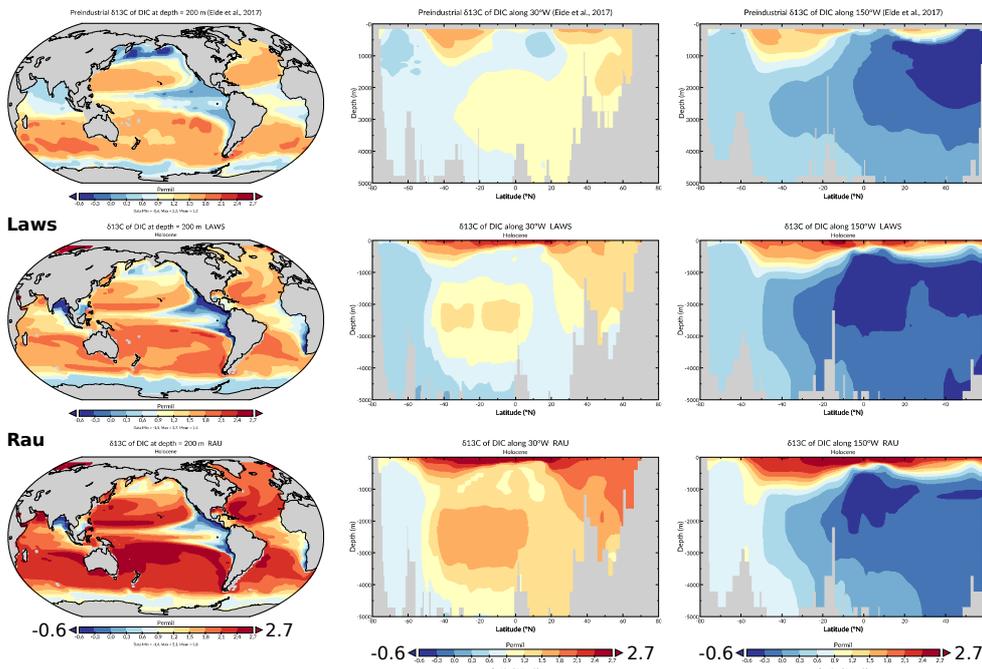
## Holocene

### Reconstruction

200 m

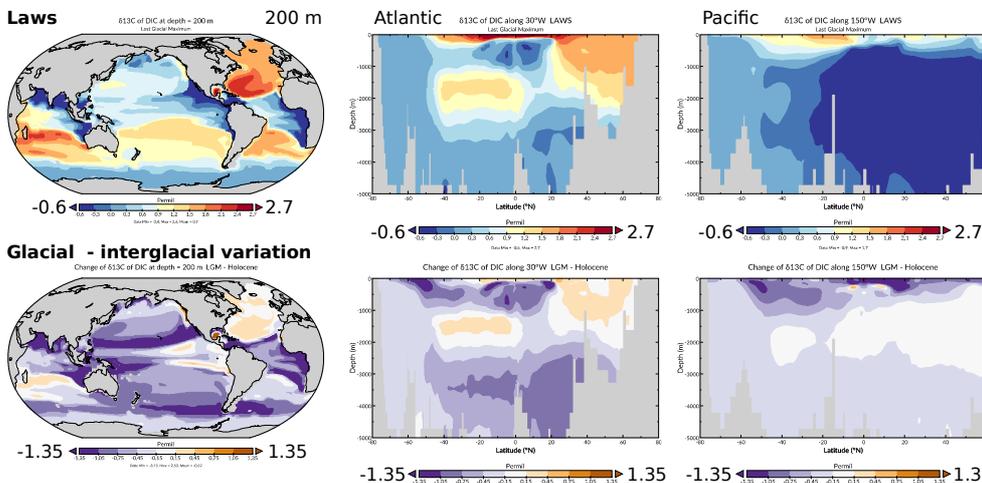
### Atlantic

### Pacific



The biogenic fractionation parametrization according to Rau leads to higher values of  $\delta^{13}\text{C}$  of DIC than the parametrization by Laws et al. The elevation is not limited to the euphotic zone but is also obvious in deeper waters. In our Holocene simulations the parametrization according to Laws et al. leads to better agreement with the reconstruction by Eide et al.

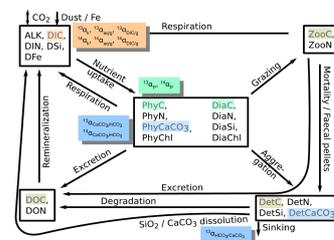
## Last Glacial Maximum



We find isotopic depletion of DIC except for small areas such as the subtropical North Atlantic and the Indian.

## Model setup

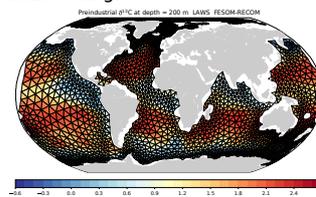
We use the Regulated Ecosystem Model RECOM which does not rely on fixed Redfield ratios for organic soft tissue. Instead, the C:N and C:Chl ratios in phytoplankton respond to light, temperature and nutrient supply. We consider isotopic fractionation during uptake and dissolution of  $\text{CO}_2$ , calcification, and photosynthesis of phytoplankton, plus radioactive decay of  $^{14}\text{C}$ .



Tracers are transported via the MITgcm, forced with climatological fields derived in fully coupled climate simulations for the Holocene (Wei & Lohmann, 2012) and the Last Glacial Maximum (Zhang et al., 2013). Our MITgcm-RECOM simulations build upon previous work by Völker & Köhler (2013). Dust fields are by Albani et al. (2016). Model resolution is  $2^\circ$  longitude x  $0.38^\circ$  to  $2.0^\circ$  latitude x 30 levels. Integration time (so far): 2 - 3 kyr.

## Current model development

RECOM has become part of the most recent version of the AWI climate model. The AWI climate model features FESOM, a multi-resolution sea ice-ocean model solving the equations of motion on unstructured meshes. The figure below gives an early impression of a test run with prescribed preindustrial climate forcing.



## References

Albani S et al. (2016), doi:10.1002/2016GL067911  
 Eide M et al. (2017), doi:10.1002/2016GB005473  
 Laws, EA et al. (1997), doi:10.4319/lo.1997.42.7.1552  
 Rau GH (1994), in: Zahn R et al. (eds), Carbon cycling in the glacial ocean: constraints on the ocean's role in global climate change Springer, Berlin, p 307-322  
 Völker C & Köhler P (2013), doi:10.1002/2013PA002556  
 Wei W & Lohmann G (2012), doi:10.1175/JCLI-D-11-00667.1  
 Zhang X et al. (2013), doi:10.5194/cp-9-2319-2013

